

REFINING CULTURAL AND ENVIRONMENTAL TEMPORALITIES AT THE
LATE ARCHAIC–EARLY WOODLAND TRANSITION
ON THE GEORGIA COAST, USA

by

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(Under the Direction of Victor D. Thompson)

ABSTRACT

The Late Archaic in the Southeast U.S. (c. 4500–3100 B.P.) was a time of increased sedentism and social complexity, long-distance trade, and relatively large-scale societies. Along the coast, societies constructed monumental shell rings that functioned as persistent places on the landscape and the focal points of village sites that were occupied year-round. Estuaries provided bountiful resources for these villages, including the oysters whose shells comprise the majority of the ring sites. In the terminal Late Archaic, however, a major cultural shift occurred. Shell ring sites were depopulated, and groups became smaller and more mobile. Evidence from various world regions points to a global climatic transition occurring in the mid-Holocene whose regional timing and manifestations varied. On the Georgia Coast, evidence has long pointed to the localized effects of this climatic shift as including one or more sea-level shifts. In this examination of culture and environment at the end of the Late Archaic, we model new and existing radiocarbon dates from shell-bearing archaeological sites and interpret use and depopulation dates within the framework of a high-resolution paleoenvironmental proxy

data source – a 5,177-year tree-ring chronology, a site-level standardized representation of annual tree growth, (and an earlier 529-year floating chronology) derived from ancient buried bald cypress (*Taxodium distichum*) trees at the mouth of the Altamaha River that provides insight into annual environmental conditions on the North Georgia Coast. This chronology and associated analyses indicate that Late Archaic societies weathered a 493-year period of enhanced environmental fluctuation that lasted from 2355 – 1863 B.C.E. and included salinity intrusion events, unreliable rainfall patterns, numerous drought years, and possible hurricanes, speaking to the resilience of both Native American societies of the ancient Southeast as well as the estuary ecosystems that provided their resource base.

INDEX WORDS: Coastal Archaeology; Environment; Late Archaic; Southeast U.S.;
Paleoclimate; Tree Rings; Shell Rings; Radiocarbon;
Mid-Holocene

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DEDICATION

Dedicated to my wonderful parents, Angela Presson and Casimir Napora.

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CHAPTER 1

INTRODUCTION

In the twenty-first century, societies around the globe grapple with questions of cultural resilience and sustainability in the face of increasing climatic instability. People living on coastlines are being forced to consider the ramifications of many facets of climate change, including rising sea levels, increasing storm activity, temperatures trending upwards, and enhanced environmental fluctuation. Confronted with this barrage of environmental blows, will we be able to adapt? Or, will we find it necessary to reconfigure our ways of life?

These considerations have special urgency today, as the global population hits 7.8 billion, and human-exacerbated climate change continues to accelerate. One third of the global population lives within sixty miles of a coastline (Nagaraja 2020). As we look towards an uncertain climatic future, the value of the information that can be gleaned from ancient societies that experienced comparable environmental conditions, and perhaps employed to help bolster twenty-first century resilience, becomes clear.

Four thousand years ago on the Georgia Coast, Indigenous societies flourished. Monumental shell rings functioned as persistent places on the landscape, representing centuries-long continuity of lifeways focused on the estuaries between the barrier islands and the mainland. The rich bounty of these estuaries provided reliable resources for the first settled villages in North America. The villages at the shell rings, however, were no longer inhabited after the Terminal Late Archaic, and the data indicate possibly more mobile societies as the norm after this period. Did environmental aspects play a role in

this cultural shift? Were these coastal hunter-gatherer societies resilient in the face of major environmental upheavals? Here, I examine this question via an integrated study combining paleoenvironmental proxies with direct radiocarbon dating and modeling of archaeological sites.

The second chapter of this dissertation introduces the theoretical and methodological frameworks by which this study analyzes cultural resilience in periods of environmental upheaval. I also discuss the site history of Butler Island in McIntosh County, Georgia, a former plantation site from which the vast majority of the ancient trees derive as well as the fieldwork undertaken for this project.

In the third chapter, I present the new 5,177-year tree-ring chronology for the Georgia Coast developed from a deposit of ancient bald cypress (*Taxodium distichum*) trees buried at the mouth of the Altamaha River as well as an earlier 529-year floating segment. I discuss the importance of this chronology to high-resolution paleoenvironmental and archaeological analyses in the Southeast U.S.

The fourth chapter considers the environmental drivers of the tree-ring chronology and discusses the results of the Optical Emission Spectroscopy (OES) elemental tests on trees, tree life histories, and environmental fluctuation and downturns visible in the multimillennial chronology. I compare tree ring-based insights into salinity intrusion events with existing sea-level curves for the Southeast Coast.

The fifth chapter delves into the use-lives and depopulations of the Late Archaic shell rings of the North Georgia Coast, modeling previously existing and new radiocarbon dates and considering them in light of the tree-ring chronology.

Finally, I conclude this dissertation with a summary of the main findings and a brief discussion of the relevance of this work to our own societal resilience in the twenty-first century.

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Nagaraja, Mamta Patel. 2020. Living Ocean. NASA Science: Share the Science. <https://science.nasa.gov/earth-science/oceanography/living-ocean>. Accessed November 9, 2020.

CHAPTER 2

LITERATURE REVIEW

Are coastal hunter-gatherer societies resilient in the face of major environmental change? How long after the onset of ecological and climatological disruption can estuarine-dependent societies continue their lifeways? This dissertation investigates the answers to these questions by examining the continuity of and shifts in littoral lifeways on the Georgia Coast during the transition from the Late Archaic (c. 4500–3100 B.P.) to the Early Woodland (c. 3100–2400 B.P.). This study frames the archaeological and paleoenvironmental evidence for temporally variable cultural transitions during this period within a new environmental proxy—a 5177-year tree-ring chronology (i.e., a site-level representation of tree growth (Speer 2010: 4)) spanning the years 3161 B.C.E. to 2016 C.E., as well as an earlier floating (i.e., unconnected to the present) 529-year tree-ring chronology we have developed for the Georgia Coast. This background chapter discusses the theoretical framework for analyses, the cultural and environmental context of the study, dendrochronology as it has been employed in the study area, the paleoenvironmental field site, and the methods employed in this research.

Beyond Collapse: Resilience Theory and Archaeology

In recent years, archaeological research focused on major cultural change has moved beyond the idea of societal collapse, instead analyzing the temporalities and interactions, both cultural and environmental, that contribute to differing occurrences and experiences of societal change. Rather than a singular linear and overarching shift as explanation, diversities and nuances (e.g., geographic, temporal, etc.) are now recognized

as crucial to the conceptualization of cultural change. A major approach to understanding transformative change in adaptive cultural-ecological systems is resilience theory, at its base conceptualized by the ‘figure eight’ adaptive ecosystem cycle model composed of growth, conservation, release, and reorganization (Hollig and Gunderson 2002). Rather than the loss implied by the term ‘collapse,’ resilience instead focuses on variation through time in complexity and form as well as continuity (Faulseit 2015; Tainter 2015). At the same time as it informs studies of the ancient past, resilience theory also holds the potential to elucidate and contribute to contemporary issues such as the ramifications of increasingly dramatic climate change (Redman 2005).

Resilience theory has recently been applied to the case of the Georgia coast at the time of the Archaic-Woodland transition. Turck and Thompson (2016; see also Thompson and Turck 2009) analyzed what they termed differential community resilience, the idea that communities in periods of transition may experience, understand, and respond to change in diverse ways, based on varying societal and environmental parameters. On the terminal Late Archaic Georgia coast, differential resilience appears to have been tied to the presence or absence of deltaic sediment deposition, which may have buffered river deltas against the deleterious effects of a drop in sea level. Russo (2010) also noted differential community resilience along the Florida coastline during the Late Archaic-Woodland transition, emphasizing that the interplay of social factors and localized environmental conditions led to a highly diverse cultural landscape during this period. In southern Florida, Russo states that shell ring formation continued through the Late Archaic-Woodland transition, probably supported by still-productive estuaries of the gradually sloped coastline. By contrast, the steeper drop-off of the continental shelf in

northeast Florida may have disrupted estuary productivity (Russo 2010:161, 171).

However, even in these more trying environmental conditions, sedentary lifeways, public mound building, and differentiation in social status continued, indicating that cultural resilience is not entirely dependent on environmental parameters. The analyses of both the Georgia and Florida coastlines indicate the importance of multiscale analyses to resilience theory. While a general trend may seem evident over a large area, case studies can provide invaluable insights into different experiences of and approaches to resilience.

Resilience and Scale

Multiscale approaches and cross-scale interactions are central to resilience studies in archaeology (Redman and Kinzig 2003). The adaptive cycles central to a resilience approach may be viewed as complex nested entities within a broader spatial and temporal framework (Redman and Kinzig 2003:1). Interactions across scales of space and time are therefore crucial to the understanding of resilience. Deep time approaches, which strive to understand the causes of change and transformation over very long time scales, have been put forth as a central element that archaeological research has the ability to contribute to the study of resilience (Redman and Kinzig 2003). By contrast, however, small-scale analyses are also vital to the study of resilience, as discussed in the Archaic-Woodland transition examples from the Georgia and Florida coasts. This study delves into the short-term temporal diversity and the localized nuances of a particular event—the cultural transition from the Archaic to the Woodland and associated environmental shifts— as visualized within what Simiand (1960) termed *histoire événementielle*, the conceptualization of time as punctuated by episodes of change. The critique that analysis centered on such a perspective overlooks the importance of the

“slow, sometimes motionless currents” of deep social structures in favor of a focus on the “surface flurries” of “short-term crises” (Braudel 1949: 1241-1242), however, is addressed here by situating ‘the event’ of focus within a multimillennial timescale that allows for the identification of short-term anomalies as well as more enduring state shifts. Long- and short-term conditions on the Georgia Coast are conceptualized in this study within a press-pulse dynamics framework, which aids in defining and identifying normative and deviant conditions for the locale. This project elucidates long-term trends and changes in decadal and yearly conditions at the mouth of the Altamaha River, providing a multiscalar dataset on the ancient coastal environment and employing a pulse-press dynamic framework to illuminate human responses to long- and short-term events.

The incorporation of tree-ring data into a study of differential resilience allows a unique opportunity to focus on a variant of the microscale usually invisible in the archaeological record—that of a year-by-year analysis—while conceptualizing this zoomed-in approach within the overarching framework of a multimillennial *histoire événementielle*, with both deep currents and surface movements illuminated. Tree rings provide one of the only annual, and positively dateable, records of locale ecological conditions available. Ringwidth chronologies, meanwhile, present this annual data as a temporally macroscale analysis. This bald cypress tree-ring study has developed a record of annual environmental conditions along the Georgia coastline extending back to 5161 B.P., before the time of the Archaic-Woodland transition. This ringwidth sequence is by far the longest tree-ring chronology on the East Coast and provides year-by-year information about ecological and climatic press-pulse dynamics.

Cultural and Environmental Context

This project examines human-environmental relationships before, during, and after the depopulation of the Late Archaic shell rings of the Georgia coast. By examining ancient remnant trees buried near the mouth of the Altamaha River, this study is able to delve into climate and environment in the region at a much finer scale than has previously been possible for this period of major cultural change.

Coastal Georgia's Terminal Late Archaic Period and the Transition to the Early Woodland

The Late Archaic is generally described as a period characterized by population growth, shifts in subsistence strategies, technological innovations, and increased social complexity (Anderson and Sassaman 2012; Jefferies 2008; Sassaman 2010). The societies of this period appear to have been linked by long-distance trade routes (see, for example Sanger et al. 2019) and to have shared certain cultural aspects, including the construction of monumental architecture and burial rituals (Kidder 2010: 24). Along the southeastern Atlantic Coast, shell rings are among the most well-known Late Archaic period sites. Geoarchaeological, faunal, and isotopic evidence indicates that these shell rings represent habitation sites utilized year-round by sedentary people (Colaninno 2012; Thompson and Andrus 2011; Turck and Thompson 2016). Shell ring sites also appear to have served as 'persistent places' on the Late Archaic coastal landscape, promoting continuity of some cultural practices with repeated reuse (Thompson 2010: 218). Most Atlantic coast shell rings appear to have been occupied for 500 years or less (Thompson 2010: 222). Between c. 3800 and 3500 B.P., during the terminal Late Archaic period,

occupation of the Georgia coast shell rings ceased, as did a general exploitation of shellfish species in some areas (Thompson and Turck 2009; Turck and Thompson 2016).

Contemporary research on the Late Archaic- Early Woodland transition in the southeastern U.S. has moved beyond an overly-broad categorization of the period as one of general decline—decline in population density, long-distance networks and connectivity, and measures of cultural complexity—instead focusing on the considerable variability of cultural trajectories of this era (e.g., Kidder 2010; Marquardt 2010; Turck and Thompson 2016). Many societies, however, did experience dramatic changes, some of which appear to have been accompanied by major climatic events.

The Mid- to Late-Holocene Climate Change Event

Around 3000–2600 B.P., a major climatological event affected different regions of the globe. The cause of this climate episode appears to have been a change in cosmic ray intensity and solar irradiation (Dergachev et al. 2004; van Geel et al. 2000).

Variability within the localized manifestations of this climatic episode most likely relate to differential solar forcing and variation in the planet's geomagnetic field and variations in underlying atmospheric and oceanic circulation patterns between regions (Debret et al. 2009; Kidder 2006: 212; Turney et al. 2005).

Kidder's Climate Hypothesis (2006, 2010:25) compiles evidence for the decline of much of the southeastern Late Archaic being entangled with the global changes in climate that appear to have altered weather patterns, precipitation, and hydrology. The Climate Hypothesis was originally based upon events in the interior riverine Southeast, particularly the demise of Poverty Point, Louisiana. Massive floods impacted this major Late Archaic site during its terminal phase from 3100–2500 B.P. (Kidder 2010: 26). A

large number of studies indicate that climatic events that appear to have accompanied the Archaic-Woodland transition in the Southeast U.S. were the regional manifestations of the major worldwide climatic transition of the mid-to-late Holocene with highly variable expressions globally (e.g., Dark 2006; Drysdale et al. 2006; Maher and Hu 2006).

The ramifications of this global climatic downturn on the coastal regions of the Southeast, however, are still little understood, but a regression in sea level at some time during this period appears to have occurred. Evidence from Murrells Inlet, South Carolina based on foraminifera from within vibracore samples indicates rising sea level until around 4200 B.P., followed by a fall of around two m lasting until 3600 B.P., at which point sea level began to gradually rise again (Gayes et al. 1992). Another study of marsh stratigraphy and archaeological sites suggests a series of sea-level fluctuation during the Late Archaic (Colquhoun and Brooks 1986), though the multiple dramatic fluctuations posited here are not supported by other studies. The most comprehensive recent study from the Southeast, Balsillie and Donoghue (2011), which compiles all available geochronological sea-level datasets from the Gulf Coast, finds evidence for a lowstand occurring at 4800 B.P. as well as a highstand at approximately 4200 B.P. followed by a subsequent, though less pronounced, regression and a relatively slow transgression. Archaeological occupational evidence indicates that by 2400 B.P., the productivity of the marsh system appears to have been restored (DePratter and Thompson 2013: 148), which coincides with a return to modern sea-level as well as much less pronounced fluctuation in the Balsillie and Donoghue (2011) model.

Dendrochronology in the U.S. Southeast

Bald cypresses, some of the longest-lived trees in the Southeast and Mid-Atlantic regions of the U.S. (Edwards et al. 2013: 470), are ideal for creating paleoclimatic reconstructions derived from tree-ring analyses. Although comparatively little tree-ring research has been undertaken in the Southeast U.S., some research does focus on the growth patterns of bald cypress, including the reactions of this species to certain climatic events and ecological conditions as well as differences in growth patterns between individuals growing in different environments. A series of bald cypress chronologies have also been created from some locales in the region. Some studies have investigated the effects of hurricanes on both annual ring patterns and on other tree ring characteristics (Collins 2014; Lewis et al. 2011; Miller et al. 2006; Tucker 2015). Finally, some extant tree-ring chronologies in the Southeast serve as proxies for rainfall in the region, suggesting environmental conditions that may have affected prehistoric and historic cultures and settlements.

Bald Cypress Tree-Ring Chronologies

Bald cypress has been successfully employed in the construction of numerous ringwidth chronologies, extending spatially from the upper limit of the species' modern range (southern Delaware) to the lower limits in southern Florida (Stahle et al. 2012). Stahle and colleagues have created a network of these chronologies, several covering over 1000 years. The longest southeastern chronology is from the Black River of North Carolina, which extends back into the late 300s A.D. As Stahle et al. (2012) note, however, the length of these chronologies is limited by the availability of appropriate samples. Buried bald cypress logs are present throughout the Southeast, but the difficulty

of locating and sampling them constrains many studies. Fortunately, with the help of the Georgia Department of Natural Resources, this project obtained over 100 once-buried samples dating to the period of interest, as is discussed in the “Field Methods” section. This rich data source provides a unique opportunity to construct sequences dating back into the deep past in the Southeast U.S.

The work of Stahle and coauthors shows that ringwidth in bald cypress is closely tied to precipitation, with growing season moisture the major driver of growth, accounting for as much as 50–70 percent of ringwidth variability in some locales. Seasonal climatic anomalies in the region as well as the zonal positions of the Bermuda High, North Atlantic Oscillation, and Southern Oscillation are also contributing factors to the width of bald cypress tree rings in the Southeast (Stahle et al. 1985; Stahle and Cleaveland 1992; Stahle and Cleaveland 1996), indicating the value of bald cypress chronologies to both paleoclimatic reconstructions of the local environment as well as analyses of larger climate patterns and shifts.

Bald Cypress and Salinity Stress

Chemical analyses of wood may record the salinity level of water taken up by bald cypresses, thus serving to elucidate certain types of salinity intrusion events during the lifetimes of ancient trees. These analyses may potentially be used to identify events which would have affected coastal societies, including sea-level rise and storm surge.

Studies of bald cypress sapling responses to freshwater inundation regimes (Young et al. 1993; Young et al. 1995) and a few studies on salinity intrusion (Allen et al 1996; Krauss et al. 2009) provide data on the types of environments that young individuals of this species can and cannot tolerate. Since nearly all of the published

research and chronology construction involving bald cypress growth patterns focuses on freshwater environments, little is known of this species' reactions to ecological conditions and shifts in oligohaline brackish coastal environments where salinity varies from 0.5 to 5.0 parts per thousand (The Venice System 1958). The few published studies on bald cypress growing in non-freshwater environments find that salinity as well as prolonged freshwater flooding decrease adult bald cypress growth and kill saplings (Thomas et al. 2015). Mature trees, however, can persist for decades, particularly on raised hummocks very similar to those along the Georgia Coast, before dying under conditions of persistent salinity stress such as those caused by rising sea levels (Hsueh et al. 2016; Krauss and Duberstein 2010).

Chemical analyses can help identify tree death from increased salinity levels. Yanosky et al. (1995) finds that the detection of chloride, sodium, and bromide in the outer tree rings of bald cypress is positively correlated with saltwater stress. Furthermore, this study finds that the longer the tree is exposed to and survives under salt-stressed conditions, the farther back into the sapwood these three elements penetrate. Bald cypress trees that survive salinity intrusion events, however, expel these salts from their wood. Thus, the presence of elevated levels of sodium, chloride, and bromide in the sapwood of dead bald cypress trees can provide evidence of salinity-induced mortality as well as an indicator of survival time under salt-stressed conditions. Yanosky et al. (1995) by extrapolation thus indicates that relative salinity levels during salinity-induced tree mortality events are possible to ascertain from elemental analyses: the longer a tree survives in increasingly saline conditions, such as those created during sea level transgression events, the farther back into the sapwood chloride, sodium, and bromide

will penetrate. By contrast, an event during which water very high in salt rapidly impacts coastal bald cypress forests, such as hurricane storm surge or a rapid sea-level transgression event, should induce relatively rapid mortality in trees, leaving them with very high levels of salts in the outer rings but little penetration into the sapwood.

Southeastern and Mid-Atlantic Archaeological Enquiries Employing Tree-Ring Data

Extant bald cypress chronologies have been employed in the Southeast and Mid-Atlantic regions to examine archaeological questions. Since, as noted, bald cypress ringwidth, at least in non-brackish environments, is closely tied to growing season precipitation, these enquiries have focused on the possible role played by drought in cases of cultural change. Blanton and Thomas (2008), using Stahle's 1985 chronology from the Altamaha, analyzed precipitation over this sequence, determining that the St. Catherines Period Drought (A.D. 1176–1220) had a major impact on Indigenous Mississippian societies in the area. This study also found another major dry period beginning around A.D. 1360 and a late 1500s megadrought coinciding with the arrival of European missionaries on the East Coast of America. Jesuits arriving on the Georgia Coast recorded the latter event. Stahle (1998) analyzed bald cypress from Jamestown, Virginia, concluding that the settlement of this colony coincided with a major, though apparently localized, dry period that would have affected the settlers' ability to support themselves—a fact reflected in the colonists' epithet for these years, the Starving Time. These archaeological studies show the value of dendrochronological analyses for the investigation of cultural inquiries in the Southeast, a method that this project will also employ.

Paleoenvironmental Field Site: The Altamaha Wildlife Management Area

The paleoenvironmental samples used in this study are cookies (lateral slices of trees) obtained from the Altamaha Wildlife Management Area (WMA) in Darien, Georgia. The Altamaha WMA includes three islands in the Altamaha River Delta (Butler, Champney, and Rhett's) managed by the state of Georgia's Department of Natural Resources for waterfowl hunting (Figure 2.1). We obtained the majority of tree samples from Butler Island while some others were recovered from Champney; no samples were obtained from Rhett's, although it was surveyed in the course of this project.

Plantations were developed beginning in the late 18th century that operationalized the unique tidal aspects of Southeast U.S. delta coastal banks and river islands, including the Altamaha WMA islands, for rice paddy agriculture. Only a small geographic area in the region was suited for such agriculture. Rice fields needed to be located far enough upstream to avoid the detrimental effects of saline water on the rice crop, yet far enough downstream to take advantage of the river's tidal rise and ebb (Wilms 1972: 45). The Altamaha WMA tidewater islands are subject to twice-daily tidal shifts ranging from approximately 1.5 to 2.5 meters (Champney Island Tides 2021), within the necessary tidal range for rice agriculture (Wilms 1972: 53). Rice plantations on these islands used this dramatic tidal flow and complex systems of dykes, canals, and irrigation ditches to manage the input of fresh water and prevent the intrusion of brackish water into the rice fields during high tide (Bell 2010: 167).

Butler and Champney islands, which are easily and quickly reached by vehicle from the WMA office buildings, consist of large, rectangular “ponds”—the former plantation-era rice fields that are usually kept flooded with a few inches of water – each

surrounded by deep canals. The ponds are drained on a DNR-determined schedule. Elevated earthen roadways run between these ponds and are currently used by state employees in vehicles and/or large earthmover equipment for site access and maintenance work as well as by members of the public, generally on foot, in periods of the year during which the WMA is open for hunting. The earthen material used for the construction of these elevated roadways is hauled out of the canals. During maintenance work at the Altamaha WMA, which is largely continuous due to erosion of the earthen roadways as well as the impact of hurricanes and strong storms, numerous stumps and trunks of trees are hauled up from the canals, which staff say can be up to 30 feet in depth. Some newly-piled earthen roadways appear to be composed largely of plant material (mainly roots and knees of baldcypress), attesting to the number of specimens that must be present in the site strata.

While many of the trees recovered from the coastal mud are now simply lying on the surface along the roadways, many others are located within the impoundments themselves, which are periodically and systematically burned to encourage the growth of plants attractive to waterfowl. Although this burning has destroyed the parts of the trees above the waterlogged ground, the portions still in the mud are almost perfectly intact.

Rhetts, which is only accessible via watercraft, is a large swampy island encircled by an elevated earthen road. Rhetts Island was visually surveyed for exposed stumps and logs by the P.I. and Craig Jacobs on a 4-wheeler. The extreme plant overgrowth at the site and presence of numerous nesting alligators along the roadway meant that little on-foot survey could be undertaken. A number of *in situ* stumps which appear to be highly degraded due to exposure and thus likely not appropriate for dendrochronological

analysis were noted to be present in the canal between the road and swamp area, but it was not possible to reach these *in situ* stumps. To obtain samples from these *in situ* stumps, which could probably be radiocarbon dated and elevation-mapped to contribute to studies on sea-level rise in the locale, it would be necessary to bring a small watercraft to Rhetts Island.

Field Site History

Butler Island is the site of the former Butler Plantation, a large agricultural operation largely focused on rice, sugarcane, and cotton where hundreds of enslaved people were forced to labor during the last decades of the 18th and first half of the 19th centuries (Bell 1987). The persons sold at the May 1859 auction in Savannah referred to as “The Weeping Time”—one of the largest sales and family separations of enslaved persons to ever occur in the United States— were men, women, and children from Butler’s Island Plantation as well as from Hampton, the other Butler-owned plantation on nearby St. Simons Island (Bell 1987: 106). Apart from the 1978-’79 dissertation fieldwork by Dr. Theresa Singleton, which surveyed the island to locate historic sites (particularly the four settlements of enslaved individuals) and conducted some testing and artifactual/zooarchaeological analysis at one of the villages (Singleton 1980), no further archaeological work has been undertaken at this part of the site.

The buried trees that were employed for this dissertation project on environmental change in the ancient past are some of the same deposits that first prompted the concept of sea-level fluctuation. Charles Lyell, often called the “father of geology,” visited Butler Island Plantation in the 1840s, remarking upon his approach to the island that:

[S]till we saw the buried stumps and stools of the cypress and pine continuing to show themselves in every section of the bank, maintaining the upright position in which they originally grew. The occurrence of these in the salt marches clearly demonstrates that trees once flourished where they would now be immediately killed by the salt water. There must have been a change in the relative level of land and sea, to account for their growth (Lyell 1850: 249)

These buried trees had long been noted by agriculturalists in the region. Famed naturalist William Bartram, who visited the Altamaha delta and other regions of the Southeast, stated that:

[I]t is plainly to be seen by every planter along the coast of Carolina, Georgia, and Florida, to the Mississippi, when they bank in these grassy tide marshes for cultivation, that they can not sink their drains above three or four feet below the surface, before they come to strata of cypress stumps and other trees... (Bartram 1792, quoted in Lyell 1850: 250)

These buried forests noted in the 18th and 19th centuries, then, may be extremely widespread throughout the tidewater Southeast, but the difficulty of locating and sampling them has prohibited much dendrochronological work on these valuable deposits from being conducted. With the help of the Altamaha WMA staff, we had the opportunity to locate and sample a number of these ancient trees whose analysis has provided the longest tree-ring chronology in the eastern U.S. to date.

Developing the Chronology: Field and Laboratory Methods

Field Methods

Sampling of specimens was largely constrained by the maintenance schedule of the WMA and the availability of WMA staff, who could usually only step away from their typical work for a few hours at a time to cut the samples with a 4.5-ft long chainsaw. Sampling thus took place over two years at short intervals. The PI and colleagues also traveled to the site to locate and mark (with tall pvc pipes and flagging tape) trees and stumps, which were later sampled with the help of WMA staff. Ponds were generally only accessible on foot, since the ground was too wet to support vehicles, so in order to obtain the most trees possible in a short period, our sampling focused on ponds with large numbers of ancient trees. We also obtained a number of specimens from beside elevated roadways.

Bald cypress trees growing in wetland areas swell outwards at the base of the trunk into the buttress (Walsh and Dawson 2014). Combined with an extensive root system, including protruding conical structures called “knees,” bald cypress buttressing ensures that even hurricane-force winds are generally incapable of overturning these trees (Fowells 1965). In this lower buttressed portion of the tree, annual growth rings are less regular than higher up in the trunk. When possible, we obtained samples above the basal swell of logs (typically between 1–2 meters above the swell). Sometimes, however, only the upper basal swell or the lower portion of the trunk was present, so we had no choice except to sample in an area where buttressed growth was present. The first sampling method attempted was coring, which we hypothesized was viable due to the high level of preservation. Although cores emerged with every appearance of freshly-cut wood,

including scent intact, they oxidized, blackened, and turned to dust within one day. This was a similar, though much more rapid, process to what we observed in newly recovered stumps and logs. Although these trees had the appearance and scent of freshly-cut cypress (including, in some cases, the presence of bark and sap), within a few weeks of exposure to air, drying, and temperature fluctuations, the specimens blackened and oxidized. After the coring attempts failed, we turned to whole or partial cookies from tree trunks, which have a better survival record. These appear to maintain their high preservation level due to the structural support of the lateral slice as well as slow drying in the (relatively) temperature-controlled environment of the laboratory and/or the woodshop where most of the samples were sanded.

Since most cookies had to be pried out of the ground with crowbars and carried out of the ponds by hand or in a backpack (sometimes hauling them the distance of 400 meters), the size of samples obtained in these locations was also limited; we usually could only obtain $\frac{1}{4}$ cookies from trees within ponds. Since, as noted, the ponds are periodically burned, destroying the portions of trees not protected by the mud, some of the cookies obtained from the ponds contained only the outer rings (usually ~200-400 annual rings) of what were clearly extremely large and long-lived individuals; based on ring curvature, some of these trees likely lived over 2,000 years. By contrast to the generally smaller samples obtained from within ponds, we were able to recover very large samples (i.e., whole or $\frac{1}{2}$ cookies) from some of the complete or nearly complete trees lying by the elevated roadways, which were accessible by 4-wheeler and/or truck.

We also obtained cookie samples from modern bald cypresses in the Altamaha River mouth area to determine environmental correlates of ringwidth and extend the

chronology to the present. Our sampling strategy avoided felling any living or dead standing trees for this project. Bald cypress trees represent a keystone species for several coastal environments and standing dead trees are important wildlife habitat. We thus sampled only recently dead trees that were fallen or leaning, including one on Butler Island knocked over during maintenance work and twelve on Sapelo Island, the nearest barrier island to the WMA.

Lab Methods

Samples were dried in the climate-controlled laboratory and/or woodshop for six weeks before preparation. They were subsequently sanded to reveal the cellular structure using either a handheld belt sander in conjunction with hand sanding or a large table-mounted belt sander followed by various grits on an orbital sander and finally, a handheld metal scraper, which opened up the cellular structure.

Tree rings were measured at .001 mm precision from high-resolution photographs obtained with a macro lens and camera extension tubes used with a Canon DSLR camera as well as a Velmex stage advancer and microscope in areas where photographs could not provide appropriate detail. Both the Image J extension of Object J, developed for tree ring and otolith measurements and funded by the Institute for Marine Research in Norway (Vischer & Nastase 2020), and Measure J2X were used to obtain measurements. Since photographs can distort images, only the central portions of photographs were used for measurements. Randomized measurements of tree rings from each sample obtained from photos and Image J were checked with the Velmex stage advancer to ensure measurement accuracy.

A combination of hand-drawn skeleton plotting (Speer 2010: 156) and computer programs was employed for crossdating. COFECHA (Grissino-Mayer 2001) was used to support and conduct quality control checks on the crossdating results obtained through skeleton plotting, while ARSTAN (Cook and Holmes 1996) was used to construct the chronology following crossdating. A 5,177-year tree-ring chronology anchored to the present as well as an earlier 529-year floating (i.e., unconnected to the present) chronology were developed from the Altamaha WMA and Sapelo Island samples.

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Figures



Figure 2.1. Location of the Altamaha WMA and the islands of Butler, Champney, and Rhett's in the coastal Southeast U.S. The plantation-era system of ponds and dykes still demarcates both Butler and Champney Islands.

CHAPTER 3

A 5177-YEAR TREE-RING CHRONOLOGY FROM THE SOUTHEAST COAST,
USA¹

¹Napora, K.G., A. Cherkinsky, R. Horan, B. Tyler, C. Jacobs, and V.D. Thompson. To be submitted to *Nature*.

Summary

While tree rings provide one of the best annually resolved climatic proxies available, dendrochronology in eastern North America has long been limited in its temporal scope because of poor tree preservation as well as the difficulty of recovering appropriate specimens for analysis. Here we present a new ringwidth chronology based on specimens recovered from a deposit of buried bald cypress (*Taxodium distichum*) on the Southeast Coast of the U.S. This 5177-year chronology, plus an additional earlier 529-year floating chronology, now represents the longest in eastern North America. Bald cypress chronologies have provided important information about rainfall and droughts and context for archaeology. This chronology pushes the timespan back an additional 2556 years for the Southeast U.S., providing insight into both environmental and cultural shifts in the region.

Abstract

We present a new 5177-year anchored tree-ring chronology and an earlier 529-year floating chronology separated by a gap of less than 80 years from ancient bald cypress (*Taxodium distichum*) trees. We recovered these samples from buried anoxic conditions at the mouth of the Altamaha River on the Georgia Coast. Crossdating of samples was facilitated with 139 radiocarbon dates. The Altamaha coastal chronology presented here is the longest tree-ring chronology developed for eastern North America. This chronology extends existing tree-ring paleoclimate records by 2556 years and provides insights into drought periods and cultural changes on the Southeast Coast of the U.S..

Keywords: dendrochronology, bald cypress, Southeast U.S., paleoclimate

Introduction

Dendrochronology, the scientific study of tree rings, provides high-resolution proxy records of environmental change dateable to calendar years and constitutes an important tool for the development of Holocene climate records. Long tree-ring chronologies created from well-preserved remnant or buried deposits of trees and wood from archaeological sites as well as exceptionally long-lived extant individuals exist for various locales around the world (e.g., Ferguson et al. 1985, Friedrich et al. 2004, Shao et al. 2007). These multimillennial chronologies, often paired with complementary methodologies like chemical and stable isotope analyses of tree rings (Pearl et al. 2020b), facilitate reconstructions of paleoenvironmental conditions across the Holocene. In North America, most of the long chronologies exist for the western portion of the region; however, a few do exist for the eastern United States. For example, a 2427 year-long chronology is available in the Northeast U.S. from subfossil wood (Pearl et al. 2020a). Relatively few, however, exist for the Southeast U.S., with the longest from this region spanning 2624 years (Stahle et al. 2019).

This study provides a temporal expansion of the tree-ring record in the Southeast U.S. and a year-by-year record of ecological conditions in a zone rarely covered by dendrochronological analyses. The development of this chronology has important applications for both paleoclimatology and archaeology in the region. Comparatively few tree-ring chronologies have been developed in the sub-tropics in general (see Pearl et al. 2020), and few tree-ring chronologies derive from locales directly adjacent to coastlines, including only a single study from the Southeast U.S. (Carr 2010) available through the International Tree-Ring Databank at the National Center for Environmental Information

(NOAA) Paleoclimatology website. Long, high-resolution records of Holocene environmental conditions in the Southeast are therefore lacking, largely due to environmental conditions detrimental to the preservation of ancient wood. In arid locales like the U.S. Southwest, millennia-old trees and wooden archaeological materials often remain uncompromised on the surface and are accessible for chronology-building (e.g., Bannister 1965; Ferguson 1969; Salzer et al. 2019). Meanwhile, in places like Western Europe natural and human-altered samples employed in tree-ring chronology construction and dendrochronological dating have been found preserved in acidic bogs, frigid alpine lakes and river gravels that hinder decomposition (e.g., Becker 1993; Leuschner & Delorme 1988; Pilcher et al. 1977, 1984; Winiger 2008). The humid subtropical climate of the Southeast U.S. (Peel et al. 2007), however, requires different conditions for preservation that occur more rarely. The development of multimillennial tree-ring chronologies in the Southeast U.S. similar in scope to those in other world regions is limited by the amount of preserved, recoverable trees available to researchers, with surface wood decaying rapidly in the humid environment and subsurface material usually difficult to locate and sample (Stahle et al. 2012). For this study, we obtained samples from buried ancient trees from a state-managed site, the Altamaha Wildlife Management Area (WMA), at the mouth of the Altamaha River on the Georgia Coast, as well as modern trees from Sapelo Island, the nearest barrier island to the Altamaha WMA (Figure 3.1). Tree samples from the WMA date back almost 7000 years and reach into the modern era. The anoxic waterlogged conditions from which the samples were recovered preserved these trees almost completely intact (Napora et al. 2019). Many of

the trees seeped resin when cut, and bark is still present after millennia underground on a number of trees.

Tree-ring chronologies of the Southeast U.S. provide information about environmental conditions, including droughts, and also contextualize cultural shifts in the archaeological and historical records. Studies have shown that bald cypress growth in the Southeast corresponds strongly to early growing season precipitation (Stahle et al. 1985), making cypress tree-ring chronologies a particularly useful tool for rainfall and drought reconstructions. Existing bald cypress tree-ring chronologies contribute to an understanding of the impact of drought periods on Native American agricultural societies (Anderson et al. 1995) as well as environmentally contextualize the experiences of European-sponsored colonization of eastern North America. Both the disappearance of the Roanoke Colony in North Carolina and the worst hardships of the Jamestown settlement in Virginia were found to have coincided with major droughts in these records (Stahle et al. 1998). More locally, the depopulation of the Savannah River Valley and influx of migrants to the Georgia Coast are also thought to be influenced, in part, by a series of droughts (Anderson et al. 1995). The existing Altamaha River tree-ring chronology, derived from living and remnant trees from approximately 45 km upriver from the Altamaha WMA and covering the time period 929–1985 C.E., has also been used to reconstruct the impact of drought over the last millennia (Stahle & Cleaveland 1992) and its relevance to the archaeology of St. Catherines Island on the Georgia Coast (Blanton & Thomas 2008). This study extends the Altamaha tree-ring chronology by several millennia and provides an important tool for conducting high-resolution paleoclimate analyses.

Results

We developed a 5177-year anchored tree-ring chronology (3161 B.C.E. – 2016 C.E.) from the mouth of the Altamaha River as well as an earlier 529-year floating segment. Radiocarbon dating of every tree sample supports the accuracy of this chronology (Table 3.1). For the anchored chronology, the series intercorrelation (i.e., the average correlation of each series with the master chronology derived from all other series) is 0.248, with a mean sensitivity (i.e., the degree of annual variation) of 0.474. The mean length of a series is 321 rings with a maximum of 1078 rings. The earlier floating segment, which the four floating chronology radiocarbon dates combined with the anchored chronology absolute dates indicate is separated by a very small margin of less than 80 years, consists of only two trees. For this floating segment, the series intercorrelation is 0.321, with an average mean sensitivity of 0.558 and a mean series length of 276.

Number of radii providing replication at any given year peaks at 23 (from 15 trees) in the early 1600s C.E. Replication varies throughout the timespan of the chronology, and there is correlation of increased variability in tree ringwidth with low sample numbers in parts of the chronology, which we account for by detrending with a smoothing spline with a 50% frequency response of 100 years and by detrending the variance of the robust mean chronology with a 100-year spline to minimize the variance changes due to fluctuating sample size (Figures 3.2, 3.3; Tables 3.2, 3.3). These same standardization methods have been used for other bald cypress chronologies on the East Coast (see Stahle et al. 2012). Other methods of variance stabilization (e.g., stabilizing the variance of both the series and chronology; normalizing the detrended series to a

mean and standard deviation of 1 (Figures 3.4, 3.5; Tables 3.4, 3.5) were also tested. The chronology retained most overall characteristics with these other methods but best minimizes differences in periodic variance shifts based on sample size with the normalization of individual series.

Correlation of the Altamaha standard chronology mirroring the detrending of Stahle's 929 – 1985 C.E. chronology (Stahle 1985) is 0.171 and visual comparison strongly supports this crossdating. Differences in the chronologies are likely due to the specific ecological factors operating at the coast compared to those impacting forests farther upriver (e.g., the more-pronounced impact of salinity at the river mouth).

Discussion

This multimillennial tree-ring chronology is currently the longest in eastern North America and provides an important tool for reconstructing the coastal paleoenvironment and contextualizing cultural change over a span of five thousand years. As indicated by the high correlation between bald cypress ringwidth and early growing season precipitation reported by other studies (Stahle et al. 1985), this chronology is particularly useful for reconstructing spring rainfall patterns and droughts on the Georgia Coast. Its deep time span means that societal shifts in the Southeast U.S. beginning as early as 3000 B.C.E. can now be interpreted against annual environmental conditions for a better understanding of coupled human-environment systems, the impact of environmental factors on societies, and the resilience of Native American cultures in the face of large-scale shifts in the coastal ecosystem.

Previous research links cultural shifts on the coastlines of the Southeast U.S., including alterations in landscape and resource usage and differential community

resilience, to environmental changes. For example, Thompson and colleagues find connections between the depopulation of monumental shell ring village sites in the Late Archaic (c. 3000 – 1000 B.C.E.) and possible climate shifts, leading to disruptions of estuary ecosystems and resource bases. At the same time, other, coastal settlements in the region during this period continued to be inhabited, speaking to diverse experiences of climatic change (Ritchison et al. 2020; Thompson & Turck 2009; Turck & Thompson 2016). The new tree-ring chronology suggests periods of enhanced environmental fluctuation during the Late Archaic, which may shed light on the timing of site depopulations. Changes in settlement patterns and estuary resource usage in this and later periods on the Southeast coastline may also be related to shifts in freshwater availability (Russo 2013: 159–161; Sanger 2013: 231–232; Schwadron 2013). The understanding of the temporal correlations between these societal and environmental changes on the shorelines of the Southeast U.S. can be enhanced with the addition of the high-resolution insights into environmental fluctuation, flood conditions, precipitation, and salinity intrusion events provided by this coastal tree-ring chronology (Napora & Jantzi, in prep).

In addition to being a source of information for Holocene coastal paleoenvironment, this tree-ring chronology also provides a tool for dating wooden artifacts or materials excavated on the Georgia Coast and nearby locales. While wood is generally best dated against a tree-ring chronology developed from the same species, dates obtained on pine (*Pinus* sp.) and eastern redcedar (*Juniperus virginiana*) samples recovered from the Altamaha WMA indicate this bald cypress tree-ring chronology's utility to other species endemic to the region. This cross-species correlation of growth patterns is reasonable given that precipitation is the major environmental driver of

ringwidth for trees across the Southeast U.S. This chronology will allow for the annual of dating a wide variety of wooden archaeological materials, including wooden artifacts, posts with preserved tree rings, and carbonized wood from hearths and other features, if sufficient annual rings are present. The presence of a final growth ring could provide an exact cutting date, while the lack of an outer ring would still give a *terminus post quem* for the use-life of a site or the creation of an artifact. Numerous wooden artifacts including canoes (Setzler 1933; Wheeler et al. 2003), wooden posts (Keene & Garrison 2013; Sears et al. 1994), stakes (Clausen et al. 1979; Doran 2002; Duggins et al. 2018), and carvings and masks of animals (Cushing 1896; Spivey-Faulkner 2018), are recovered from waterlogged contexts throughout the region (Wheeler & Ostapkowicz 2019).

Mean sensitivity is high for the Altamaha coastal trees, but the crossdating is strongly supported by the radiocarbon dates from each tree. The relatively low interseries correlation (0.248, compared to a standard of around 0.4 [“International Tree-ring Data Bank” 2020]) is almost certainly due to the combined effects of the tidewater location as well as the characteristics of many of the samples used in this study. The single publicly available coastal bald cypress chronology previously developed in the Southeast U.S. (Carr 2010) has significantly lower correlation than chronologies even slightly further inland (e.g., Carr & Stahle 2010), even with ideal sampling methods. The interseries correlation of Carr’s 2010 chronology is similar to the correlation of the modern Altamaha samples, from which we were able to obtain ideal samples, with the tree-ring chronology. Sampling in living baldcypress is often undertaken approximately three meters up the trunk to obtain core samples above the basal swell, where growth is less regular (Stahle et al. 2019). The samples we recovered were often the lowest section of

the cypress trunks, including the basal swell. These limitations are partially overcome with replication across radii and specimens and could be lessened further with further sampling at the Altamaha WMA. The radiocarbon dates that we ran on every tree provide strong support for the accuracy of the chronology.

The oldest sample recovered from the Altamaha WMA is a large knee (ALT028), an aboveground projection of the bald cypress root system, with an inner raw radiocarbon age of $5940 \text{ BP} \pm 25$, which, when calibrated, yields a two sigma date range of 4900 – 4720 cal. B.C.E. The WMA therefore clearly contains buried trees dating back to at least this time. Thus, there exists the potential to extend the anchored Altamaha coastal chronology by at least another 2,000 years through the recovery of additional connecting samples. Increased replication would also bolster the crossdating and development of this chronology, particularly in less well-replicated periods (see Esper & Gärtner 2001). There are many more trees still buried at the Altamaha WMA that could extend and refine this chronology. These should be collected before this area is lost to sea level rise, as it is one of the many threatened sites on the Georgia Coast (Anderson et al. 2017).

Conclusion

The Altamaha WMA bald cypress samples have allowed for the development of a 5,177-year anchored chronology (3161 B.C.E.– 2016 C.E.) and an earlier 529-year floating chronology separated from the main chronology by less than 80 years. Radiocarbon dating of all samples supports the accuracy of this chronology. The Altamaha coastal chronology greatly extends the time period over which paleoclimatic variables, particularly precipitation and droughts, can be studied and provides annual environmental context for cultural changes and continuity over a much longer timespan

than was previously available to researchers. It also provides an opportunity for annual dating of wooden artifacts and features found in the area. Continued fieldwork at the Altamaha WMA has the potential to extend this chronology even further into the past.

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Figures and Tables



Figure 3.1. Field site map. All bald cypress (*Taxodium distichum*) samples used to construct this tree-ring chronology were obtained from the islands of Butler and Champney in the Altamaha Wildlife Management Area (WMA) and Sapelo Island on the Georgia Coast in the Southeast USA.

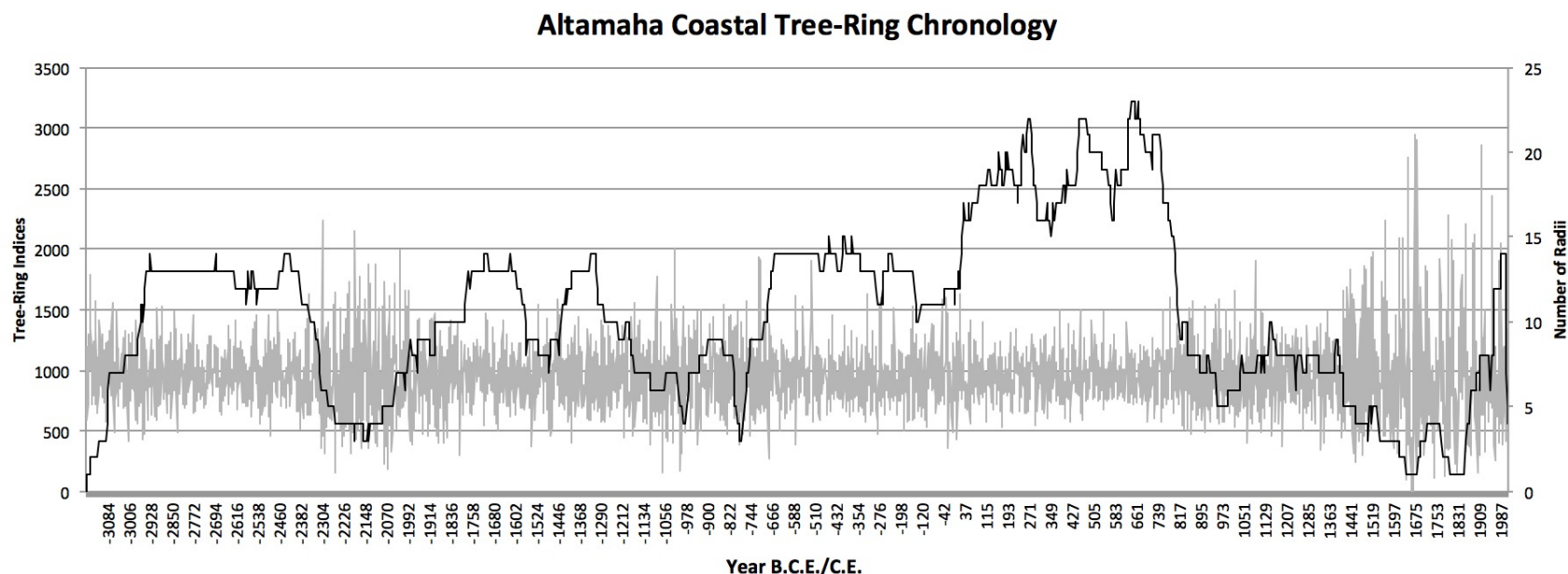


Figure 3.2. 5177 year chronology from the mouth of the Altamaha River, Georgia, USA, also showing the number of radii (1 or 2 per tree) covering each year of the chronology. Here, the chronology has been detrended using a smoothing spline with a 50% frequency response of 100 years and by detrending the variance of the robust mean chronology with a 100-year spline to minimize the variance changes due to fluctuating sample size.

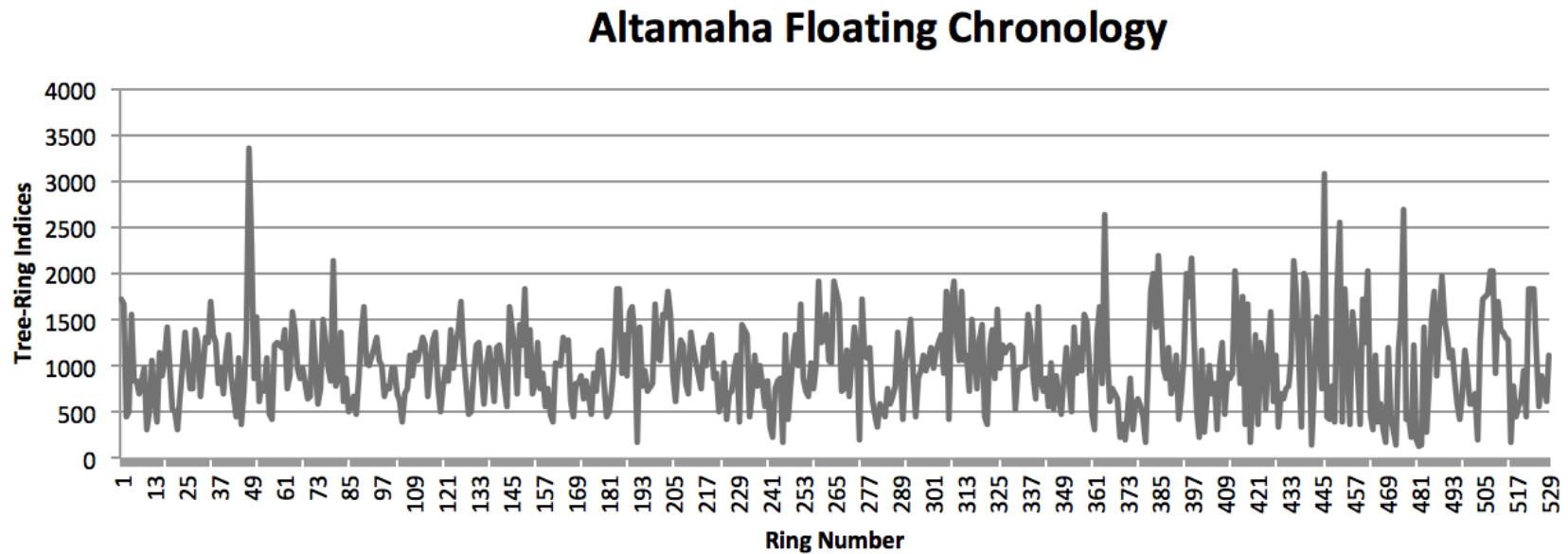


Figure 3.3. 529-year floating chronology based off of the 2 oldest trees from the Altamaha WMA; as in Figure 1, the chronology has been detrended using a smoothing spline with a 50% frequency response of 100 years and by detrending the variance of the robust mean chronology with a 100-year spline.

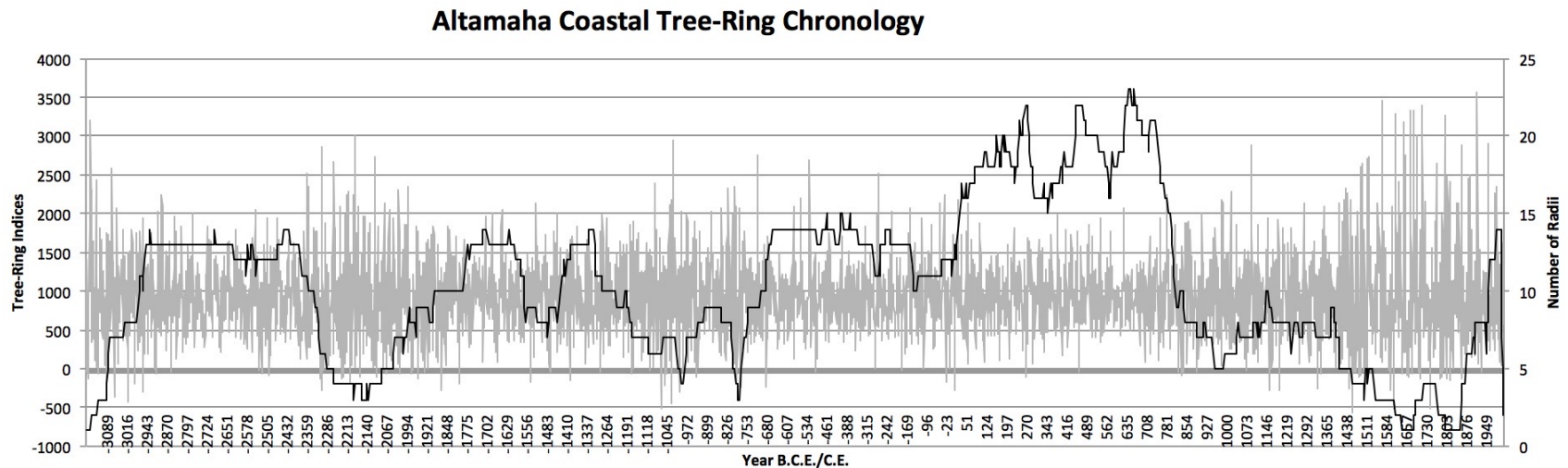


Figure 3.4. The 5177-year tree-ring sequence detrended using a smoothing spline with a 50% frequency response of 100 years and by normalizing the detrended series to a mean and standard deviation of 1 to minimize the variance changes due to fluctuating sample size. Many of the same general characteristics as in Figure 1 are evident, while changes in periodic shifts in variance have been minimized.

Altamaha Floating Chronology

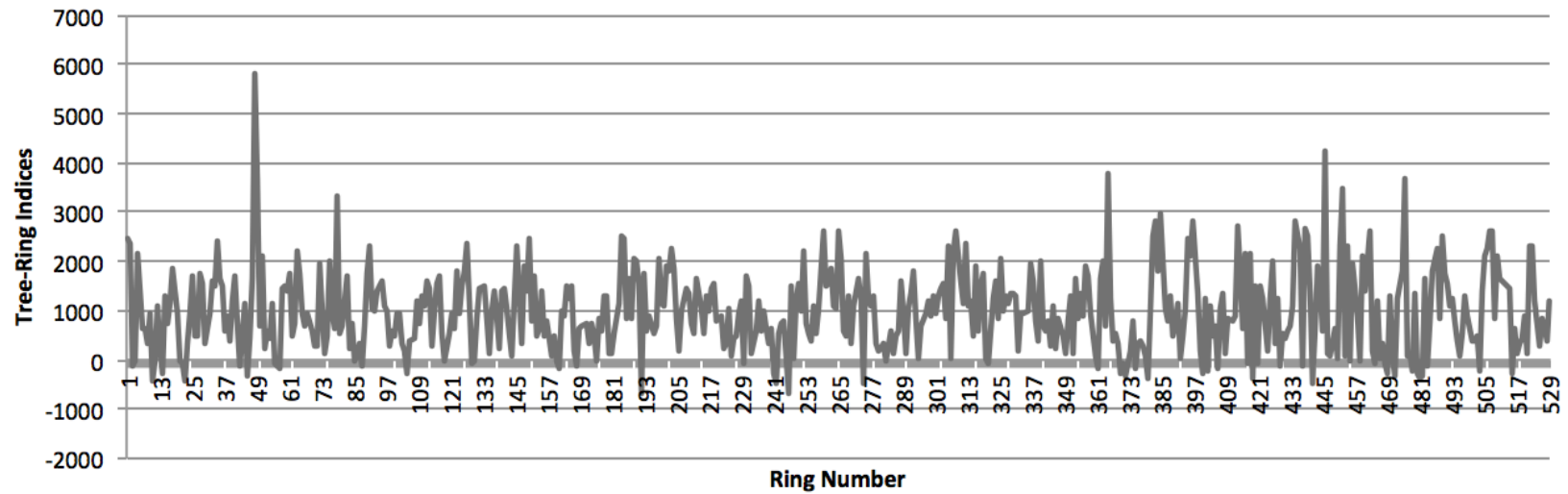


Figure 3.5. 529-year floating chronology; as in Figure 3, the chronology has been detrended using a smoothing spline with a 50% frequency response of 100 years and by normalizing the detrended series to a mean and standard deviation of 1 to minimize the variance changes.

Table 3.1. Information on all tree specimens from the Altamaha WMA as well as modern specimens from northern Sapelo Island (ALT #s 113 – 123). Locational data is not available for several samples because they were obtained either after having been previously cut at unknown locations or dredged from the river; these samples were taken from piles behind the main ranger building at the Altamaha WMA. Radiocarbon dates for each sample were run at the Center for Applied Isotope Studies at UGA, with two sigma date ranges calibrated using OxCal Version 4.4 (Bronk Ramsey 2009) and employing the IntCal20 Northern Hemisphere curve (Reimer et al. 2020) (rounded off by 5 years). For the “In Chron” column, trees that were incorporated into the final chronology are marked with an asterisk (*), and those not incorporated into the final chronology are marked with an “X” as well as a superscript number explaining the reason for the exclusion (1=non-cypress specimen; 2=knee; 3=dateable but poor correlation with chronology due to irregular growth patterns; 4=not dateable due to irregular growth). The dendrochronological dates for all trees correspond with the calibrated radiocarbon date ranges. In the cases of slight discrepancies between the dendrochronological and radiocarbon dates, it was always either the case that 1) the tree either began earlier than the measurements or that there were more rings than the measurements which in either case matched the associated radiocarbon date; in these cases, unmeasured rings were either indiscernible or showed abnormal growth patterns inappropriate for chronology integration; or 2) for a few trees, the date fell within the 99.7% confidence interval but not within the 95.4% interval.

ALT #	Locational Data		Inner Ring Radiocarbon (14C) Dates							Outer Ring Radiocarbon (14C) Dates							Dendrochronological Dates			
	Northing	Westing	Sample	UGAMS#	Raw 14C (BP)	±	δ13C,‰	from	to	Sample	UGAMS#	Raw 14C (BP)	±	δ13C,‰	from	to	Inner Date	Outer Date	# Years	In Chron
1	31.3595	-81.4661	ALT Date 1	26913	3640	25	-27.5	-2135	-1925	ALT Date 4	27104	3140	25	-24.3	-1500	-1305	-1981	-1414	568	*
2	31.3588	-81.4753	ALT Date 3	27103	2430	25	-23.6	-750	-405	ALT Date 2	26725	1700	25	-24.5	255	415	-743	190	933	*
3	31.3599	-81.4762	ALT Date 5	28491	1630	20	-23.52	405	540	ALT Date 6	28492	1220	20	-24.08	705	885	454	686	233	X ³
4	31.3618	-81.4622	ALT Date 7	28493	3930	25	-27.71	-2560	-2300	ALT Date 8	28494	3760	25	-23.9	-2285	-2040	-2560	-2280	281	*
5	31.3588	-81.4757	ALT Date 9	28495	4040	25	-26.49	-2630	-2470	ALT Date 10	28496	3840	25	-25.09	-2455	-2200	-2572	-2153	420	*
6	31.3667	-81.4662	ALT Date 11	28635	3590	25	-27.79	-2030	-1880	ALT Date 12	28636	2950	25	-24.64	-1260	-1050	-1781	-1232	550	*
7	31.3681	-81.5362	ALT Date 13	28637	310	20	-26.47	1500	1650	ALT Date 14	28638	170	20	-25.89	1660	...	1522	1786	265	*
8	31.3614	-81.4595	ALT Date 31	29295	1840	25	-25.07	125	305	ALT Date 32	29296	1520	25	-24.96	435	605	233	570	338	*
9	31.3613	-81.4594	ALT Date 27	29291	2470	25	-25.47	-765	-425	ALT Date 28	29292	2210	25	-24.79	-375	-180	-679	-281	399	*
10	31.3566	-81.4607	ALT Date 33	29297	3630	25	-28.28	-2130	-1900	ALT Date 34	29298	3190	25	-24.44	-1505	-1420	-1954	-1442	513	*
11	31.3558	-81.4609	ALT Date 41	31781	140	20	-25.68	1670	1945	ALT Date 42	31782	10	n/a	-23.98	1895	1905	1787	1909	123	*
12	31.3558	-81.4609	ALT Date 19	29283	4390	25	-26.47	-3095	-2915	ALT Date 20	29284	4030	25	-25.78	-2625	-2470	-2964	-2549	416	*
13	31.3549	-81.4625	ALT Date 43	31783	3850	25	-24.89	-2460	-2200	ALT Date 44	31784	3390	25	-24.37	-1750	-1615	-2181	-1692	490	*
14	31.3617	-81.4624	ALT Date 35	31641	4420	25	-25.16	-3320	-2920	ALT Date 36	31642	3880	25	-23.13	-2465	-2235	-2942	-2313	630	*
15	31.3613	-81.4630	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 57	34994	4060	20	-24	-2835	-2490	-3080	-2577	504	*

16	31.3617	-81.4263	ALT Date 17	29281	4490	25	-26.09	-3350	-3090	ALT Date 18	29282	3670	25	-25.09	-2140	-1960	-3018	-2255	764	*
17	31.3580	-81.4677	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 58	34995	1980	20	-23.48	-40	120	-372	41	414	*
18	31.3552	-81.4655	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 59	34996	2880	20	-24.81	-1190	-980	-1424	-1107	318	*
19	31.3595	-81.4599	ALT Date 29	29293	4980	25	-26.32	-3905	-3650	ALT Date 30	29294	4610	30	-23.41	-3515	-3195	n/a	n/a	293	*
20	31.3550	-81.4654	ALT Date 23	29287	4740	25	-24.59	-3635	-3375	ALT Date 24	29288	4500	25	-24.09	-3350	-3095	n/a	n/a	398	*
21	31.3663	-81.4665	ALT Date 45	31785	1060	20	-24.92	895	1030	ALT Date 46	31786	1100	20	-23.59	890	995	1005	1069	65	X ¹
22	31.3655	-81.4665	ALT Date 21	29285	940	25	-25.92	1030	1165	ALT Date 22	29286	610	25	-24.24	1300	1400	1129	1394	266	*
23	31.3444	-81.4438	ALT Date 47	31787	4350	25	-24.9	-3075	-2900	ALT Date 48	31788	3710	25	-23.88	-2200	-2025	-2951	-2307	645	*
24	31.3442	-81.4582	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 60	34997	3010	20	-23.68	-1380	-1130	-1616	-1222	395	*
25	31.3427	-81.4589	ALT Date 37	31643	3540	25	-27.98	-1955	-1770	ALT Date 38	31644	3040	20	-23.89	-1395	-1220	-1712	-1304	409	*
26	31.3398	-81.4572	ALT Date 39	31645	4120	25	-24.37	-2870	-2575	ALT Date 40	31646	3600	25	-24.74	-2030	-1890	-2534	-1909	626	*
27	31.3383	-81.4560	ALT Date 25	29289	2960	25	-24.98	-1265	-1055	ALT Date 26	29290	2460	25	-24.3	-760	-415	-1231	-781	451	*
28	31.3360	-81.4598	ALT Date 49	31789	5940	25	-25.3	-4900	-4720	ALT Date 50	31790	5520	25	-24.14	-4445	-4330	n/a	n/a	n/a	X ²
29	31.3401	-81.4640	ALT Date 51	31791	310	20	-24.08	1500	1650	ALT Date 52	31792	230	20	-23.46	1640	1800	1509	1642	134	*
30	31.3413	-81.4620	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 61	34998	2070	20	-24.98	-160	5	-668	-138	531	*
31	31.3489	-81.4634	ALT Date 15	29279	2250	25	-26.71	-390	-205	ALT Date 16	29280	2010	25	-23.58	-55	80	-259	53	313	*
32	31.3572	-81.4803	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 62	34999	modern	n/a	-23.77	n/a	n/a	1911	1957	47	*
33	31.3556	-81.4797	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 63	35000	3210	20	-23.54	-1510	-1430	-2038	-1515	524	*
34	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 64	35001	3660	20	-22.95	-2135	-1950	-2460	-2131	330	*
35	31.3624	-81.4635	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 65	35002	3340	20	-23.52	-1690	-1530	-2032	-1591	442	*
36	31.3587	-81.4687	ALT Date 53	31793	2040	20	-24.07	-105	30	ALT Date 54	31794	1690	20	-25.6	260	420	12	349	338	*
37	31.3647	-81.4685	ALT Date 55	31795	3190	25	-22.91	-1505	-1420	ALT Date 56	31796	2950	25	-23.69	-1260	-1050	-1414	-1276	139	*
38	31.3644	-81.4625	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 66	35003	3810	25	-24.52	-2345	-2140	-2927	-2327	601	*
39	31.3623	-81.4717	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 67	35004	2780	20	-24.06	-1005	-840	-1194	-1000	195	*
40	31.3632	-81.4746	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 68	35005	320	20	-24.46	1495	1645	1100	1505	406	*
41	31.3587	-81.4642	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 69	35006	modern	n/a	-24.66	n/a	n/a	1865	1951	87	*
42	31.3601	-81.4639	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 70	35007	1550	20	-24.77	430	580	160	491	332	*
43	31.3593	-81.4648	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 71	35008	920	20	-24.23	1035	1200	n/a	n/a	118	X ⁴
44	31.3576	-81.4701	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 72	37933	2150	20	-23.77	-350	-55	-472	-223	250	*
45	31.3532	-81.4783	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 73	37934	2670	20	-23.59	-900	-795	-967	-791	177	*

46	31.3605	-81.4712	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 74	37935	1100	20	-24.47	890	995	630	930	301	*
47	31.3602	-81.4718	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 75	37936	1590	20	-23.93	425	545	242	534	293	*
48	31.3601	-81.4723	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 76	37937	1710	20	-24.10	255	410	-123	302	426	*
49	31.3603	-81.4724	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 77	37938	1990	20	-22.67	-45	110	-456	17	474	*
50	31.3613	-81.4738	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 78	37939	2100	20	-23.49	-175	-45	-757	-143	615	*
51	31.3616	-81.4733	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 79	37940	2100	20	-22.72	-175	-45	-655	-147	509	*
52	31.3616	-81.4733	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 80	37941	1400	20	-22.74	605	665	64	658	595	*
53	31.3606	-81.4725	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 81	37942	1230	20	-23.58	700	885	608	846	239	*
54	31.3609	-81.4726	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 82	37943	1690	20	-24.34	260	420	24	366	343	*
55	31.3588	-81.4759	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 83	37944	2200	20	-23.71	-365	-175	-766	-261	506	*
56	31.3618	-81.4622	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 84	37945	350	20	-24.56	1470	1635	1121	1545	425	*
57	31.3694	-81.4720	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 85	37946	1240	20	-24.83	680	880	449	850	402	*
58	31.3690	-81.4718	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 86	37947	1220	20	-23.27	705	885	633	760	128	*
59	31.3682	-81.4731	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 87	37948	660	20	-23.61	1280	1390	1146	1266	121	*
60	31.3684	-81.4732	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 88	37949	550	20	-24.49	1320	1425	1282	1413	132	*
61	31.3678	-81.4729	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 89	37950	870	20	-23.75	1050	1225	920	1126	207	*
62	31.3673	-81.4725	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 90	37951	590	20	-24.51	1305	1410	397	1404	1008	*
63	31.3677	-81.4726	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 91	37952	600	20	-23.13	1300	1405	1256	1381	126	X ³
64	31.3666	-81.4725	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 92	37953	1220	20	-25.42	705	885	585	749	165	*
65	31.3677	-81.4720	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 93	37954	1230	20	-24.44	700	885	721	816	96	*
66	31.3667	-81.4725	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 94	37955	880	20	-22.94	1050	1220	832	1159	328	*
67	31.3675	-81.4713	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 95	37956	990	20	-23.77	990	1155	336	1048	713	*
68	31.3676	-81.4716	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 96	37957	1190	20	-23.58	770	890	246	792	547	*
69	31.3675	-81.4717	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 97	37958	1100	20	-24.44	890	995	451	958	508	*
70	31.3686	-81.4716	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 98	37959	1390	20	-23.96	605	670	537	647	111	X ³
71	31.3686	-81.4700	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 99	37960	160	20	-24.79	1665	...	1688	1778	91	*
72	31.3685	-81.4694	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 100	37961	270	20	-23.27	1520	1795	1386	1542	157	*
73	31.3686	-81.4699	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 101	37962	210	20	-23.77	1645	...	1724	1800	77	*
74	31.3636	-81.4720	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 102	37963	3180	20	-24.50	-1500	-1415	-1893	-1479	415	*
75	31.3610	-81.4701	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 103	37964	4390	20	-25.42	-3095	-2915	-3160	-2925	236	*

76	31.3557	-81.4659	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 104	37965	2600	20	-23.82	-810	-775	-926	-804	123	*
77	31.3613	-81.4717	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 105	37966	980	20	-24.19	1020	1155	360	1117	758	*
78	31.3613	-81.4717	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 106	37967	1180	20	-25.37	770	945	355	802	448	*
79	31.3617	-81.4726	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 107	37968	2410	20	-25.43	-720	-400	-760	-493	268	*
80	31.3617	-81.4717	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 108	37969	1100	20	-23.68	890	995	n/a	n/a	271	X ⁴
81	31.3607	-81.4708	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 109	37970	1760	20	-23.08	235	355	4	230	227	*
82	31.3643	-81.4679	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 110	37971	3910	20	-23.35	-2470	-2305	-3116	-2415	702	*
83	31.3644	-81.4680	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 111	37972	2110	20	-24.27	-195	-45	n/a	n/a	308	X ⁴
84	31.3644	-81.4680	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 112	37973	3310	20	-24.49	-1625	-1515	-1774	-1558	217	*
85	31.3644	-81.4680	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 113	37974	2290	20	-23.28	-405	-230	-1472	-395	1078	*
86	31.3615	-81.4804	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 114	37975	3590	20	-22.88	-2025	-1885	-2134	-1957	178	*
87	31.3507	-81.4705	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 115	37976	2960	20	-23.00	-1265	-1055	-1469	-1167	303	*
88	31.3603	-81.4730	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 116	37977	760	20	-23.79	1225	1285	1042	1241	200	*
89	31.3600	-81.4736	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 117	37978	570	20	-23.54	1315	1420	1230	1303	74	X ³
90	31.3592	-81.4743	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 118	37979	930	20	-23.81	1035	1165	995	1131	137	*
91	31.3589	-81.4746	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 119	37980	1140	20	-22.65	770	990	583	958	376	*
92	31.3596	-81.4744	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 120	37981	1170	20	-24.78	770	960	370	810	441	*
93	31.3599	-81.4751	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 121	37982	1690	20	-24.68	260	420	28	340	313	*
94	31.3600	-81.4751	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 122	37983	1190	20	-23.51	770	890	647	804	158	*
95	31.3601	-81.4745	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 123	37984	450	20	-24.15	1420	1460	1245	1460	216	*
96	31.3604	-81.4749	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 124	37985	850	20	-22.93	1160	1260	1164	1234	71	X ³
97	31.3607	-81.4755	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 125	37986	1540	20	-24.31	435	595	185	591	407	*
98	31.3611	-81.4751	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 126	37987	1700	20	-23.41	255	415	122	297	176	*
99	31.3610	-81.4750	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 127	37988	1840	20	-23.33	125	245	-406	198	605	*
100	31.3611	-81.4750	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 128	37989	1610	20	-24.40	415	540	21	482	462	*
101	31.3617	-81.4736	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 129	37990	2200	20	-24.63	-365	-175	-975	-342	633	*
102	31.3615	-81.4735	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 130	37991	1180	20	-23.57	770	945	162	758	597	*
103	31.3612	-81.4734	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 131	37992	1300	20	-23.76	660	775	406	717	312	*
104	31.3577	-81.4760	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 132	37993	2020	20	-23.83	-55	65	n/a	n/a	286	X ⁴
105	31.3580	-81.4756	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 133	37994	1910	20	-23.63	65	210	-241	134	376	*

106	31.3579	-81.4755	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 134	37995	1750	20	-23.49	240	375	-37	252	290	*
107	31.3549	-81.4755	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 135	37996	4050	20	-24.00	-2630	-2475	-3083	-2541	543	*
108	31.3557	-81.4737	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 136	37997	2810	20	-25.10	-1015	-905	-1406	-990	417	*
109	31.3558	-81.4736	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 137	37998	1280	20	-22.98	670	775	629	758	130	X
110	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 138	37999	130	20	-25.63	1680	1940	n/a	n/a	167	X ¹
111	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	ALT Date 139	38000	240	20	-24.51	1635	1800	1691	1737	47	X ¹
112	31.3523	-81.4785	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	1991	2016	26	X ³
113	31.4566	-81.2443	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	1962	2010	49	*
114	31.4563	-81.2438	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	74	*
115	31.4561	-81.2441	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	98	X ⁴
116	31.4559	-81.2440	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	1968	2016	49	X ³
117	31.4724	-81.2491	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	1949	2008	60	X ³
118	31.4720	-81.2494	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	67	X ⁴
119	31.4712	-81.2490	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	1965	2006	42	*
120	31.4712	-81.2490	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	1958	2013	56	*
121	31.4712	-81.2490	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	1878	2008	131	*
122	31.4712	-81.2490	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	1954	2015	62	*
123	31.4717	-81.2496	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	n/a	1862	2009	148	*

Table 3.2. Version 1 of the 5177-year tree-ring chronology indices in Tucson compact format; the chronology has been detrended using a smoothing spline with a 50% frequency response of 100 years and by detrending the variance of the robust mean chronology with a 100-year spline to minimize the variance changes due to fluctuating sample size. The 3 versions of this chronology shown respectively in the table are 1) the standard chronology (ALT1Std), with no autoregressive modeling, 2) the residual chronology (ALT1Res) (which removes autocorrelation with autoregressive modeling), and 3) the ARSTAN chronology (ALT1Ars) (which reincorporates pooled autoregression into the residual chronology). Note that in decadal format a “Year 0” is automatically inserted between -1 (1 B.C.E.) and 1 (1 C.E.), throwing off labeled dates by 1 year.

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5177=N      -3160=I ALT1Std      -5(13F6.0)~
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77016128786 78431111082 80232148661 87017127951120535 84591100521 88694123080
111482 76462126301 85040113356110155 82520 69815 98068 86744124528106100118108
98171 65393113144 91179 96316123126 94639 97472125579113168 80729 98301 82499
97493105190 69492 81180 85140 83120 59324104031 69374 82480 94374 64445101902
75335119505 91466 74340103317 75647 84203 44405 96502 58238127401165371100996
57597103665 68523 88296 49666117297 76242143437 74072 95685102666 88357 79092
114457168141143123 74636 90218 84028121691 96703154808182563167208 63837 95686
69522107672160416118482162972105801 67866 50079 31183 63052 93365 58133158347
159431123002120155116597 38832138392 25049109725 54707105059 67651 80463 73321
53925 55654 41685 60528 44155 78791114317142652 51701 48451 48018 50846 94601
90507107120 95065 29915 47780 46860 44153 59542 45270 86823185921 91904114818
107331 81767183792133058 90785 98806 66705 99596107747110689 58721151124149098
151131 77801 79175 57862100730 62701 83319 80520 78350 50707143706192761 99642
182787180519197860165697104530 81038125946108396 71511 88043 78343 87555 87288
75348 93918115253 74717 70581 81443 89285 69795 73490 85578110583108319 48191
49132 87083 80301 74966 89403 89343 84273128647173150168790135815 98718160676
135481136488138862 46077117591 68142 45726 48822 48240224153163252131645 61714
87007138944156888 89081103531 64447102111127623130435 71432 94562 41205 93545
85319 76841 99129 76237 83820108724 95079100817 75322101790 50537 35894 80564
88566 91261114017 82741 83663 53527119258139488145119 70689 90629 82960 87349
91192 91432110457 31706 91975103929209774143608148566113453145265126326131209
119646 60539115177134341 90563 77648123004209328125705123314138434157097 73330
158975 73044115613 70418 68847 13580 10130 48644 10190100254119874276123 75573
166845 38570249995185544 50514110083 86647 35750 68008 35046 12884 41474 418
44795 12732 3879 24041 0 42797 78194 31364130901112904259914294454177124
155889141838 52060 18305 86660 37669108821132221290870122608136432 89646154120
111142101325 37229 73377 61689 76789 84767113466168607117099106397 76062 96779
65052 55963 81563 49053 69202114794 75778 86129 29602 54017141081136851185978
121438108328146407182132110388 61751 87892 55991 65438 73840 75382 92423 95841
63125142054 55281 64373 76329 91605 98267 87115103816 67671 83916 39769 48474
70303 99244 53532 11335 68429 46430 90627 44183 90046 83263 68426 95784126688
56318 82927 49789 94396108260 71045 75262 52601 97701143390165212143931 97074
191990157421162523 74906 86091 77305 79639112785108586 87013 55524 65585 97555
124616 51474 22658 55993 26422 12654 56116 73786106186 35285149535147505106368
115663128453 49416 34149207387227587138618 87900 94366103022 35326126550 76483
120384166390207565194458173508126572 79273 93009129152190754116837125182 63971
23590 35988 52588 35727 44919 85320 85936 73161 54266 25083 15104 34207 19299
69156 91720103688135221103617 86647 82504 84109136578142341118229 89082164518
178279 98087171113128312 97360 76762152050143709 94855118054135010108762 90699
32442 47720 43976220256 50060 85475 38095 57321114736 41105 54769 38213 37777
66646 77941 95620 50065 37408 55328 29693100932102359 35659 52314 71005 92587
204699 83233153128 95234116287203847130500163504212382105034106847 63490 50732
42487124112 50477 60516 48056 15931 83878 68488 48826 55073 70938 94757 30838
69695 48947 34309 95561 83241 80080 63982286407185238 74085137428 90091127638
80161 83001 64007102153149670 75055109341 33321 90720 83765104353 75105 84531
111267 97776 84674 81604 87293102249 78155 57845 71478109588 90322 96436 89817
107744107712112810 73387 88486110857 69393 78447244467 70927121153 89533 79260

83900 68256 82769 39119 26178 97925119426 85662 68989 79489 61330 77204 82722
 95850 91351 65225104308 39498166997106273190453111168 69814 91311 90071 55953
 74715204994163600101948117849 92273 54450 60589 84372 38283 92136 93756 93234
 120020114713 79920 53653 92601 83598 41621 52346 92754110478111439133991 75433
 74043

Table 3.3. Version 1 of the floating 529-year chronology indices in Tucson compact format. The same information as in Table 2 applies.

529=N	1=I ALTF1Std	-5(13F6.0)~
173402168117	45724 49514156850 82721 82559 68872 97396 30601 50983106362 63499	
38784114643	89185108196143089 97706 52340 49625 31094 68378 99203135082 76505	
76253137909126489	68513 98968131192126253169107132694126173 80827 96505 71497	
110697134560	88468 45175107606 36954 73375132188335717231711 86014154249 62702	
80089 74162108425	47848 43139123101126048119834137869 76318 88745159830138183	
97032 85771 98254 79962	65065 67200147797100765 59210 75424150751 97596 83601	
213135 77762 86447135244	62255 87933 51505 56640 67634 46715 86417137633164504	
102307101905120987130599105084100693	66290 79823 76501 97132 97999 69551 61229	
39884 70672 74648111023	89129114355105540131187123865 66799 99664128813135457	
78240 49919 77935 98307	84162140365 98294124980140371168538110315 47681 52211	
92924122451125121	92264 59149 99285120338 94416 62299119620122599103623 78625	
57122164171118549	69050145676123655183199 89130140626 70935 80632124091 74844	
91236 54999 75482 48128	39762102856 99183129615117606128216 62132 44520 82513	
82035 88841 66036 83036 68744	47309 91698 74019115396117535 46308 50968 71582	
109966184721183347	91549132419 88959157998164697133552 18951143155 79275 94032	
72089 79955166091109127106247156286152604179774155817	90032 60822 99311128311	
122116 82214 71293136744116873103461	74774119665 99489125923134561 85926 93520	
50838 62536103234 43440 70723 71821	98237110866 40057144979132445 45378 72875	
112200 79004100513 81311 56964 82648	34686 22599 75947 83834 86868 16485132448	
43801111391134099100754168014	86387 72938 66206104338 75992101548191114124270	
127429155343105688103467192054167803	73388 74805116701 67536106427142818115760	
19569171070112900107759120166 64647 46990	33422 58335 45063 75312 58220 68430	
78257135669100868 43914103005123473149971	92136 44219 85842 95291111625 93770	
120899 98248118488126349134711 93496180104	42898168922192303137193107226179733	
104895110607 72974150241 75752128862145191	46014 36544120721140450 8702562251	
96774122454113362120881122287118961	53507 95276 96878 99595156673136693 88604	
65308163487 84841 73919 86819	57166104213 54186 90521 67157 48021 86453119033	
49806141258 91082121112 94284156963147740	99175 46851 30333134697163170 80807	
264461112525 61694 74988 63789 24338 35696	20477 50878 87447 31466 59708 63375	
57693 44211 19015 97260180590198693143198219933	99022 86846120557 68759 83879	
111383 43764 70014107904199383174837217607127714	53077 24532117815 29194100898	
70301 81963 31316100309124677 47431 93447 87933	91788202675149373 80352173777	
35858168072 18219134057 37137126400109545	53007110925158970 60665110334 32916	
71025 64672 77030 79529109792214771177296	34710200790191010126626 15135 79055	
153554134269 75076306865 46514 41504 78636	38718182769256509 41106182050 36222	
159362128525 92916 36177171724124927201802	48330 31663112287 39294 58095 31223	
18754119085 45282 16094116142151772270134	42291 44258 22799121523 20120 13352	
15375142872 28793 84695151690180611 89959197470149319134987107802116631	88083	
59974 41461 71113118360 92711 59868 60214 69061	21516127043171110179041202614	
202432 90976168917140053136737132031127150	18577 77700 46501 56992 64340 94213	
45616182991182371113580 55885 90497 77577	61079112452	
528=N	2=I ALTF1Res	-5(13F6.0)~
146840 25936 64202170470 66106 86826 73184105563 30734 69931119569	61106 48465	
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143602115533 61085106879130673116964161188112884116169	72644101004 71985117839	
130836 78398 47535121587 34286 89456138521326429167471	49971157397 47250 89627	
78437114725 44492 56713137422119086111593131608 65321	94406162195121311 86622	
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47346 91855138294 51877 97752 54315 69170 78694 54678100248140191153038	83922	
100678119536124663 96190 98495 65635 88060 81253102577 98197 69368	68795 49530	
86199 81713117069 85808116396101153128716114919 59985108184128609127007	68220	
55376 90723103216 83498143604 86029124722133207156805 90612 43708	65661105091	
100826111973 84587 64515111134119373 60911 63271129638125564 96108	79318 62791	
167240114929 57310152785105502184248 67829143574 66345 89654129876	69231 98500	
56416 86768 54866 55230120375 96910128011106611120212 51898	54222 98188 86551	
94250 68210107269 92672 57164101414 79302137358101286 60645	58159 95779127103	
199643154384 77806127788 91940161132145517109301 2318175336	67803 80399 71897	
86438166757 89619 88566149234139934160840131786 88374 64576	96762125720116200	
73507 71702140018 96570108735 73609123570 94382128639115677	65857 86077 53959	
73240108728 50101 85052 80914107356127864 35848156706136805	32197 79329120575	
74250 88358 78731 61784 80211 54199 44971101106 97687 96831	6952152769 42108	
129472138191 90453175201 68043 76802 74560 89774 65548118461175706124782119342		
152698 97363104002190216141776 43666 98335124929 69019113085142052104450	10140	
188053108776 94571116590 53879 60550 23601 85955 50510 95826 57870	79356 95182	
127994103539 38114116941131868145423 97436 44632 99846104414109270	90541121188	
94783116693119497127048 81401181103 20709177616172809105335	97856170627 83502	
103449 70068158165 57957135823135823 34448 49631138980144345	71676156078 83602	
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70412 45815 83548 77669 34082 56975 42054 77127103979 40159 78322 77356 69509
58410 36989123942182900171349110841206057 71471 72850126765 66403 87447118351
44423 82422118888203614152883187050 89911 32618 29466139693 33515113663 70063
94423 37628120024127567 47229 99408 97314 91677201431126049 58290175589 22126
174596 1890151620 37610137523102536 47030128656154271 44827117575 30040 91083
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125048 78527 33475185746118902186421 23631 27888133499 46185 68753 48242 38687
145366 52581 24560141000160052257007 -297 32140 42257145138 26346 26676 44331
172360 31650 93230164686171915 65883188067130424110951 93487109946 83389 59940
51243 88666132384 92809 58635 69369 82754 33332147443176376160388176136169969
55242156242126915118624118655115211 8308 90553 62376 71016 80575107457 51655
194880172692 84747 41354 97544 85571 66885123480
529=N 1=I ALTF1Ars -5(13F6.0)~
173402162300 39058 51366160227 78791 82359 69469 99132 30551 55303110154 63245
40724118801112184112660142710 94642 50968 51205 33533 72286101158134839 73731
76223138594123662 66069 99732130616123413166120126810121816 77239 96210 71187
111771133315 85415 44463109890 36369 76054133478333480216647 74540152034 58210
80825 74399109333 46458 45436125930124547116763135139 72722 88660159807133907
93764 84435 98077 79139 65267 68516149101 97092 57351 76624151807 94675 82552
213746 71304 85811135305 59313 89182 52037 59068 70073 48375 89375137953161032
96777 99999119536128777102251 98969 65418 80776 77204 97775 97728 68890 62243
41578 73894 76215112059 88348113942104089129578121149 64439100694128755133064
75184 50149 80223 99050 83298140086 94472123558138169164844104270 44607 53994
95401 99857111943 87102 61799103088120023 65128 55927120355129852102396 79823
58542158598127252 63049145003114981187404 86239140676 74913 84370126584 74830
93199 54983 78196 50274 44756108739 98751127748112455122835 56708 45103 86626
83734 90824 66277100166 92707 55628 92068 77632132646108162 62364 50232 85297
124007204699176436 93905126504 97522160610158283121577 6863155719 79539 76090
66861 79458162430102768 89149146949149822171334146810 98234 64204 89223123450
121139 77960 67060133080103538109481 75606118432 98264128273121632 70413 79845
49714 62649100861 50282 74580 75560102208128329 41815144451146167 41921 67097
113645 77124 83540 75264 56574 71064 48105 34040 87213 94994 95776 6062132983
49055118742142138 99328175060 83852 73401 68958 83236 62017110461177910141191
128018158599109705106046191489161046 56523 89178122650 73789107564143645113643
13013169731123463 99512116487 57352 51568 13400 67715 43710 83970 54494 69771
88815125638108940 39997104303132775152326108457 46413 88559102005109692 92582
119626 98917116465122965131885 88117178600 37264164402186374123527102811171219
98502103134 70728152000 68909129275141989 43292 37687125856149791 82163152321
94622114493116243118205115649117095 46967 97559 95984101251156844134514 81339
61739165871 84607 70727 83143 57676107906 57113 93921 72626 47640 93694126141
51476144983 93323115752 91180157115144817 93523 40281 29243140439162189 80767
256969103473 46547 72290 71833 28150 41842 29805 62342 96047 39326 65542 70098
63211 50661 26597108482184686189186129626212297 95123 71823120830 70791 81295
114411 47459 71355112855206322175277202905111585 35058 15788121956 38139100634
70196 88146 35131106361128907 53317 89576 95118 90649199461146998 68189168888
36636161250 14790133673 44702125876107986 48712117853158031 57050108529 31836
76726 73094 81114 86284112703222206171236 15120193568188569110911 6842 78380
159653131844 71863295849 48214 12450 79629 46891187084266186 31986157764 39538
148127135185 85937 30513171110133879193556 43337 15954115797 49512 58119 39421
25928129764 58850 15893123285164957270689 35654 18588 25109129365 32531 12466
25894156752 43603 81351160758184713 83725184639148251121114 97934109511 85392
56863 42157 76483127431 98587 58337 60593 74454 27951132267183172177906192545
189462 74085150783137612126546124246120318 12588 72142 56509 61855 72541101673
52007184771190547103818 42158 85362 82488 63196115728

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Table 3.4. Version 2 of the 5177-year tree-ring chronology indices in Tucson compact format; the chronology has been detrended using a smoothing spline with a 50% frequency response of 100 years and by normalizing the detrended series to a mean and standard deviation of 1 to minimize the variance changes due to fluctuating sample size. The 3 versions of this chronology are 1) the standard chronology, with no autoregressive modeling, 2) the residual chronology (which removes autocorrelation with autoregressive modeling), and 3) the ARSTAN chronology (which reincorporates pooled autoregression into the residual chronology).

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5177=N -3160=I ALT2Std -5(13F6.0)~
104934 -3488140779110676 49238-13615 20140 70301 39030 90834 92991179966114571
169426102899193265320612204741229615184963194718125859153551 33831 57219 95007
90733163707 39221 55616 78308 79041104965 84451 12833 10151243776167302 92313
24710 1849119468 -7473 -92 38500101676 63267 79636 96135 95188182446167634
72595 88568 96703 76902118406102366166292 49667 37749126087119936 81943 62792

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114767 33110 52953 5091 7919 60283 87692 96072175109167817-19054104875 75211
39267172409 20244 76589 98287153071 89475163453120361 95278107222 51025 78624
120708140236165521200751257991106344 46789108124128481139211 80915-36117116675
58534 60362 78698 39311 42026 87334206254 93162 82566207175 50158 36742 6241
52788 82565131594101800 59834 63783107292106131 95669 96079135157140079113711
95034103732 74979179164108119 50402 36434 90680103172 76718120771104884145902
42009114923112528 74295 57411 83965 66288 43176 42579 25230-41906 17645 84434
148671148657 88624 92020151398 42939130602144694 97001 71610148778 75597178996
161897 85407122773 92218 65686116359 88875 99045112540-18670 93661 91505108018
152693175930116157160687132692 23183100796 76873149073 91223 76794 41292 93876
33310 81789 86427176302137625 85980 83760 43220 82899 94827145235-29297194580
162357 76739 85814 19579 63201 28683 87099150141 84877 82566130437 95448 90686
69990124280134162 18088 52379 28755 97799 81341113312107944108481 76122 80871
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203202 96431159302 86870 54230 97597 47197 99714 60262109445148447157164 66221
42179 73959 9450517144822339220886315481115929 61412 74133 86733 83431 90622
68473 62254 98115 58181151823142947104508 24478 65087 93637 47483 77992135755
106901 69897145279104210 67403142746 90269163687109210 86094 86420 36658 -4495
58275 83593108608113095 60147 98208 95640 98032129478 72020 86161 68563196377
107663135344176464132225151282 97176 12361 63208107985 77317 43391 94978 98413
116400105225117859122470 86849 66339125349130076128823129067 90745183310 81909
52896 77026 20623 77795 87473 94916 39491 98954 66403 47073 71273 81498 32100
124885 36613 79817 80986 41043110764 71725144421 62184158558 60684 64792 73621
112948141341 85117101129 90461 86415109900 89987144444116649139442200686133156
50411113066137765138014 70622 91120 89017 82509 26999100279100293120348105053
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Table 3.5. Version 2 of the floating 529-year chronology indices in Tucson compact format. The same information as in Table 4 applies.

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156944105586 2337130276150292171683 94039 6168 95615109779116809 85168133470
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220914 45255 60478 85453 36519130980 26163104959 62285 23546108099145664 5586
189444 73478130919 76585196536148382 70064 2225 3664201118184953 49031372292
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127688 42297418425-46823-20951 96953 22959255081333316-62572213766-14617196048
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171593 25167-19053164702194770347776-58280 -7091 8874171234-16234-15713 12148
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23056 82114151106 88652 34721 51660 72784 -5210174870220530195300220152210420
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59835 47249118207 0 0152028149907134587172271 44808 76549222194171030
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0 45309 50568124588 76019128617108275160897144190 27950101407158809168440
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89687 99118121204 73280 21981105602140631 28311 9626140544161549105530 59138
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18955 34472101281 20564 55932 56264103590157464 0167185172761 8896 47475
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39880 18407 0119733257585266340159328285834 98671 55763132667 53796 70517
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Methods

Fieldwork

Very large bald cypress trunks, roots, and knees are often recovered during ongoing Georgia Department of Natural Resources (DNR) maintenance work at the Altamaha Wildlife Management Area (WMA) on the Georgia Coast. We obtained cookies (lateral slices) of buried bald cypress trunk specimens via chainsawing. The vast majority of these specimens came from Butler Island, a large river island and former rice plantation in the Altamaha lying south of the town of Darien.

To extend the chronology to the present as well as to determine environmental conditions driving ringwidth in this specific locale, we also obtained samples from modern (known fell dates and standing dead) cypress trees on Butler Island and in a stand on northern Sapelo Island, the nearest barrier island to the Altamaha WMA and closest sampling-accessible modern trees.

Radiocarbon Anchoring

To facilitate crossdating of the recovered specimens, whose age was unknown, we radiocarbon dated trees at the Center for Applied Isotope Studies (CAIS) at the University of Georgia. For the first 26 trees, we dated both an inner and an outer ring, and once we determined that ring counts matched the timespan indicated by the radiocarbon dates, we continued dating trees at only an outer ring. A plot of the radiocarbon dates, calibrated via IntCal13 and rounded off by 5 years, were presented in Napora et al. 2019. Table 1 provides a range of information on each tree from the Altamaha WMA, including raw ^{14}C BP dates from CAIS and calibrated date ranges using OxCal Version 4.4 (Bronk Ramsey 2009), employing the IntCal20 Northern Hemisphere curve (Reimer et al. 2020).

Measurement, Crossdating, and Chronology Development

Hand-drawn skeleton plots were used to crossdate specimens, with COFECHA (Grissino-Mayer 2001) used to verify the results. Tree rings were measured at an accuracy of 0.001 mm using a combination of a Velmex stage advancer and Measure J2X software with the Object J extension of Image J (Vischer & Nastase 2020); randomized digitally measured rings were verified for accuracy. Either two or one radii (the latter in the case where only a single line represented regular growth) was measured per tree. For the more recent samples, we used both inter-sample dating as well as Stahle's chronology from ~45 km upriver (AD 929—1985) at the Altamaha (Stahle et al. 1985) for crossdating.

ARSTAN (Cook & Holmes 1996) was employed for chronology creation. For Version 1 of the chronology, we used a 100-year smoothing spline with a 50% frequency response, which has been shown to remove the majority of age-related trends as well as

retain almost all of the annual and decadal variation (Drake & Naiman 2007; see Cook 1985 and Cook & Briffa 1990) as well as a stabilization of the chronology variance with a 100-year spline to minimize changes due to fluctuating sample size (as used in Stahle et al. 2012). For Version 2 of the chronology, we used a 100-year smoothing spline with a 50% frequency response and the detrended series were normalized to a mean and standard deviation of one to better account for fluctuating sample sizes throughout the chronology.

CHAPTER 4

FIVE MILLENNIA OF ENVIRONMENTAL CHANGE

ON THE SOUTHEAST COAST, USA: A TREE-RING BASED ANALYSIS²

²Napora, K.G. and S. Jantzi. To be submitted to *Palaeogeography, Palaeoclimatology, Palaeoecology*

Abstract

This study employs a multimillennial tree-ring chronology developed from ancient buried bald cypress (*Taxodium distichum*) trees on the Georgia Coast to reconstruct elements of the coastal environment over the past five millennia. This analysis provides a complementary set of high-resolution environmental proxies previously lacking for the coastal Southeast United States. We examine ties to precipitation, salinity level, and hurricane activity as well as tree life histories and ringwidth extremities and fluctuation, situating environmental changes within a press-pulse dynamics framework. Elemental analysis of sodium in the outer rings of ancient trees is used to reconstruct salinity intrusion events in a novel application of this methodology. These integrated analyses provide insights into regional ecological conditions in antiquity.

Keywords: Climate change, coastal environment, dendrochronology, Southeast U.S.

Introduction

High-resolution paleoenvironmental proxies bolster the understanding of the ecological events in antiquity. In this article, we use multiple lines of high-resolution proxy evidence from tree rings, including novel elemental analyses, to identify and analyze environmental changes on the Georgia Coast over the last five millennia. Many of the environmental events which occurred on the southeastern coastlines can be addressed on a high-resolution scale by tree-ring analyses. These events include the timing and intensity of droughts in the past and the impact of hurricane and other storm events (Liu and Fearn 2000; Saunders 2010), and shifts in sea-level, the timing and number of which have been variously stated from different proxies (Balsillie and Donoghue 2004; Colquhoun and Brooks 1986; DePratter and Howard 1981; DePratter and Thompson 2013; Gayes et al. 1992; Thompson and Worth 2010).

Here, we examine a 5,177-year tree-ring chronology (i.e., a site-level representation of yearly tree growth (Speer 2010: 4)) that covers the years 3161 BCE to 2016 CE) from the north Georgia Coast (Napora et al. in prep) and discuss the environmental information gleaned from this chronology. This information includes characteristics of the chronology itself as well as specialized tests which provide insight into sea-level changes, droughts, and salinity intrusion events (e.g., storm surge and/or hurricane events) in the region. We use modern environmental datasets to determine the factors influencing ringwidth at this coastal location, analyze concentrations in salts in the outer rings of the earliest trees to investigate salinity intrusion events, and interpret ecological changes visible in cypress life histories.

We conceptualize environmental changes on the Georgia Coast using a press-pulse dynamics framework that distinguishes between more enduring state shifts and relatively short-term anomalies (Bender et al. 1984; Thompson et al. 2016; Yang et al. 2008). Coastal dynamics are viewed through Simiand's (1960) *histoire événementielle*, which visualizes time as punctuated by episodes of change. Tree-ring data identify ecological *événements* at timescales often invisible using other proxy data, including annual and sub-annual events.

Our analyses point to several major environmental occurrences and state shifts on the Georgia Coast over the past five millennia. Tree-ring data may indicate a multicentury period in the early 3rd millennium BCE characterized by continuously flooded conditions. Major environmental upheavals on the Georgia Coast include at least two periods of unstable environmental conditions indicated by enhanced ringwidth fluctuation. Elevated sodium levels in tree rings suggest four periods of enhanced salinity conditions leading to mass tree mortality events prior to 700 BCE, at least one of which likely indicates a period of rapid sea-level transgression. Finally, our data indicate that a shift to less stable or more frequently punctuated environment in which the lives of mature bald cypress tree were regularly cut short commences with the onset of the Vandal Minimum in the 6th century CE. This state prevails into the 21st century. These insights into the deep-time environmental history of the Georgia Coast represent high-resolution reconstructions for the region and establish long-term baseline ecological conditions operating on the coastline of the Southeast U.S.

Materials and Methods

Our tree ringwidth chronology is largely derived from a buried deposit at the state-owned Altamaha Wildlife Management Area (WMA) directly south of Darien, Georgia, at the mouth of the Altamaha River. Tree specimens, the vast majority of which are bald cypress (*Taxodium distichum*), are dredged up from the coastal sediments during maintenance work at the WMA. Most of the specimens used to develop this chronology are from Butler Island, a marsh island within the WMA. Some specimens were obtained from Champney Island, another WMA marsh island. Some modern bald cypress samples were also obtained from Sapelo Island, the closest barrier island to the WMA and closest accessible sampling location to obtain modern comparative specimens (Figure 4.1; see Napora et al. in prep for sample details). The time span covered by each Altamaha cypress tree in this study is shown in Figure 4.2.

The islands of the Altamaha WMA lie within the river's delta, at the center of the Georgia Bight, the curved coastline that extends from Cape Fear, North Carolina to Cape Canaveral, Florida. As the wide, shallow continental shelf in the area increases the tidal amplitude, these islands are subject to dramatic twice-daily tidal shifts up to 3 meters (Champney Island Tides 2021; Hayes 1994: 236–237). On the barrier islands and deltaic river islands, bald cypress trees grow in ponds and sloughs, which function as reservoirs of fresh water. The water level in these habitats fluctuates with the daily tidal cycle as well as throughout the year (Johnson et al. 1974: 50–51). Although mature cypress trees frequently grow in inundated conditions, seeds can only sprout, and saplings become established, during prolonged dry periods (Fowells 1965; Johnson et al. 1974: 51). Planters cleared thousands of acres of these Georgia Coast forests and cypress swamps in

the 18th and 19th centuries to develop plantations based around the labor of enslaved individuals (Bell 2010; Johnson et al. 1974: 8; Wilms 1972).

Previous research on trees in the Southeast U.S., including bald cypress, indicates that precipitation and ringwidth are highly correlated (Stahle 2012). The dataset used for this study, however, derives from within the tidal zone, unlike other publically accessible bald cypress studies, all but one of which (Carr 2010) are much farther inland. Ecological analyses of the Altamaha River estuary report that a complex variety of global and regional climatic factors affect freshwater delivery differently throughout the year (Sheldon & Burd 2014). Due to the complex interacting environmental and climatic phenomena at this location, we hypothesized that salinity and storm impacts were also factors influencing ringwidth for this dataset. Because of the different temporal periods over which various factors influence freshwater delivery, we analyzed precipitation and salinity at different monthly time periods to determine which periods had the highest correlation with yearly tree growth.

We used a historical monthly total precipitation dataset from a station in Brunswick, Georgia, obtained from the Southeast Regional Climate Center (SERCC) (The Southeast Regional Climate Center 2020) to examine the effect of local rainfall on tree ringwidth and growth patterns in the study locale. The dataset covers the years 1895—2012, but gaps in the data meant that only the data from 1906—2011 are included in this analysis.

Salinity data are from the Georgia Coastal Ecosystems Long Term Ecological Research (LTER) Data Portal (Georgia Coastal Ecosystems LTER 2019), part of the National Science Foundation's LTER network. This dataset, GCE3_Hydro, North

Sapelo, represented the closest geographic source of salinity information to the modern comparative sampling location – about two miles north of the cypress stand on Sapelo Island where modern recently felled and standing dead specimens were obtained. This dataset covers the years 2004 – 2016 but is missing data for 2010.

The impact of hurricanes on coastal ecosystems can be profound (e.g., Dame et al. 2000: 803), and previous studies indicate that strong storms affect the growth of tree rings in the Southeast U.S. Two master's theses on pine trees in southern Georgia and Mississippi found significantly narrowed rings in the year after a hurricane (i.e., a cyclone in the North Atlantic or east Pacific [Hyndman & Hyndman 2009]) or tropical cyclone (i.e., a pressure cell originating over warm tropical ocean water [Hyndman & Hyndman 2009]) event (Collins 2014; Tucker 2015). For the sake of brevity, we will hereafter refer to all strong storms as “hurricanes.” To determine whether hurricanes produced an identifiable ringwidth pattern in coastal Georgia bald cypress, we compared known storm histories with chronology tree-ring indices (standardized units representing average ringwidth). Historical information on hurricanes impacting the Georgia Coast from 1750 to 2012 CE was obtained from Welford et al. (2017). Over this period, 24 hurricanes in 23 separate years are known to have made landfall in Georgia or impacted the Georgia coastline. These hurricanes and the relative tree-ring indices for these years in the normalized (i.e., normalization of the detrended series to a mean and standard deviation of one) tree-ring chronology are listed in Table 4.1.

Coastal forests in the past experienced periodic increases in salinity, and associated mortality, from storm surge, sea-level transgression as they do today (e.g., Doyle et al. 2010; Myers 2019; Saha et al. 2011). Based on existing research, we

hypothesized that periods of increased, prolonged salinity influx to coastal forests could be identified through elemental analyses of tree rings. Bald cypress seedlings can survive only short-term salinity levels of approximately 4 parts per thousand (ppt) (Hsueh et al. 2016:1453), while at least some mature trees tolerate up to 8 ppt (Conner and Inabinette 2005). Coastal cypress forests generally transition to marsh when soil salinity continually exceeds 2 ppt, although forests can survive short, infrequent pulses of higher salinity (Hsueh et al. 2016: 1453). Bald cypress exposed to saltwater uptake sodium, chloride, and bromide in detectable levels in the outer rings; the longer the trees are exposed to borderline survivable levels of salt water before they die, the farther back into the trunk wood these elements penetrate (Yanosky et al. 1995). If, however, bald cypress trees survive saltwater intrusion and then experience declines in saline levels, the trees will eliminate these salts. Bald cypresses growing on the Altamaha WMA islands and Sapelo Island are acclimated to twice-daily tidal shifts ranging between 1.5 and 2.5 meters (“Champney Island” 2021; “Old Tower” 2021) and associated shifts in salinity. Salinity-enhancing events like sea-level transgression or lingering storm surge, however, would disrupt trees’ ability to expel accumulated salts by prolonging exposure time to salt water and limiting freshwater intake. Based on the existing cypress salinity studies, we hypothesized that sampling ancient trees at standardized intervals, beginning at the outer rings (for samples containing the final year’s growth or sapwood—the softer, lighter outer wood where vascular functions occur (Speer 2010: 76)) and inward toward the pith, would provide information about whether individual trees succumbed to salinity intrusion as well as relative rates of salinity increase in the coastal forests. Relative rates of salinity increase could in turn suggest relative rates of sea-level transgression. High levels of

sodium, chloride, and bromide concentrated in the outer rings, with low or undetectable levels of these elements farther back into the tree, would suggest that a tree experienced mortality due to relatively rapid salinity increase. If a tree had elevated levels of these elements in the outer rings and detectable though less elevated levels in test locations nearer to the pith, this would suggest that the tree died after relatively prolonged exposure to saltwater (i.e., slower rates of salinity increase that nevertheless were extensive enough to eventually kill the tree). If these elements are not detected in the samples, or were detected in relatively low levels, this would indicate that the tree did not experience salinity-related mortality, though it may have experienced increased or fluctuating salinity levels during its lifespan, but high salt levels were not the cause of death.

For each of the 33 trees that grew prior to 1000 BCE based on our original radiocarbon dates before dendrochronological analysis was completed, we took three samples beginning at the outer extant rings (Salinity Test 1- Outer Rings), followed by a second sample taken 5 cm farther back towards the pith (Salinity Test 2) and a third sample taken 10 cm back towards the pith (Salinity Test 3). In total, 99 samples were thus obtained from these 33 trees.

Samples were prepared and analyzed at the Center for Applied Isotope Studies (CAIS) at the University of Georgia, USA. Each dry sample, method blank, and certified reference material (CRM) was transferred to a pre-weighed PFA digestion vessels (Savillex, USA) and weighed again. Each method blank consisted of 100 μL 18 M Ω deionized water. Each CRM consisted of 100 mg of NIST SRM 1572 Citrus Leaves (NIST, USA). Concentrated, trace metal grade nitric acid (Fisher Scientific, USA) was added to each vessel (3 mL / g sample). The vessels were covered and left to react

overnight at room temperature. An additional volume of acid was added (2 mL / g sample) and the vessels were covered and heated at 95 °C for six hours. After cooling, Suprapur grade 30% hydrogen peroxide (Merck, USA) was added (4 mL / g sample) and the vessels were covered. Once effervescence subsided, the vessels were heated at 95°C for two hours. After cooling, vessels were weighed, and the contents were transferred to 15 mL centrifuge tubes and then diluted to 10 mL with 10 MΩ deionized water. Samples were centrifuged just prior to analysis to remove any remaining particulates.

Samples were analyzed for a range of elements (Al, B, Ba, Br, Ca, Cd, Cl, Co, Cr, Cu, Fe, K, Mg, Mn, Mo, Na, Ni, P, Pb, S, Si, Sr, and Zn) by inductively coupled plasma optical emission spectrometry (ICP-OES) using an Optima 8300 machine (PerkinElmer, USA). Calibration and QC stock solutions for all elements except S were obtained from Inorganic Ventures (USA). Sulfur calibration stock solution was prepared from 99% pure Na₂SO₄ (Sigma-Aldrich, USA) and sulfur QC stock solution was obtained from Aldrich (USA). All calibration and QC solutions were prepared by diluting the stock solutions in 2% w/w HNO₃ into the appropriate mg/kg ranges. The blank was 2 % w/w HNO₃. After calibration, QC solutions were used to check the performance before and after each group of samples were run. All calibrations, QCs, blanks, samples, and CRMs were run in triplicate and averaged. Method blanks were subtracted from the sample results. The instrument detection limit for each element was determined as 3 times the standard deviation of 10 replicates of the blank solution. Concentrations determined by ICP-OES were multiplied by the dilution factors in order to determine the concentration of each element in the original dry sample. We discuss the results of the analysis for the targeted elements (Na, Cl, and Br) in the next section.

Results

Environmental Ringwidth Drivers at the Altamaha River Delta

Precipitation

As expected, precipitation is the main environmental factor associated with ringwidth on the Georgia Coast. The highest correlation of ringwidth is with January – May precipitation (Pearson correlation coefficient (r)= 0.348; this statistic is used in all further discussions of correlation in this article). This indicates that rainfall prior to the growing season (stated variably as either March – August [Stahle et al. 2012] or May – September [Jackson 1952]) impacts annual ringwidth at the mouth of the Altamaha.

Salinity

Correlation between annual precipitation and salinity is -.382. The highest correlation of salinity and ringwidth for the Altamaha coastal chronology is -0.167 for the time period from May – September (the growing season according to Jackson 1952) mean salinity. Meanwhile, January – May salinity has a positive correlation of 0.152 (a likely random correlation, as salinity has a negative impact on baldcypress growth (Allen et al. 1996)), apparently indicating that prior-to-growing-season salinity does not have the same impact on ringwidth as precipitation, but that salinity during the growing season negatively impacts growth. As the time depth of the salinity dataset is poor (only 12 years), however, continued monitoring of monthly salinity levels is necessary for a better understanding of the role of salinity in the annual growth of coastal bald cypress.

Hurricanes

Although hurricanes cannot be reliably identified solely from the ringwidth patterns of trees growing in this location, hurricane impact years and subsequent years are

associated with certain ringwidth patterns. The hurricane season, which runs from June 1st– November 30th in the Atlantic basin (National Hurricane Center and Central Pacific Hurricane Center 2020), overlaps with the late portion of the growing season and thus might be expected to exert an influence over the latewood, the darker, denser part of the annual growth band laid down in late summer and early autumn. Of the 23 years directly impacted by a hurricane on the Georgia Coast between 1750 and 2012 CE (Welford et al. 2017), 7 (30%) show a decreased ring in the year of impact and 14 (61%) show a large ring in the year of impact. 14 (61%) show a decreased ring in the year after impact. In 11 (48%) of these years, the chronology shows a large ring in the year of impact followed by a smaller ring the next year. The smaller ring following the year of impact is a trait shared with pine trees (Collins 2014, Tucker 2015). The high percentage of hurricane years associated with a large ring in the year of impact followed by a smaller ring the next year may indicate a ringwidth signature unique to this species or this location. This pattern may stem from increased rainfall from hurricane/storm activity leading to a burst of latewood development followed by increased salinity from lingering storm surge that negatively affects the next year's early season growth.

Isotopic analyses would very likely help to differentiate hurricane events from precipitation-based ringwidth changes. Low $\delta^{18}\text{O}$, due to the high relative percentage of ^{16}O evaporated from ocean water during hurricanes and subsequently precipitated, is associated with hurricane events (Speer 2010: 243). Pine trees in interior Georgia (Miller et al. 2006) and in Texas (Lewis et al. 2011) record hurricane signatures to a high level of accuracy with sub-annual isotopic testing. Future research on the Altamaha coastal trees should include seasonal oxygen isotope studies of both modern trees which survived

hurricanes at known intervals as well as ancient trees whose ringwidth signatures suggest they may have experienced hurricanes in the past. Such a study could help to distinguish hurricane events from precipitation and/or salinity levels that could cause similar ringwidth patterns.

Periods of Major Fluctuation in Tree-Ring Indices

The normalized tree-ring indices (Napora et al. in prep: Figures 3.3 and 3.4, Tables 3.4 and 3.5) provide insights into environmental conditions over a period of five and a half millennia while best accounting for the fluctuating sample replication.

Two periods in the tree-ring chronology are the most notable for the increased interannual variability and very small and very large annual rings they exhibit. These periods can be characterized as large-scale state shifts to multicentury environmental upheaval. The first of these periods (2355–1863 BCE) is marked by high interannual variability and numerous very small rings; these characteristics do not appear to be due solely to the decreased replication at the central part of this period: even with the series normalized, this period stands out for the environmental fluctuation it indicates, including numerous very dry years as well as a number of very large rings followed by very small rings. The series of extremely small index numbers also continues beyond the period of poor replication, further supporting the interpretation of this period as one of decreased stability and frequently impacted by environmental events. The period from 1086–678 BCE also exhibits high interannual variability as well as very large and very small rings. A third visually similar period between ca. 1451–1958 CE corresponds with the Little Ice Age's intensification in the North Atlantic (Miller et al. 2012), but low replication during much of this period makes interpretation less certain.

Sodium Concentrations

OES test results of the 33 oldest trees from the Altamaha WMA deposits found varying levels of elements in the three samples taken from each tree. Of the three elements of particular interest (Yanosky et al. 1995), only sodium was present in detectable level. Levels of chloride and bromide were too low to be detected with this methodology in every sample. At least one of three tests had elevated sodium levels in 21 of 33 trees (64%) (Table 4.2). We consider “elevated” to be sodium levels >1000 mg/kg (0.1%), the approximate value found in the stems of sapling bald cypress grown in non-saline conditions (Allen et al. 1997: 316).

Our initial concern about this analysis—which has never been reported in the literature for ancient trees—was that trees buried for centuries or millennia in brackish coastal waterlogged anoxic environments might absorb salts after being buried. Thus, salts within these trees might reflect post-depositional conditions rather than environmental conditions during the tree’s lifespan. If it was the case that salts in these trees were incorporated post-burial, we would expect concentrations of salts to be the same throughout a tree and levels of salts in specimens buried in the same deposit in the same locale to be very similar. Instead, we find that sodium levels vary significantly among trees as well as between tests from the same tree. Furthermore, we find several temporal clusters of tree deaths with elevated salts in the outer rings. We interpret these results as confirmation that the detection of these elements is indeed related to saline conditions at the end of a tree’s life rather than to depositional environment and that clustered deaths indicate salinity intrusion events that caused mass mortality among bald cypress trees at the mouth of the Altamaha.

Yanosky et al. 1995 indicates that sodium levels would be most concentrated in the outermost rings in a tree that died of exposure to continuously increasing salinity levels. Our study found that this was not always the case. In 13 cases (62%), the outermost ring was the most elevated in sodium, while in 5 cases (24%), the middle sample was the most elevated. In 3 cases (14%), the innermost sample had the highest concentration of sodium. What these alternative patterns might indicate is uncertain. One possible interpretation is that the tree experienced rapidly salinity influx that was then somewhat slowed or mitigated (through, perhaps, an event like high rainfall), leading to an unsuccessful attempt to expel high concentrations of accumulated salts from the sodium-saturated wood before the tree ultimately succumbed. Such a sequence of events must have occurred within a relatively short time period.

The sodium levels in two trees are extremely elevated compared to the other samples. Tree 15 (outer ring = 2578 BCE) had outer ring sodium levels of over 15000 mg/kg. The outer ring sodium level in Tree 23 (last ring= 2306 BCE) was over 46000 mg/kg. It may be notable that Tree 23 died at the very beginning of the period of increased environmental fluctuation that lasted from c. 2300 to 1850 BCE.

Our analyses found no single year or short period between 3161 BCE and 700 BCE when all or most sodium spikes cluster. There are, however, four periods of 13 years or less in which at least two trees died with elevated sodium levels in at least one of three samples; in two of these cases, another tree also dies with elevated sodium levels within 28 years. These clusters of tree deaths, including the third sodium spikes, occur within the time periods: 1) 2578 – 2542 BCE, 2) 2314 – 2281 BCE, 3) 1233 – 1223 BCE, and 4) 805 to 792 BCE (Figure 4.3). These clusters of sodium spikes in the outer rings of

trees indicate salinity intrusion events, possibly sea-level transgression or periods of increased storm activity where saltwater was pushed into the coastal forests with more frequency than in other periods.

Altamaha Bald Cypress Life Histories

Since most samples from the Altamaha WMA include near-pith rings (years close to the sprouting date of the tree) as well as sapwood (years close to the final year of life of the tree), it is possible to include shifts in tree lifespans in this analysis.

The tree-ring chronology indicates several gaps in bald cypress tree sprouting, which may point to periods in which the environment did not support the flourishing of bald cypress seedlings into more resilient mature trees. Cohorts of trees that sprouted in the same year, with each cohort separated by decades or more, are common in modern cypress forests. This indicates the sensitivity of seedlings and young trees to environmental fluctuations, including flooded or saturated conditions and increased salinity, and the necessity of stable, relatively dry (though not too dry) conditions over a period of decades for bald cypress trees to flourish (Sharitz and Lee 1985).

Discussion

The specific river-mouth location of the field site, a few kilometers upstream of one of the most productive estuaries on the western Atlantic coast but subject to the impact of hurricanes, means that a wider array of environmental factors influence tree growth in this locale than has previously been found for tree-ring studies in the Southeast U.S. The seasonality of significant ringwidth drivers differs: prior-to-growing season and early growing season rainfall impacts ringwidth, while salinity throughout the growing season limits annual growth. Hurricanes, meanwhile, are sometimes (but not always)

discernible via ringwidth patterns, notably appearing as a very large year-of-impact ring and a very small subsequent ring in 48% of modern cases.

This study provides high-resolution, tree ring-based environmental proxies that elucidate the temporal relationships between cultural changes evident in the archaeological record and environmental conditions. Via this multi-proxy dataset, we gain insights into some of the factors which impacted the Georgia Coast over the millennia, including punctuated episodes of stability and change, hurricane events, inundated conditions, and salinity impact on trees, including evidence for sea-level changes.

Studies in the Southeast U.S. use a variety of data sources to develop models of sea level in the ancient past. Garrison et al. (2012) use sediment cores from offshore locations on the Georgia Coast to model paleoshorelines and pollen data to detect shifts in ecological communities during the Late Pleistocene and early Holocene. Sea-level reconstructions of the Atlantic Southeast Coast focused on the later Holocene indicate a generally steadily rising sea over the past several millennia, but presence and timing of sea-level oscillations in these models vary. Engelhart and Horton (2012), employing a comprehensive database for East Coast sea-level points, find more rapidly rising sea-level prior to 2000 BCE, with a slower rate of rise from 2000 BCE to 1900 CE; data points from the Georgia Coast in this model, however, are lacking. Other sea-level reconstructions, meanwhile, find evidence for fluctuations. These include DePratter and Howard (1981), using archaeological site and buried tree locations; Colquhoun and Brooks (1986), based on marsh stratigraphic and archaeological evidence from South Carolina; and Gayes et al. (1992), using marsh foraminifera. These models are

summarized by Turck and Thompson (2016b: Figure 3). Both DePratter and Howard (1981) and Gayes et al. (1992) posit a single sea-level lowstand but at different times. The Colquhoun and Brooks (1986) sea-level model suggests several sea-level fluctuations over the period in question.

Correlation between the bald cypress tree-ring data and existing Late Holocene sea-level models that find evidence for fluctuation (Figure 4.4) as well as the other major insights gained into the ancient coastal environment from this array of analyses, progressing forward temporally from the beginning of the tree-ring chronology, are described as follows:

Possible Multicentury Inundation Followed by Rapid Salinity Intrusion c. 2900 – 2500 BCE

The 355-year gap in bald cypress trees sprouting from 2928 – 2573 BCE appears from the tree-ring indices to be one of relative stability—a period characterized by press conditions. Such a gap in establishment may indicate that forests flooded frequently—an interpretation supported by a dearth of many dry years indicated by the tree-ring chronology. Such flooded conditions would not affect established trees, but would prevent successful seed germination and sapling development (Fowells 1965).

The first of the four clusters of trees with elevated levels of sodium in their outer rings occurs at the very end of this gap in new tree growth (2578 – 2542 BCE). None of the existing Atlantic sea-level models for the later Holocene find evidence of rising seas at this time. Balsillie and Donoghue’s comprehensive Gulf Coast study (2004), however, does find evidence for a sea-level lowstand at c. 2800 – 2500 BCE. It is therefore possible that this was a period of rising sea-level on the Southeast Coast of similar timing

to that in the Gulf that was not detected by the less high-resolution proxies on the Atlantic Coast. It is also a possibility that the trees which died at this time succumbed to a salinity intrusion event such as storm surge.

Environmental Fluctuation and Rapid Salinity Intrusion c. 2350 – 1850 BCE

Beginning c. 2350 BCE, the tree-ring chronology indicates that the growth environment becomes less stable, with more interannual variation and a number of very small and very large rings. The second cluster of trees with elevated sodium levels also dies at the beginning of this period (2314 – 2281 BCE), which corresponds exactly with the beginning of a period of rapid sea-level transgression in the Colquhoun and Brooks (1986) model. As precipitation prior to and during the early part of the growing season is the main ringwidth driver at the Altamaha WMA and Sapelo Island, the variability in the chronology very likely indicates less reliable year-to-year rainfall patterns compared to earlier centuries. In addition, a number of very large rings followed by very small rings punctuate the chronology. Extrapolating from modern data, at least some of these may be indicative of hurricanes or other storms impacting the coast. Taken together, these proxies indicate a major period of environmental upheaval on the coastline of the Southeast commencing with rapid salinity intrusion and characterized by pulse conditions including unpredictable rainfall patterns, numerous drought years, and possible hurricane events.

A Rapid Salinity Intrusion Event c. 1230 – 1220 BCE

The third cluster of trees with elevated sodium in the outer rings occurs in the period 1233 – 1223 BCE. As with the 2314 – 2281 BCE cluster, this corresponds closely with the Colquhoun and Brooks (1986) sea-level model and, paired with the extant data,

supports the presence of another pulse period of rapidly increasing salinity. Meanwhile, some studies indicate enhanced catastrophic storm frequency on the northeastern coast of the Gulf of Mexico beginning around 1400 BCE and continuing until 1000 CE (Liu and Fearn 2000; see also McFadden and Jaeger 2014, Saunders 2010). Increased numbers and strength of storms also may have impacted the Southeast Atlantic coastline and may be the driver behind some of the rapid salinity intrusion events evidenced by the tree rings.

Environmental Fluctuation c. 1086 – 678 BCE

Another period of enhanced environmental fluctuation occurs c. 1086 – 678 BCE. As with the earlier period of instability, it is marked by increased interannual variability as well as a number of very small rings. It likely indicates several centuries of environmental flux.

Rapid Salinity Intrusion c. 800 – 790 BCE

The fourth and final cluster of increased sodium in the outer rings of trees occurs from 805 – 792 BCE. All existing sea-levels models propose a lowstand and subsequent transgression occurring slightly after this time. The tree-ring sodium data likely indicate that the lowstand was slightly earlier than previous models suggest, and that a pulse period of rapidly rising sea level had begun by 800 BCE.

A Shift to a More Variable or More Frequently Impacted Environment post-Vandal

Minimum (after c. 500 CE.)

Climatic factors associated with the Vandal Minimum (c. 500 – 800 CE) impacted the environment in Florida, with persistent dry summers throughout the Vandal Minimum and cooling of winter temperatures throughout most of the period evident in shell and fish otolith oxygen isotopes (Wang et al. 2011, 2013). This study supports the onset of a

negative environmental period shifting the expected lifespans of trees in Georgia at approximately the same time as in Southwest Florida.

Tree lifespans at the mouth of the Altamaha never again reach pre-Vandal Minimum lengths (although sample number are admittedly relatively small). Unlike the gap from 2928 – 2573 BCE, during which tree establishment was apparently unable to occur but through which mature, long-lived trees could persist, the relatively short lifespans and differing years of death after c. 530 CE may indicate ongoing environmental fluctuation, elements of which mature trees could not survive, or an environment characterized by pulse conditions—frequent impacts on coastal ecosystems of negative events capable of causing mortality to mature trees (e.g., hurricanes, storm surge, and/or prolonged or extreme drought).

Enhanced Fluctuation in the Little Ice Age (after c. 1550 CE)

The only discernible exceptions to the short lifespans in the post-c. 500 CE period are Trees 40 and 56, which include 406 and 425 rings, respectively (Tree 40 lived for somewhat longer as the near-pith is not intact). These individuals died around 1505 and 1545 CE, respectively, when the third and final period of enhanced ringwidth variability begins. Although sample replication is especially low in this part of the chronology, and thus the magnitude of enhanced fluctuation difficult to ascertain, the fact that the timing of this episode coincides with the onset of the Little Ice Age, which lasted from approximately the 16th to the 19th centuries (Mann 2002), lends support to an interpretation of increased environmental variability on the Southeast Coast. The continued relatively short lifespans of the trees growing at the mouth of the Altamaha indicate that environmental conditions at the river mouth post- c. 530 CE never returned

to the type of long-term stability required for the establishment and continued growth of mature, long-lived bald cypress at this location. Geochemical sea-surface temperature proxy records indicate significant cooling at the Little Ice Age and multi-decadal to centennial climate perturbations in the Gulf of Mexico and Atlantic (Cronin et al. 2010, 2012; Richey et al. 2007, 2009), conditions which the cypress lifespan data support. Paleoenvironmental records from North America and the Caribbean indicate a shift in tropical cyclone activity occurring c. 1400 CE during the transition out of the Medieval Warm Period (c. 900–1400 CE) (Schmitt et al. 2020); the environmental instability indicated by the cypress tree rings and shortened lifespans could indicate a similar transition to a more frequently storm-impacted environment, but more research is required to assess the factors contributing to enhanced variability on the Georgia Coast.

As indicated by the interplay of environmental factors influencing ringwidth at this coastal location as compared to the much more linear connection of inland cypress trees to the single variable of precipitation, the coastal environment is a less stable growth environment for bald cypress than locales even slightly more upriver and inland. Stands of very old cypress (probably at least 1000 years old) still grow only a few miles upriver, and individuals up to 2500 years have been located in swamps further inland in the Southeast (Stahle et al. 2019). This study suggests that tree age at the mouth of the Altamaha had already declined significantly prior to the large-scale logging of cypress stands that decimated the old-growth forests upriver from Darien in the mid-19th to early 20th centuries (Sullivan 2000).

Conclusion

A multi-proxy study of Altamaha bald cypress trees indicate unstable environmental conditions during several multicentury periods as well as several salinity-enhancing pulse events between 2600 and 700 BCE, likely including at least one rapid sea-level transgression event. Our data also support existing studies that point to several major global climatic shifts impacting the environment of the Southeastern U.S.—namely, the Vandal Minimum and the Little Ice Age. Evidence for the impact of these periods has previously been found in the Gulf of Mexico; this study indicates that their influence was felt on the Georgia Coast.

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Figures and Tables

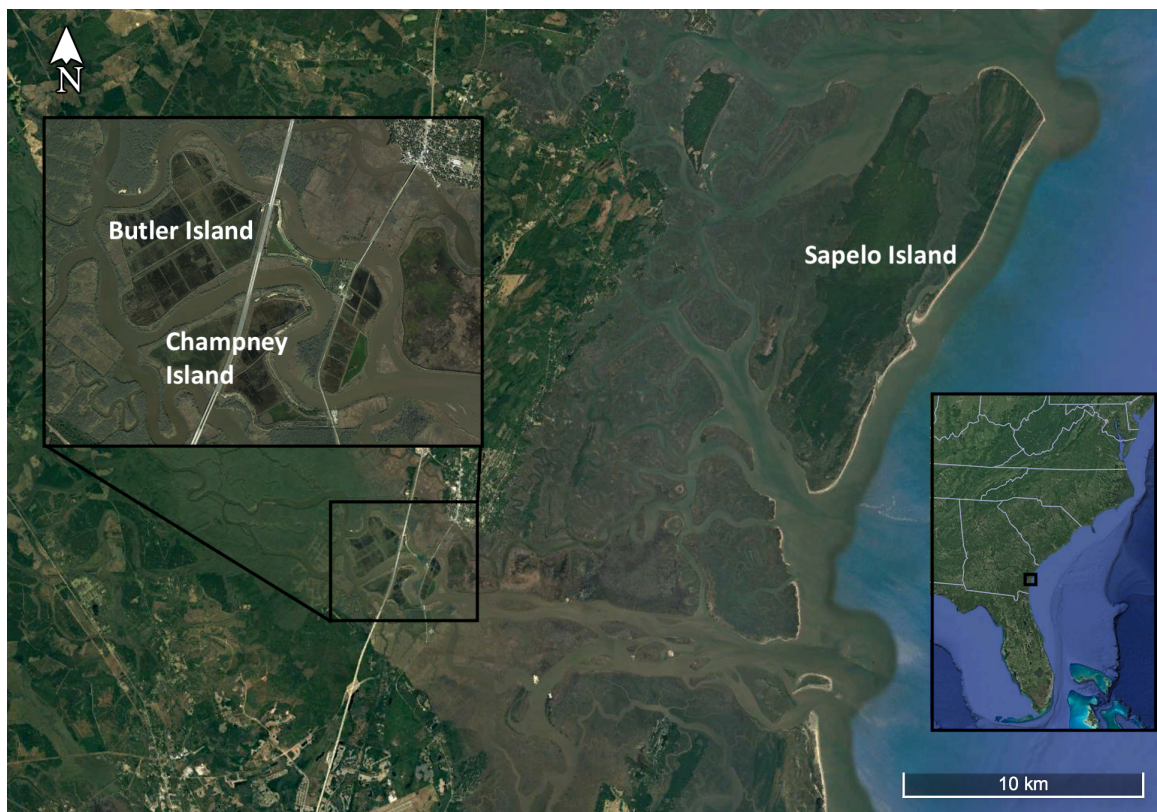


Figure 4.1. Location of Butler and Champney Islands, part of the Altamaha Wildlife Management Area (WMA), as well as Sapelo Island. All tree samples used in this study were recovered from these three locations.

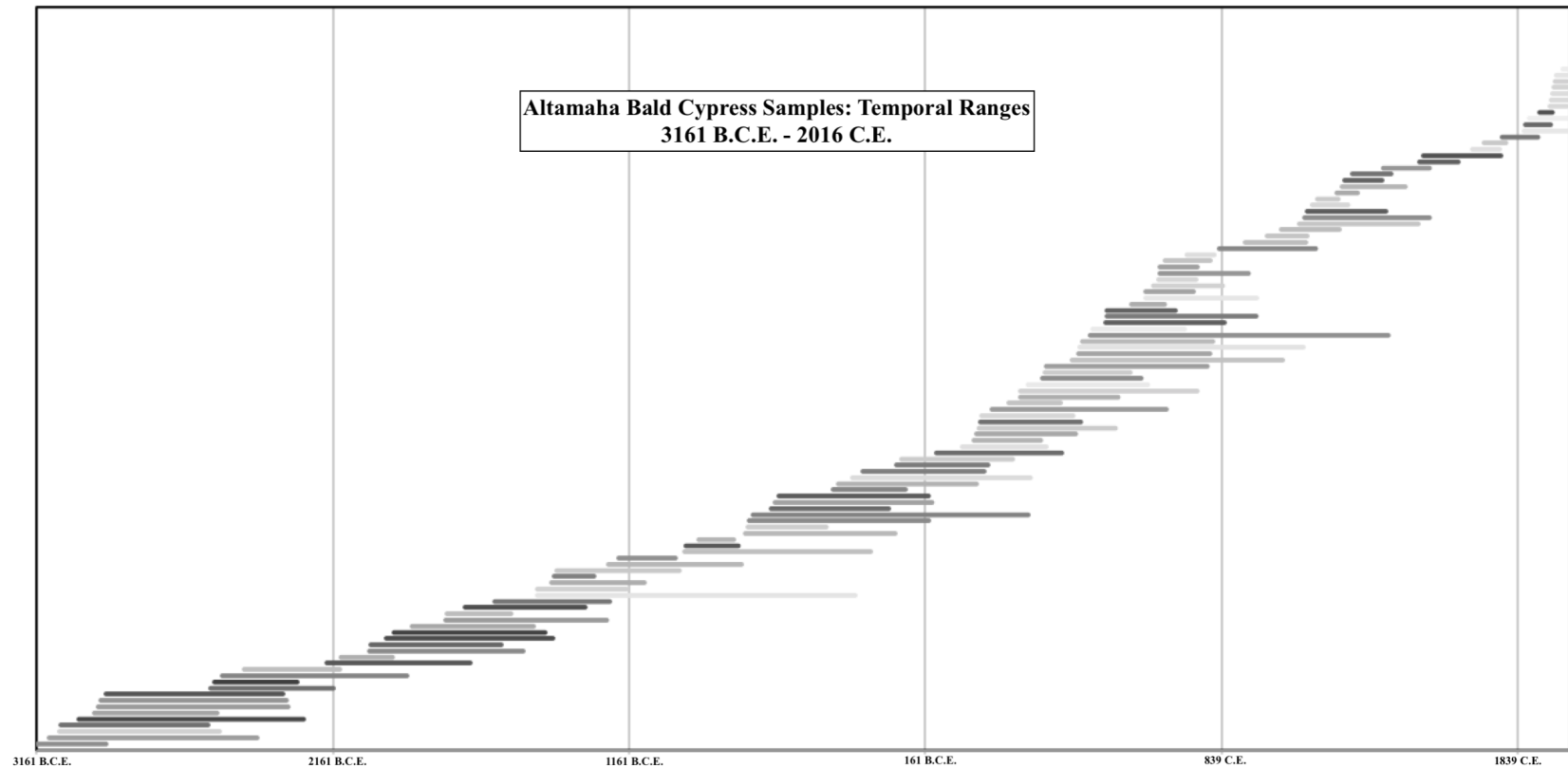


Figure 4.2. Time spans covered by the bald cypress specimens from the Altamaha WMA over the period 3161 B.C.E. to 2016 C.E.

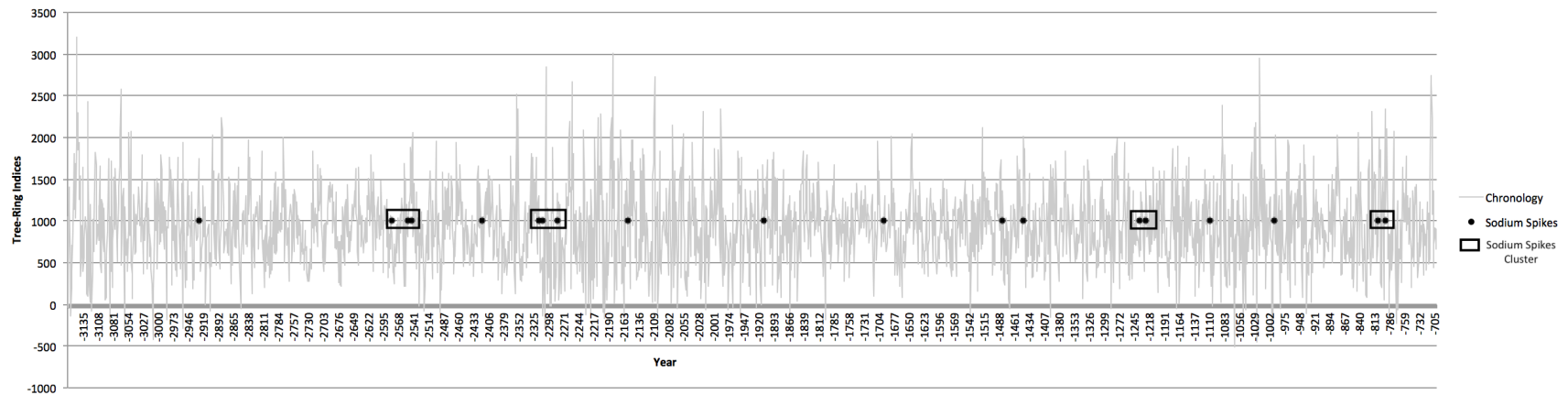


Figure 4.3. Elevated sodium levels (at least 1 of 3 samples for the tree above 1000 mg/kg in Na) in the outer rings of trees that grew from 3161 B.C.E. to 700 B.C.E superimposed against the normalized tree-ring chronology for this time period. Four clusters of sodium spikes, in which multiple trees died in a very short time period with elevated sodium levels in their outer rings, are present.

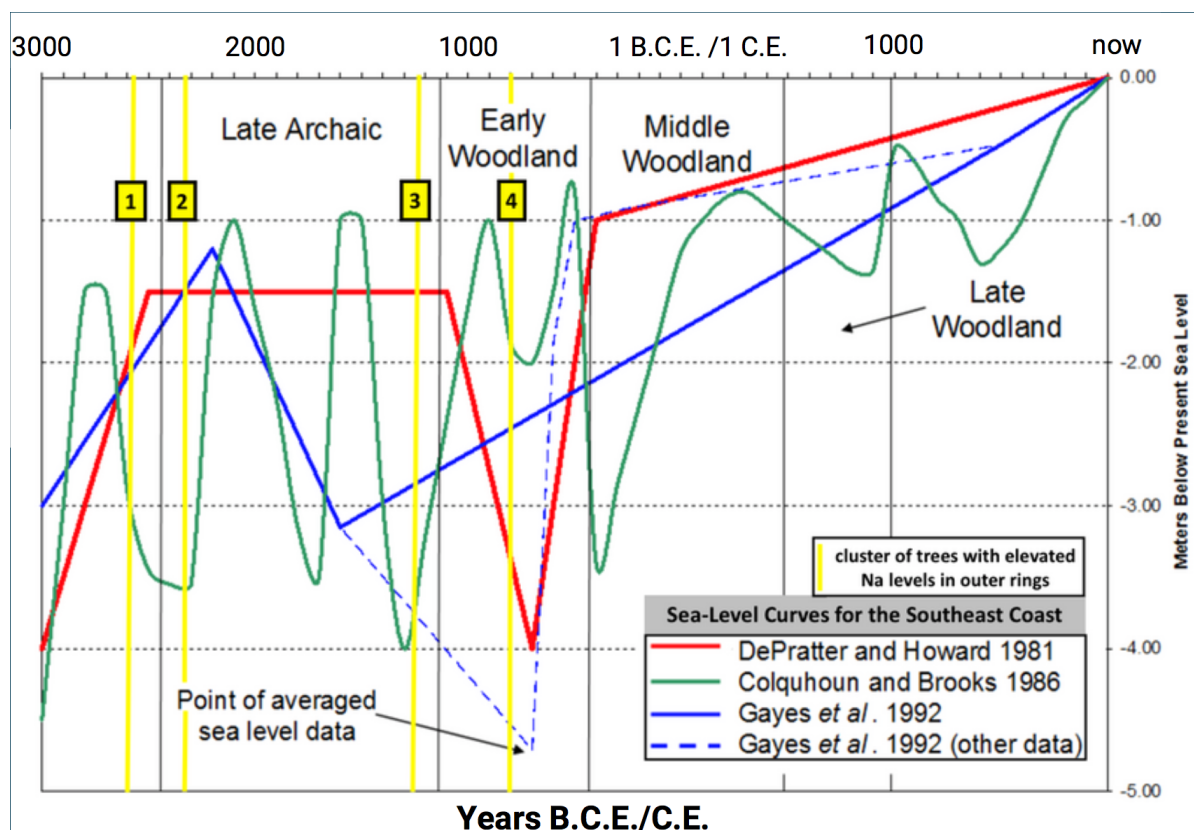


Figure 4.4. The timing of the 4 clusters of sodium spikes shown with the existing sea-level curves for the prehistoric Southeast Atlantic Coast that indicate sea-level fluctuation. The clusters indicate salinity intrusion events that killed multiple trees. Two of these clusters (2 and 3) correlate very closely with the beginning of rising sea-levels in the Colquhoun and Brooks 1986 model and support the timing of this model in these areas. Cluster 1 may indicate a previously unknown sea-level fluctuation, as evidence suggests occurred on the Gulf Coast (Balsillie and Donoghue 2004), or it may be a tree die-off from another type of salinity intrusion event (e.g., storm surge after a hurricane). The timing of Cluster 4 may indicate that sea-level rise began around a century earlier than existing models suggest and thus that the hypothesized lowstand was also approximately 100 years earlier. Figure adapted from Thompson and Turck 2016b, Figure 3.

Table 4.1. Hurricanes that impacted the Georgia Coast over the period 1750 – 2012 and corresponding normalized chronology characteristics. Hurricane information summarized from Welford et al. 2017, pg. 43, Table 2.2.

Hurricane Information			Normalized Tree-Ring Indices		
Year of Impact	Dates	County of Impact	Decreased Year of Impact	Large Year of Impact	Decreased Subsequent Year
1752	Sep. 8 – 16	Unknown			X
1804	Sep. 2 -12	McIntosh		X	X
1811	Oct. 4 – 5	St. Mary's		X	X
1813	Sep. 16 – 17	St. Mary's			X
1817	Aug. 1 – 9	St. Mary's		X	X
1824	Sep. 7 – 15	McIntosh	X		
1830	Aug. 11 – 19	Chatham	X		
1837	Aug. 1 – 7	Duval		X	X
1842	Sep. 30 – Oct. 10	Glynn		X	X
1846	Oct. 5 – 13	Levy		X	X
1853	Oct. 19 – 22	offshore		X	X
1854	Sep. 7 – 12	Chatham		X	X
1878	Sep. 1 – 13	offshore		X	X
1881	Aug. 21 – 29	Liberty	X		X
1885	Aug. 21 – 27	North Georgia		X	
1893	Aug. 15 – 30	Chatham		X	
1896	Sep. 22 – 30	inland		X	X
1898 (2 hurricanes)	Aug. 30 – Sep. 1 Sep. 25 – Oct. 6	North Georgia Camden	X		
1911	Aug. 23 – 30	North Georgia	X		
1928	Sep. 6 – 9	offshore		X	
1940	Aug. 5 – 14	North Georgia	X		
1947	Oct. 9 – 16	Chatham		X	X
1979	Aug. 25 – Sep. 6	McIntosh	X		

Table 4.2. Sodium results from Optical Emission Spectroscopy at the Center from Applied Isotope Studies at UGA for the earliest 35 trees from the Altamaha. Elevated sodium levels are indicated by yellow, orange, or red emphasis corresponding to 1000+, 5000+, and 10,000+ mg/kg, respectively.

SAMPLE ID	YEAR	Sodium (mg/kg)
Tree 19 - Salinity Test 1 - Outer Rings	c. 3475-	4250
Tree 19 - Salinity Test 2	3360 B.C.E.	2465
Tree 19 - Salinity Test 3		1194
Tree 20 - Salinity Test 1 - Outer Rings	c. 3240-	5237
Tree 20 - Salinity Test 2	3160 B.C.E.	2169
Tree 20 - Salinity Test 3		1238
Tree 75 - Salinity Test 1 - Outer Rings	2926 B.C.E.	2748
Tree 75 - Salinity Test 2		385
Tree 75 - Salinity Test 3		748
Tree 15 - Salinity Test 1 - Outer Rings	2578 B.C.E.	15173
Tree 15 - Salinity Test 2		2683
Tree 15 - Salinity Test 3		10260
Tree 12 - Salinity Test 1 - Outer Rings	2550 B.C.E.	4255
Tree 12 - Salinity Test 2		1903
Tree 12 - Salinity Test 3		547
Tree 107 - Salinity Test 1 - Outer Rings	2542 B.C.E.	548
Tree 107 - Salinity Test 2		920
Tree 107 - Salinity Test 3		1525
Tree 82 - Salinity Test 1 - Outer Rings	2416 B.C.E.	2521
Tree 82 - Salinity Test 2		905
Tree 82 - Salinity Test 3		355
Tree 38 - Salinity Test 1 - Outer Rings	2328 B.C.E.	100
Tree 38 - Salinity Test 2		298
Tree 38 - Salinity Test 3		368
Tree 14 - Salinity Test 1 - Outer Rings	2314 B.C.E.	1530
Tree 14 - Salinity Test 2		2049
Tree 14 - Salinity Test 3		1261
Tree 23 - Salinity Test 1 - Outer Rings	2308 B.C.E.	46110
Tree 23 - Salinity Test 2		13251
Tree 23 - Salinity Test 3		2095
Tree 4 - Salinity Test 1 - Outer Rings	2281 B.C.E.	1911
Tree 4 - Salinity Test 2		683
Tree 4 - Salinity Test 3		470
Tree 16 - Salinity Test 1 - Outer Rings	2256 B.C.E.	901
Tree 16 - Salinity Test 2		731
Tree 16 - Salinity Test 3		588
Tree 5 - Salinity Test 1 - Outer Rings	2154 B.C.E.	700
Tree 5 - Salinity Test 2		1327
Tree 5 - Salinity Test 3		809

Tree 86 - Salinity Test 1 - Outer Rings	1958 B.C.E.	682
Tree 86 - Salinity Test 2		64.6
Tree 86 - Salinity Test 3		35.9
Tree 26 - Salinity Test 1 - Outer Rings	1910 B.C.E.	4092
Tree 26 - Salinity Test 2		1805
Tree 26 - Salinity Test 3		1213
Tree 13 - Salinity Test 1 - Outer Rings	1693 B.C.E.	7862
Tree 13 - Salinity Test 2		2408
Tree 13 - Salinity Test 3		395
Tree 35 - Salinity Test 1 - Outer Rings	1592 B.C.E.	598
Tree 35 - Salinity Test 2		518
Tree 35 - Salinity Test 3		502
Tree 84 - Salinity Test 1 - Outer Rings	1559 B.C.E.	945
Tree 84 - Salinity Test 2		777
Tree 84 - Salinity Test 3		865
Tree 33 - Salinity Test 1 - Outer Rings	1516 B.C.E.	678
Tree 33 - Salinity Test 2		316
Tree 33 - Salinity Test 3		273
Tree 74 - Salinity Test 1 - Outer Rings	1480 B.C.E.	2272
Tree 74 - Salinity Test 2		1571
Tree 74 - Salinity Test 3		987
Tree 10 - Salinity Test 1 - Outer Rings	1443 B.C.E.	7042
Tree 10 - Salinity Test 2		2292
Tree 10 - Salinity Test 3		510
Tree 1 - Salinity Test 1 - Outer Rings	1415 B.C.E.	625
Tree 1 - Salinity Test 2		191
Tree 1 - Salinity Test 3		213
Tree 25 - Salinity Test 1 - Outer Rings	1305 B.C.E.	296
Tree 25 - Salinity Test 2		478
Tree 25 - Salinity Test 3		803
Tree 37 - Salinity Test 1 - Outer Rings	1277 B.C.E.	95.3
Tree 37 - Salinity Test 2		134
Tree 37 - Salinity Test 3		138
Tree 6 - Salinity Test 1 - Outer Rings	1233 B.C.E.	436
Tree 6 - Salinity Test 2		1654
Tree 6 - Salinity Test 3		79.4
Tree 24 - Salinity Test 1 - Outer Rings	1223 B.C.E.	523
Tree 24 - Salinity Test 2		3978
Tree 24 - Salinity Test 3		260
Tree 87 - Salinity Test 1 - Outer Rings	1168 B.C.E.	198
Tree 87 - Salinity Test 2		136
Tree 87 - Salinity Test 3		413

Tree 18 - Salinity Test 1 - Outer Rings	1108 B.C.E.	3843
Tree 18 - Salinity Test 2		3686
Tree 18 - Salinity Test 3		2553
Tree 39 - Salinity Test 1 - Outer Rings	1001 B.C.E.	111
Tree 39 - Salinity Test 2		239
Tree 39 - Salinity Test 3		143
Tree 108 - Salinity Test 1 - Outer Rings	991 B.C.E.	1395
Tree 108 - Salinity Test 2		1698
Tree 108 - Salinity Test 3		300
Tree 76 - Salinity Test 1 - Outer Rings	805 B.C.E.	519
Tree 76 - Salinity Test 2		904
Tree 76 - Salinity Test 3		4316
Tree 45 - Salinity Test 1 - Outer Rings	792 B.C.E.	196
Tree 45 - Salinity Test 2		1052
Tree 45 - Salinity Test 3		1695
Tree 27 - Salinity Test 1 - Outer Rings	782 B.C.E.	255
Tree 27 - Salinity Test 2		369
Tree 27 - Salinity Test 3		768

CHAPTER 5

A HIGH-RESOLUTION ANALYSIS OF SETTLEMENT SHIFTS AND
PALEOENVIRONMENTAL CHANGES AT SHELL-BEARING SITES AT THE END
OF THE LATE ARCHAIC (C. 3000 – 1500 BCE) ON THE NORTHERN GEORGIA
COAST, USA³

³Napora, K.G. and V.D. Thompson. To be submitted to *Antiquity*.

Abstract

The Late Archaic period in the Southeast U.S. saw the depopulation of the coastal shell rings inhabited by complex fisher-hunter-gatherer societies. A new tree-ring chronology from the Altamaha River delta provides the framework for interpreting modeled radiocarbon dates from shell-bearing Late Archaic sites on the northern Georgia Coast. Our analysis indicates two distinct settlement periods, with northern sites settled centuries before those farther south. People at all sites weathered a 450-year period (c. 2300 – 1850 BCE) characterized by variable environmental conditions. This study emphasizes the resilience of Native American coastal societies during environmental upheaval.

Keywords: Late Archaic, shell rings, Southeast U.S., climate change, tree rings, Bayesian modeling

Introduction

Late Archaic peoples in the coastal Southeast U.S. (c. 3000 – 1100 BCE) are characterized by complex social formations, sedentism, long-distance exchange, and reliance on estuarine resources (Thompson and Worth 2011). Large sites with arcuate to circular deposits of shellfish, vertebrates, other artifacts, and features were focal points of Late Archaic coastal life. Zooarchaeological evidence and isotopic studies of mollusks indicate that Native Americans used these sites throughout the year (Colaninno 2012; Russo 1998; Thompson and Andrus 2011; Trinkley 1980). By around 1500 BCE, they seem to have abandoned shell rings (Sanger 2010; Thompson and Turck 2009). This apparent shift towards different settlement patterns and perhaps more terrestrial-based economies may have been a response to sea-level changes (Thompson and Turck 2009).

Turck and Thompson (2016) show that the use-life and terminal habitation dates of Late Archaic shell-bearing sites on the Georgia Coast can be divided into two distinct groups based on subregional variation— those located in deltas (i.e., of the Savannah, Ogeechee, Altamaha, and Satilla rivers) and those in non-deltaic areas (for a map of these micro-habitats, see Thompson and Turck 2016: 40, Figure 1). In deltaic areas, large-scale shellfishing began earlier (c. 3000 cal. BCE) and lasted longer (until c. 1500 cal. BCE) than in non-deltaic areas, where intensive shellfishing commenced c. 2500 cal. BCE and ceased by c. 1800 cal. BCE.

In this article, we model new and existing radiocarbon dates from Late Archaic shell-bearing sites on the northern Georgia Coast and interpret them against the environmental proxy data from a multimillennial tree-ring chronology to better understand what environmental changes occurred during the terminal Late Archaic and

whether coastal societies of the period endured through these shifts. We present evidence for large-scale environmental instability during the Late Archaic, including two periods of rapid salinity increase at the Altamaha river mouth. More pronounced environmental fluctuation began c. 2300 BCE and continued for 450 years, overlapping with the use-life of all shell-bearing sites. Our results generally support Turck and Thompson's (2016) conclusion that deltaic shell rings were founded earlier and persisted later than non-deltaic shell rings. Our new dates also indicate that at least some of the Georgia Coast shell-bearing sites were likely occupied for a shorter amount of time than previous studies have indicated. The resilience of both coastal Native American societies as well as the north Georgia estuary ecosystems is evidenced by the use of shell ring sites and estuarine resources for the centuries of environmental instability when numerous drought years and possible strong storms dominated the coastal environment.

Our primary paleoenvironmental proxy for this study is a coastal tree-ring chronology developed from ancient bald cypress (*Taxodium distichum*) trees buried in anoxic conditions at the Altamaha Wildlife Management Area (WMA) in the Altamaha Delta (Napora et al. 2019, in prep). Ringwidth in these trees primarily reflects prior-to-growing season and early growing season precipitation as well as growing season salinity. Hurricanes sometimes also are discernible in the tree rings. In 48% of years in which the Georgia Coast was directly impacted by a hurricane between 1750 and 2012 CE (Welford et al. 2017), the annual ring is very large, and the subsequent year's ring is very small. Sodium concentrations in the outer rings of trees indicate two rapid salinity intrusion events between 2600 and 2200 BCE, likely either periods of enhanced storm frequency or sea-level transgressions (Napora and Jantzi in prep).

Existing archaeological and paleoenvironmental evidence indicates that Late Archaic societies on the Georgia Coast practiced sustainable lifeways on the scale of centuries, and possibly millennia, illustrated by the many persisting traditions and the longstanding cycle of oral transfer of traditional knowledge. Sustainable resource harvesting and continued ecosystem productivity during the occupation of Late Archaic sites bolstered this societal resilience. Oyster shell sizes and salinity-based spatial origins at Late Archaic (and temporally later) sites indicates that Native American peoples in the area engaged in sustainable oyster harvesting over relatively large areas and across very large time frames (Thompson et al. 2020). Systems of resource management and governance based on established fishing territories, as well as the maintenance of inter-village relationships, likely contributed to the long-term sustainability of these groups as well as their estuarine resource bases (Thompson et al. 2020; Turck and Thompson 2016).

While the extant evidence supports the resilience of Native American communities and their sustainable resource usage throughout the Late Archaic, it is evident that there were shifts during the terminal Late Archaic, and onwards into the Early Woodland. People in this region continued to harvest coastal resources even in the face of environmental changes (Turck and Thompson 2016). The transition to the terminal Late Archaic is characterized by an alteration of the spatial layout of settlements between the early Late Archaic and terminal Late Archaic sites, from sites with circular layouts with dense shell deposits to more amorphous and linear forms evidencing more limited use of shellfish. This transition was one of community and societal

reconfiguration, likely set within a shifting ecosystem and resource base (Ritchison et al. 2020; Turck and Thompson 2016).

Here, we establish the resilience of Georgia's coastal Late Archaic shell-bearing site-based societies and illuminate the connection between cultural and environmental changes towards the end of this time period. We model existing and new radiocarbon dates from archaeological sites from this period and compare them to lines of evidence for ecological changes—the multi-proxy tree-ring data covering the same time frame that provides information on year-by-year conditions on the North Georgia Coast. This study addresses the following questions:

- 1) Does tree-ring data indicate environmental stress or upheavals during the time periods which saw the depopulation of the large shell-bearing sites of the Late Archaic?
- 2) To what extent are shifts in settlement during the Late Archaic tied to observable environmental events?

Methods

Radiocarbon Dating

We include in this study nine Late Archaic sites from the Georgia Coast which are geographically close to the paleoenvironmental field site of the Altamaha WMA and, are either relatively well-dated, with radiocarbon dates run on appropriate materials, or they have been excavated and have well-provenienced material available and accessible for dating. North to south, these sites are: Bilbo (9CH4), a shell-bearing, non-ring midden mound at the mouth of the Savannah River, 21 km upriver from the Atlantic (Crook 2009; Waring 1977); Pagan Plum Point (9CH61), another midden mound along the

eastern marsh edge of Skidaway Island (Martin 1980); Odingsell Shell Ring (9CH111), on southern Skidaway Island (DePratter 1975); St. Catherines Shell Ring (9LI231), on western St. Catherines Island (Sanger 2015; Sanger and Thomas 2010); McQueen Shell Ring (9LI648), on eastern St. Catherines Island (Sanger 2015; Sanger and Hurst Thomas 2010); A. Busch Krick Shell Ring (9MC87), on Creighton Island (Crusoe and DePratter 1976; Turck and Thompson 2014); and Sapelo Island Shell Ring Complex (9MC23), a three-ring complex on eastern Sapelo Island (Thompson 2007, 2018; Thompson and Andrus 2011) (Figure 5.1).

Four sites were either previously undated or did not have non-shell dates: these include Pagan Plum Point Midden, Odingsell Shell Ring, A. Busch Krick Shell Ring and Sapelo Shell Ring II. For these sites, we ran new radiocarbon dates from materials in the extant collections held at the University of Georgia Laboratory of Archaeology.

Bayesian Modeling

The new dates and existing selected dates from the Georgia Coast database (Turck and Thompson 2014) were modeled in OxCal Version 4.4 (Bronk Ramsey 2009), employing the IntCal20 Northern Hemisphere curve (Reimer et al. 2020) to develop a chronological framework for the development, occupation, and dates of final usage for Late Archaic shell-bearing sites in geographic proximity to the Altamaha WMA. Existing Late Archaic radiocarbon dates from shell-bearing sites on the northern Georgia Coast were obtained from the Georgia Coastal Ecosystems Long-Term Ecological Research (LTER) online project data repository, a dataset of radiocarbon dates from archaeological sites located in the coastal plain of Georgia (i.e., the portion of the state located south and east of the Fall Line, which marks the greatest advance of the ancient shoreline in the

Mesozoic Era [Johnson et al. 1974]). This dataset incorporated all published and/or publicly available dates from the region and corrected for isotopic fractionation originally uncorrected dates (Turck and Thompson 2014).

Dates on unknown materials, bulk soil samples, and shell were excluded from this analysis. Marine reservoir effects in oyster (*Crassostrea virginica*) are often highly localized and may differ on the scale of centuries (e.g., Hadden and Cherkinsky 2017a; Rick et al. 2012). This renders oyster dates problematic in high-resolution modeling of archaeological chronologies. Mobile gastropod predators like conchs (Strombidae) and whelks (Busyconidae), meanwhile, are able to move between reservoir areas and incorporate various, indeterminate, relative amounts of old carbon, thus making the reservoir effects to which they are subjected even more difficult to specify (Hadden and Cherkinsky 2017b). Even under the best of circumstances, such dates are less precise than terrestrial-based assays (Thompson and Krus 2018).

We incorporated a total of 40 radiocarbon dates in this analysis (Table 5.1). Data for each archaeological site in this analysis were input as a separate Phase into the OxCal model; both dates in sequence and those whose relationship to other dates was unknown were included in these Phases. OxCal's Sequence function was used to model series of dates obtained from materials excavated in stratigraphic order. Such sequences of dates were preferentially selected for when this project selected materials and ran new dates from undated or underdated sites. Start and End Boundaries were inserted for each site in the model to provide date ranges for when occupation at the site most likely commenced and concluded.

Results

The modeled one sigma date ranges (68.3%) (Figure 5.2) indicate that at least some of the Late Archaic shell-bearing sites on the Georgia Coast were settled later and/or depopulated earlier than previous studies have indicated. Modeled start and end dates point to two distinct periods of settlement divided by latitude (Figure 5.3), with more northerly sites settled first, followed centuries later by more southerly sites. This division also correlates to deltaic (Bilbo, Pagan Plum Point, and Odingsell) and non-deltaic (McQueen Shell Ring, A. Busch Krick, and the three at the Sapelo Island Shell Ring complex) microenvironments; St. Catherines shell ring, likely settled in the earlier period, is the sole outlier among the non-deltaic sites. The deltaic sites may have been settled several hundred years later than Turck and Thompson's (2016) analysis indicates, possibly between 2850 and 2285 cal. BCE. This is likely due to the inclusion of shell in the modeling of these dates using the Marine13 calibration curve, which tends to give ranges that are considerably earlier than the most recent Marine20 calibration (Heaton et al. 2020). Thus, our modeling also indicates that the use-lives of the non-deltaic sites may be even shorter than previously believed, with some sites likely having been depopulated by 2000 cal. BCE.

Discussion

Two Distinct Periods of Shell-Bearing Site Settlement Prior to or During Two Major Salinity-Intrusion Events

Modeled date ranges for Late Archaic shell-bearing sites on the northern Georgia Coast indicate that these sites were founded in two periods: the first possibly prior to 2600 BCE (Bilbo, Pagan Plum Point, Odingsell, and St. Catherines) and the second

between 2400 and 2300 BCE (McQueen, A. Busch Krick, and the three Sapelo shell rings).

The settlement of these sites overlaps with two short time periods (2578 – 2542 BCE and 2314 – 2281 BCE) in which multiple cypress trees died of salt uptake in the Altamaha WMA . These two clusters contain the trees with the highest sodium levels in the outer rings of all trees sampled. These two short time periods indicate major salinity intrusion events—either periods of increased storm frequency, during which storm surge into coastal forests caused frequent cypress mortality, or rapid sea-level transgression (Napora and Jantzi in prep). Both possibilities are supported by extant evidence. The Gulf Coast may have experienced increased catastrophic storm activity in the Late Archaic (Liu and Fearn 2000; McFadden and Jaeger 2014; Saunders 2010); the Georgia Coast data may document similar levels of storm activity. At the same time, numerous lines of proxy data support a sea-level lowstand at c. 2800 – 2600 BCE on the Gulf Coast (Balsillie and Donoghue 2004). Meanwhile, one sea-level model for the Atlantic Coast based on marsh stratigraphy and archaeological site data (Colquhoun and Brooks 1986) finds a lowstand coinciding with the second of these events. The latter study, however, proposes more fluctuations than are supported by any other extant sea-level model. If each of these salinity-intrusion events represents a distinct episode of sea-level transgression, their temporal proximity may have made them difficult to distinguish. Colquhoun and Brooks' (1986) data and other proxy data sources used for previous research on sea-level changes in this region, including foraminifera from vibracore samples (Gayes et al. 1992) and archaeological occupational evidence (DePratter and Howard 1981; DePratter and Thompson 2013), may not have had high enough resolution

to identify the first of these two possible lowstands and the following rapid rise, which would have been followed relatively quickly by a second similar event. Rapid sea-level transgression events occurring at both or either of these periods by extrapolation would indicate sea-level lowstands occurring either prior to 2578 BCE and/or between 2542 and 2314 BCE.

Our data indicate that the Late Archaic shell-bearing villages on the northern Georgia Coast were settled in two distinct, latitude-distinguished, periods, each either directly prior to or during major salinity-intrusion events, one (or, less likely, both) of which may represent rapid sea-level transgression. If these events represent sea-level rise, sites were settled during two sea-level lowstand events or directly afterwards as sea-level began to rise. If the events are instead indicators of increased storm frequency, sites were settled either during periods of relative environmental stability or at the beginnings of increased fluctuation.

Enhanced Environmental Fluctuation During the Use-Lives of Shell-Bearing Sites

The Altamaha cypress tree-ring chronology also indicates that each of the two salinity intrusion events occurred at the very beginning of a period of enhanced environmental fluctuation. These periods cover the years c. 2560 – 2490 and c. 2300 – 1850 BCE. The latter 450-year period is particularly notable for numerous drought years as well as enhanced interannual variability, indicating very wet years interspersed with very dry years. Some of these years may also be associated with hurricane or other major storm events. This multicentury period of environmental instability corresponds with the timespan in which the majority of northern Georgia Coast shell-bearing coastal sites were occupied, emphasizing the resilience of Native American coastal hunter-gatherer-fishers.

Native American practices of estuarine resource harvesting and their concomitant social relationships allowed for long-term occupations and a continued stability of cultural practices. This, however, does not mean that Native American systems were unchanging on the coast, as new climate regimes emerged at the end of this period.

Coincidence of the Depopulation of the Last Late Archaic Shell-Bearing Sites with Environmental Downturn

At the end of the use-lives of the last-occupied Late Archaic shell ring and midden-mound villages (post c. 1540 BCE), the tree-ring data show an environmental downturn compared to the period of 1800 – 1540 BCE, which was characterized by a relatively stable coastal environment with more reliable rainfall patterns. The post-1540 BCE period, while not indicating as intense a level of interannual fluctuation as the 2300 – 1850 BCE period, points to an increased number of droughts compared to the preceding centuries. Very small annual rings in the years 1539, 1526, 1483, 1479, 1469, and 1396 BCE indicate either extremely dry years or possible hurricane events in the preceding year.

Conclusions

The settlement of the shell-bearing Late Archaic villages occurred in two distinct periods either directly before or at two major salinity intrusion events. These salinity intrusion events indicate either increased numbers of strong storms, with associated storm surge, or rapid sea-level transgression. The periods before the salinity intrusion events (prior to 2600 BP and 2400 – 2300 BCE) thus represent either times of relative environmental stability or periods of lowered sea-levels. Based on existing sea-level models for the Southeast, it is likely that at least one of these salinity intrusion events is

indeed a rapid sea-level transgression. More northerly archaeological sites in the study area, including the deltaic sites modeled here, were settled first, followed several centuries later by more southern sites. Whether this site settlement pattern speaks more to environmental change experienced differentially by latitude, microenvironment, or other factors is currently unclear. Each salinity intrusion event was followed by a decrease in environmental stability. Given the interconnectedness of Late Archaic villages (Ritchison et al. 2020; Thompson et al. 2020), people farther south on the Georgia Coast were certainly aware that communities living at the existing shell-ring village sites farther north had been able to continue their lifeways during the salinity intrusion event c. 2600 BCE and the following 70 years of environmental upheaval. With this knowledge, more southerly groups may have made the decision to emulate extant, and clearly resilient, communities when the environment again destabilized following another salinity intrusion event c. 2300 BCE.

People living on the northern Georgia Coast inhabited shell-ring sites year-round and focused their economies on estuarine resources during a 450-year period of enhanced environmental instability. The persistence of these permanent, relatively large-scale societies through this long period of environmental instability speaks to the resilience of both the societies themselves as well as that of the northern Georgia Coast estuary ecosystems. The unstable environmental conditions endured from 2300 – 1850 BCE include unreliable year-to-year rainfall patterns, many dry years, and possibly multiple hurricanes or other storm events.

Given the endurance of coastal-focused societies during this long period of environmental upheaval, estuary ecosystems and the resources they provided were likely

less dependent on consistent interannual rainfall patterns, and thus more resilient during this period, than at least some mainland terrestrial resources. Both decreased salinity from heightened multiyear precipitation as well as increased salinity from lasting drought conditions (leading to reduced surface flows from rivers and streams entering estuaries as well as reduced groundwater flows and aquifer recharge) can have major negative impacts on estuary species like oysters, including increased mortality and decreased recruitment (Levinton et al. 2011; Murry et al. 2020). The continued use of the same estuarine resources over this multicentury period of environmental upheaval thus implies a buffering effect of estuaries against year-to-year rainfall variation. The fluctuation in tree-ring indices indicates that, even with frequent droughts occurring, years with high or very high levels of precipitation punctuated the environment regularly; this influx of water was likely enough to replenish groundwater within the extensive aquifer systems that extend along the Southeast Atlantic coastline (“Georgia’s Aquifers” 2002), in turn partially stabilizing estuaries against the impacts of changing salinity levels. As the tree-ring chronology is largely indicative of January to May rainfall (Napora and Jantzi in prep), there may also have been large influxes of water in some years during the summer or fall that are undetectable via the chronology. Such influxes would have served to replenish depleted reservoirs. Studies also indicate that oyster reefs in the region are exceptionally resilient to fluctuations in sea level (Ridge et al. 2017), which suggests that Georgia Coast oyster beds could have persisted through the two possible sea-level shifts of this period.

Subannual oxygen isotopic analyses would help to distinguish between trees killed by sea-level rise and storms. Years in which hurricanes occur show low $\delta^{18}\text{O}$

values in tree rings due to the high relative percentage of ^{16}O evaporated from ocean water during such events and the heavier ^{18}O preferentially precipitated during high-rainfall events (Speer 2010: 243). In an seasonal analysis, low $\delta^{18}\text{O}$ values in the outer rings of a tree with elevated sodium levels would therefore indicate that the tree likely died of hurricane-associated storm surge, while normal $\delta^{18}\text{O}$ would indicate another source of salinity increase, likely rapid sea-level transgression. An oxygen isotope study could help to clarify the timing and number of sea-level lowstands that occurred during this period.

Decreased environmental stability post-1540 BCE combined with a series of annual events that impacted the coast, perhaps hurricanes or other major storms, may have meant that continued habitation of the shell ring and midden mound villages, which seem to have been ideal places to reside during long-term environmental upheaval, was no longer necessary or desirable. It is also likely, given the ability of Native American coastal communities to weather extreme environmental fluctuation, that cultural choices rather than ecological drivers played the major role in societies' decisions to transition aspects of their everyday lives in the terminal Late Archaic.

This study shows that relatively large-scale hunter-gatherer-fisher societies flourished on the northern Georgia Coast during a period of environmental instability lasting from c. 2300 – 1850 BCE that included numerous drought years and unreliable rainfall patterns as well as possible spates of strong storms. Occupation of some of the coastal shell-bearing sites of this period may be shorter than previously indicated. The depopulation of the last occupied shell rings of the region likely occurred during a period punctuated with a series of major ecological events, some of which may have been

hurricanes. Previous research by Thompson and colleagues emphasizes large-scale environmental shifts and the abandonment of shell ring and midden mound villages; however, this picture was based largely on low-resolution environmental data. The new tree ring chronology demonstrates that the relationship of coastal communities to shifting climate and estuarine resources was more complicated than previously thought. This new high-resolution temporal perspective demonstrates the flexibility that such communities possessed in the face of large-scale environmental changes.

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Figures and Tables

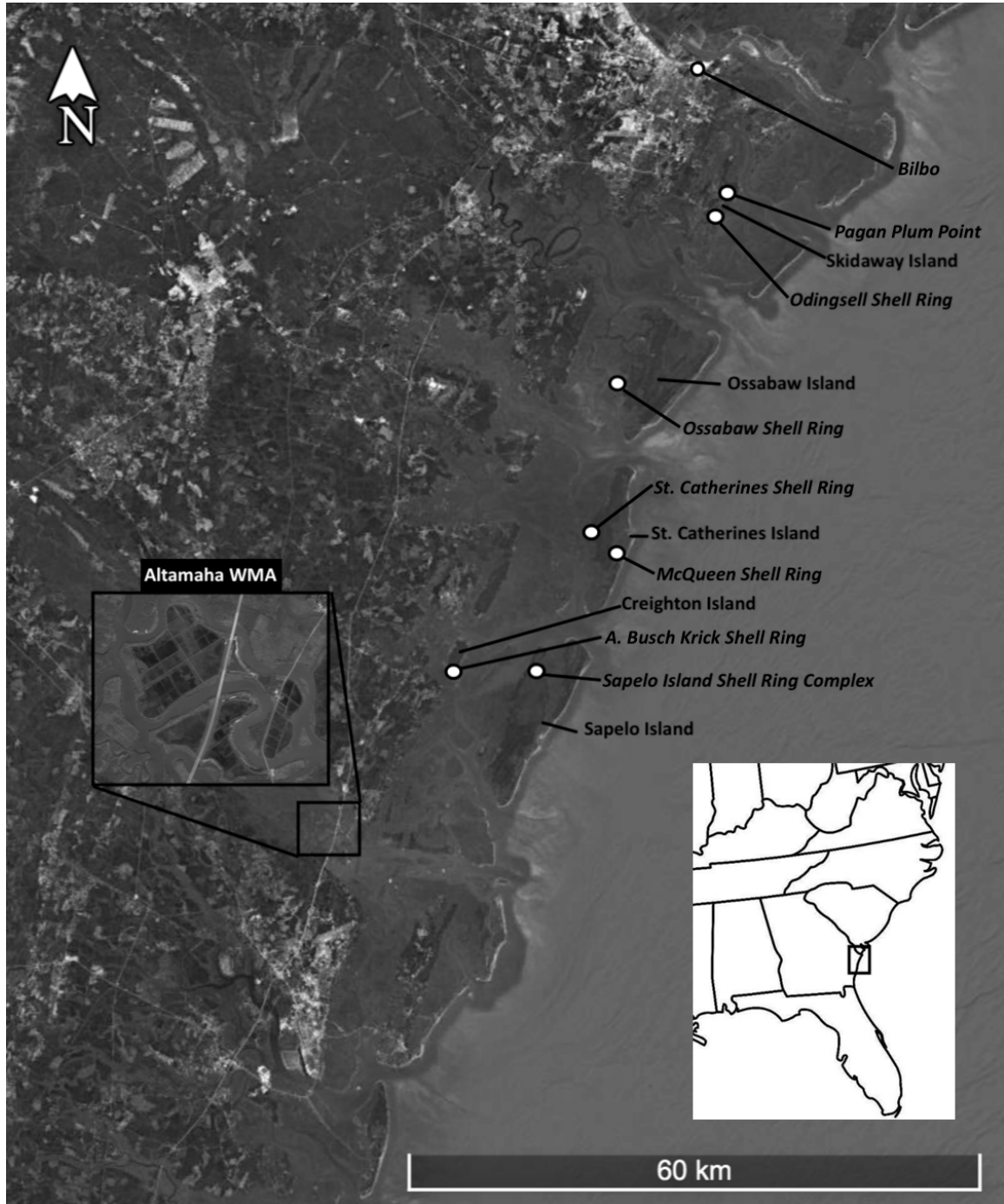


Figure 5.1. The northern Georgia Coast in the Southeast USA, indicating the locations of Late Archaic shell-bearing sites from which radiocarbon dates were run and modeled for this study. Adapted from a Google Earth basemap.

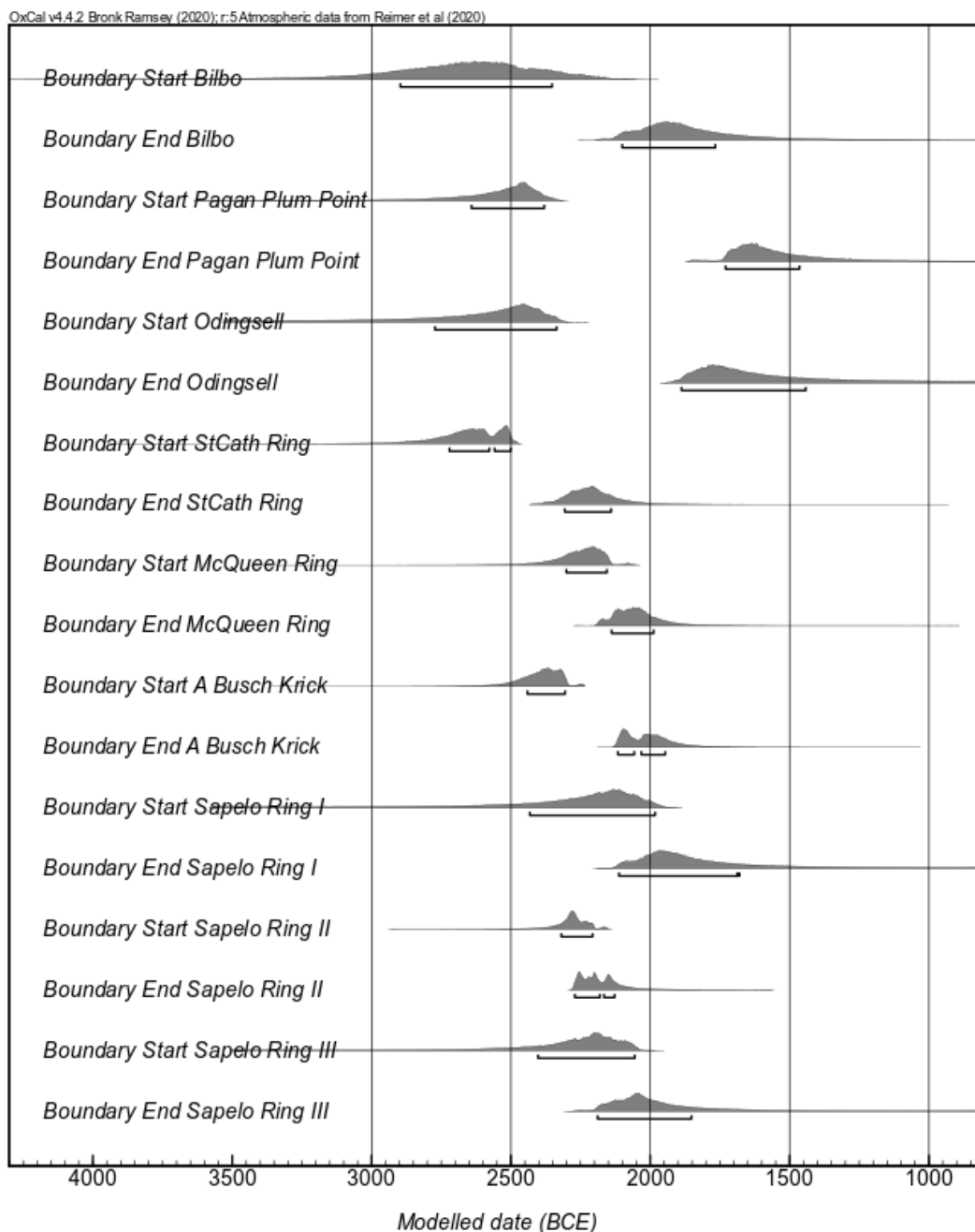


Figure 5.2. Bayesian modeled start and end dates of North Georgia Coast Late Archaic shell-bearing sites (north to south) in OxCal Version 4.4 (Bronk Ramsey 2009), employing the IntCal20 Northern Hemisphere curve (Reimer et al. 2020). Ranges indicated are the 68.3% (one sigma) probability intervals.

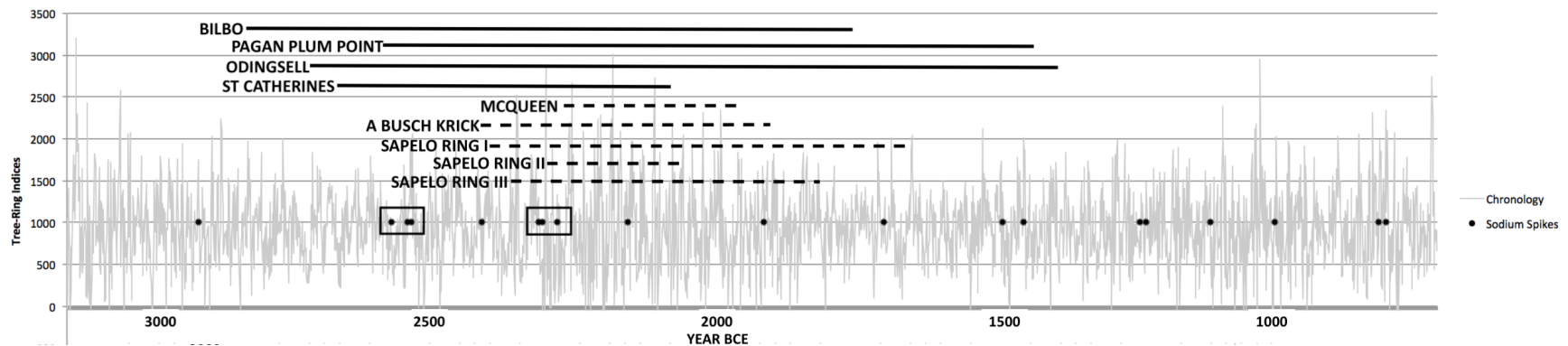


Figure 5.3. Maximum modeled one sigma use-life extents of northern Georgia Coast Late Archaic shell-bearing sites (north to south). Site use-lives are shown against the tree-ring chronology (units = standardized indices, largely indicative of winter-spring precipitation). 2300 – 1850 BCE is characterized by increased interannual variability and numerous very dry years. Also shown are the final growth years of trees that died from salinity intrusion; two clusters of such deaths are emphasized in black boxes and indicate major salinity intrusion events.

Table 5.1. Calibration and modeling of radiocarbon dates of shell-bearing Late Archaic archaeological sites on the North Georgia Coast conducted in OxCal Version 4.4 (Bronk Ramsey 2009), employing the IntCal20 Northern Hemisphere curve (Reimer et al. 2020). New dates are highlighted in black. Information on existing dates was obtained from Turck and Thompson (2014).

SITE INFORMATION		SAMPLE ID	14C AGE (years BP)	±	d13C	2 σ Unmodelled Date Ranges (BCE)		2 σ Modelled Date Ranges (BCE)		1 σ Modelled Date Ranges (BCE)		MATERIAL	SAMPLE CONTEXT (FROM ORIGINAL EXCAVATION)		
SITE: BILBO							from	to	from	to	from	to			
Site Number: 9C114		M-1109	3700	129	-25	Boundary Start Bilbo:		-2470	-1750	-2470	-2900	-2350			
Location: southern bank, mouth of Savannah River		M-1111	3820	129	-25			-2625	-1890	-2470	-2290	-1980 charred wood	Bilbo: 3.0-3.5 ft. bs (Zone 3, Waring)		
Site Type: shell mound, non-ring, deltaic		O-1046	5550	119	-25	(possibly prior to site use-life or old wood)						charcoal	Bilbo: 5.5-6.0 ft. bs (Zone 1, Waring)		
Number of Dates: 6 (5 included in model)		O-1047	4125	119	-25			-3015	-2340	-2880	-2145	-2710	Complex lenses of shell midden and alluvial sand 2.0 to 2.5 feet bs, Haag Zones 7-8/ Waring Zone 3 (3-3.5 feet above msl)		
		UGA-10676	3630	50	-25			-2190	-1830	-2195	-1890	-2135	Nearly solid oyster midden, 5.5 to 6.0 feet bs, Haag Zone 4 (0-0.5 ft below msl)		
		UGA-10677	3730	140	-25			-2565	-1750	-2490	-1960	-2340	Test Pit #14- Upper limit of undisturbed shell midden, 70-80cm below datum		
						Boundary End Bilbo:		-2190	-1240	-2105	-1765		Test Pit #14- Base of dense shell deposit, 170-180cm below datum		
SITE: PAGAN PLUM POINT							from	to	from	to	from	to			
Site Number: 9C1161		UGAMS#39510	3400	20	-22.62	Boundary Start Pagan Plum Point:		-3175	-2375	-2320	-2645	-2380			
Location: eastern marsh edge, Skidaway Island		UGAMS#39511	3790	20	-26.25			-1745	-1865	-1745	-1745	-1640 large mammal longbone fragment cf. deer	Pit 1, Level 1, 0-5' BS, 22076		
Site Type: shell mound, non-ring, deltaic		UGAMS#39512	3870	20	-22.67			-2295	-2140	-2290	-2285	-2145 carbonized wood (<i>Pinus</i> sp.)	Pit 1, Level 4, 2.1' BS		
Number of Dates: 5		UGAMS#39513	3890	20	-25.97			-2465	-2295	-2440	-2310	-2310 carbonized wood (<i>Pinus</i> sp.)	Pit 1, Level 5, 1.5-2.0' BS, 22080		
		UGAMS#39514	3870	20	-22.51			-2460	-2235	-2465	-2330	-2460	Pit 1, Level 5, 2.5' BS		
						Boundary End Pagan Plum Point:		-1860	-925	-1730	-1465		Pit 1, Level 6, 2.5-3.0' BS, 22081		
SITE: ODINGSSELL							from	to	from	to	from	to			
Site Number: 9C1111		UGAMS#39523	3510	30	-22.2	Boundary Start Odingsell:		-3530	-2315	-2315	-2775	-2335			
Location: southern Skidaway Island		UGAMS#39524	3890	20	-26.29			-1925	-1745	-1930	-1895	-1770 large mammal longbone fragment cf. deer	6-12" BS, soil sample rough sorted, 21630		
Site Type: shell ring, deltaic		UGAMS#39525	3860	20	-25.77			-2465	-2295	-2440	-2235	-2385	Ph Lewis Site 29, Level 5, 24"-30" NE, 21634		
Number of Dates: 3						Boundary End Odingsell:		-1920	-640	-1890	-1440		Ph Lewis Site 29, Level 1, c. 34" BS, 21637		
SITE: ST CATHERINES SHELL RING							from	to	from	to	from	to			
Site Number: 9L231		Beta-238322	3870	40	-25.7	Boundary Start St. Catharines:		-2920	-2475	-2475	-2725	-2500			
Location: western St. Catharines Island		Beta-238327	3820	40	-24.2			-2465	-2205	-2470	-2210	-2455	Feature 60, 2.0-1.9 m		
Site Type: shell ring, non-deltaic		Beta-238328	4120	40	-24.5			-2875	-2570	-2765	-2465	-2655	W92 S2, 2.3-2.2 m		
Number of Dates: 7		Beta-238331	3820	40	-25.4			-2455	-2140	-2460	-2200	-2405	Feature 76, 1.9-1.8 m		
		Beta-238332	3880	40	-26.0			-2470	-2205	-2470	-2210	-2460	Feature 88, 1.8-1.7 m		
		Beta-238337	3860	40	-26.8			-2465	-2200	-2465	-2210	-2455	Feature 73, 1.8-1.7 m		
		Beta-239276	3930	40	-25.0			-2570	-2290	-2565	-2290	-2475	N771 E819, 2.39-2.3		
						Boundary End St. Catharines:		-2400	-1975	-2310	-2140		Feature 82, NE Quad 1.9-1.8 m		
SITE: MCQUEEN SHELL RING							from	to	from	to	from	to			
Site Number: 9L648		Beta-244620	3800	40	-25.3	Boundary Start McQueen:		-2495	-2060	-2060	-2305	-2155			
Location: eastern St. Catharines Island		Beta-244745	6060	40	-24.4	(possibly prior to site use-life or old wood)		-2455	-2050	-2305	-2235	-2135	charred material	Feature 21, 4.0-3.9 m	
Site Type: shell ring, non-deltaic		Beta-251761	3720	40	-23.9			-2280	-1975	-2270	-2025	-2200	charred material	Feature 19N, 4.0-3.9 m	
Number of Dates: 6 (5 included in model)		Beta-251764	3710	40	-25.0			-2275	-1970	-2205	-2030	-2195	-2065	charred material	N243 E233, 4.5-4.4 m
		Beta-251766	3800	40	-27.5			-2455	-2050	-2455	-2055	-2235	-2140	charred material	N272 E200, 5.3-5.2 m
		Beta-251767	3680	40	-24.8			-2200	-1945	-2205	-1985	-2200	-2070	charred material	N272 E200, 5.1-5.0 m
						Boundary End McQueen:		-2200	-1840	-2140	-1985			N243 E233, 4.4-4.3 m (shell)	
SITE: A. BUSCH KRICK							from	to	from	to	from	to			
Site Number: 9MC87		UGAMS#39515	3650	20	-22.57	Boundary Start A. Busch Krick:		-2595	...	...	-2445	-2305			
Location: Creighton Island		UGAMS#39516	3810	20	-26.02			-2135	-1945	-2135	-1955	-2130	-1985	deer scapula glenoid cavity	Pit 1, Zone 2 (Level 2), shell sample, Level 4 BS 0.9-1.2 BS
Site Type: shell ring, non-deltaic		UGAMS#39517	3710	20	-26.02			-2200	-2030	-2190	-2030	-2150	-2050	carbonized wood (<i>Pinus</i> sp.)	Pit 1, Soil-shell area, Southern 5', Level 8, 2.1'-2.4' BS
Number of Dates: 8 (7 included in model)		UGAMS#39518	3810	20	-26.03			-2340	-2145	-2240	-2135	-2170	-2145	deer longbone fragment	Pit 1, Level 12, 3.3-3.6 BS, SO 1/3, #22103
		UGAMS#39519	3770	20	-26.03			-2290	-2275	-2275	-2145	-2190	-2150	carbonized wood (<i>Pinus</i> sp.)	Pit 1, LVL 14 at 4'1"
		UGAMS#39520	2760	35	-23.68	(possibly translocation/bioturbation)							mammal bone fragment; cf. deer	Pit 1, Level 16, 4.5-4.8 BS, Zone 3, Black midden	
		UGAMS#39521	3750	20	-26.27			-2280	-2040	-2285	-2155	-2205	-2170	carbonized wood (<i>Pinus</i> sp.)	Pit 1, LVL 16 at 4'8"
		UGAMS#39522	3800	20	-25.11			-2300	-2140	-2300	-2195	-2290	-2200	carbonized pigunt hickory shell	Pit 1, LVL 17 at 5'3", Zone 3
		UGAMS#39522	3900	20	-24.77			-2470	-2295	-2455	-2235	-2385	-2290	carbonized wood (<i>Pinus</i> sp.)	Pit 1, LVL 20 at 5.9
						Boundary End A. Busch Krick:		-2130	-1815	-2120	-1945				
SITE: SAPELO RING I							from	to	from	to	from	to			
Site Number: 9MC23		UGA-15084	3610	50	-17.0	Boundary Start Sapelo Ring I:		-3580	-1945	-1945	-2435	-1980			
Location: Sapelo Island		UGA-15085	3730	60	-18.9			-2140	-1775	-2145	-1830	-2125	-1925	sooted sherd	Sapelo Ring I: Unit 1, Level 2
Site Type: shell ring, non-deltaic								-2340	-1945	-2285	-1940	-2195	-1980	sooted sherd	Sapelo Ring I: Unit 1, Level 2
Number of Dates: 2						Boundary End Sapelo Ring I:		-2135	-615	-2115	-1680				
SITE: SAPELO RING II							from	to	from	to	from	to			
Site Number: 9MC23		UGAMS#42750	3800	20	-25.75	Boundary Start Sapelo Ring II:		-2525	-2150	-2150	-2320	-2205			
Location: Sapelo Island		UGAMS#42751	3770	20	-26.07			-2300	-2140	-2280	-2140	-2275	-2145	carbonized wood (<i>Pinus</i> sp.)	Unit 1, 37.5 cmbd
Site Type: shell ring, non-deltaic		UGAMS#42752	3810	20	-26.99			-2290	-2060	-2285	-2150	-2280	-2160	carbonized wood (<i>Pinus</i> sp.)	Unit 1, 47 cmbd
Number of Dates: 4 (3 included in model)		UGAMS#42753	6220	25	-26.24	(likely prior to site use-life)		-2340	-2145	-2305	-2150	-2295	-2200	carbonized wood (<i>Pinus</i> sp.)	Unit 1, 70-75 cmbd
													carbonized wood (<i>Pinus</i> sp.)	Unit 1, 80-85 cmbd	
						Boundary End Sapelo Ring II:		-2285	-1930	-2275	-2125				
SITE: SAPELO RING III							from	to	from	to	from	to			
Site Number: 9MC23		UGA-15082	3560	50	-27.5	Boundary Start Sapelo Ring III:		-3500	-2030	-2030	-2405	-2050			
Location: Sapelo Island		UGA-15083	3730	60	-25.5			-2340	-2035	-2175	-2015	-1775	charcoal	Sapelo Ring III: Unit 9, Level 4	
Site Type: shell ring, non-deltaic		UGA-15086	3730	50	-25.6			-2295	-1970	-2285	-1980	-2200	-2040	charcoal	Sapelo Ring III: Unit 9, Level 7
Number of Dates: 3						Boundary End Sapelo Ring III:		-2270	-770	-2190	-1850			Sapelo Ring III: Unit 11, Level 4	

Outlier Dates Removed from Analysis

CHAPTER 6

CONCLUSION

This dissertation presents a new 5,177-year tree-ring chronology, as well as an earlier 529-year floating chronology, developed from bald cypress (*Taxodium distichum*) samples buried in the anoxic coastal sediments at the Altamaha Wildlife Management Area. We use this chronology, and associated proxy data (information on environmental drivers of ringwidth, elemental analysis, and tree life histories) to interpret environmental conditions on the Georgia Coast over five millennia. Pairing the tree-ring proxy data with Bayesian modeling of dates from Late Archaic shell-bearing archaeological sites, we also interpret the terminal portion of this period in light of the environmental data.

Insights into long-term environmental history gleaned from the tree-ring data presented here include further evidence pointing to the impact of several major global climatic shifts in the Southeastern U.S. and support for the sea-level model proposed by Colquhoun and Brooks (1986), likely indicating that rather than a single episode of sea-level rise and fall during the Late Archaic, the Southeast coast was subject to several sea-level fluctuations.

Focusing in on the Late Archaic period on the Georgia Coast, we show that people living at the shell rings of the North Georgia coast were able to weather a period of apparently extreme environmental downturn lasting approximately 450 years (2300 – 1850 B.C.E.) that included numerous drought years and unreliable year-to-year rainfall patterns as well as possible hurricane events. The final occupations of these coastal shell sites came after a period of restored environmental stability but may have been

influenced by several years of very negative conditions. The resilience of both the Indigenous peoples of coastal Georgia's Late Archaic as well as the estuary systems of the Southeast in the face of a multicentury environmental downturn is evident from this study.

This study indicates that coastal estuaries can be extremely resilient during episodes of major environmental fluctuation. In our own period of increasing climatic instability, studies like this one which may point to areas where certain resources could continue to be utilized should continue to be researched and, from there, operationalized to help develop more sustainable harvesting practices so that places which in the past have proved resilient can continue to be so in the future. For instance, the fact that the more northerly and deltaic-situated shell-bearing sites were used over much longer periods of time, including during a long, extremely negative environmental period, may be indicative of the greater ability of these specific areas to weather a wider variety of ecological events over a longer period. Understanding differences in locales could direct resource management organizations to preferentially select these areas for projects such as oyster bed restoration and reforestation of cypress forests with more salt-tolerant individuals.

The findings of this dissertation, however, may also speak, to the impact of even individual negative events like hurricanes or extreme drought years on the lifeways of coastal societies. While people living at the shell rings of the Georgia Coast remained there over the 450-year period of environmental turbulence, they made the decision to reconfigure their way of life during a period that appears to have been relatively stable from an environmental standpoint, but punctuated by several very negative years. Was

the decision to shift to new ways of living impacted by what appears to have been only a few events? Perhaps these very negative years indicate hurricanes that would have devastated coastal communities, and direct hits by a series of such storms could have contributed to the decision of people living in the terminal Late Archaic to change their established ways of life. Today, the stability of coastal ecosystems is threatened by anthropogenically amplified climate change, leading to changes that include less predictable environmental patterns and a higher frequency of hurricanes and storms. If a few, perhaps catastrophic, environmental events contributed to the decision three and a half thousand years ago to shift from a way of life that had endured sustainably for centuries, what might a series of such events do to our own coastal societies, which are already struggling with resource depletion, environmental degradation, and the effects of rising and warming global seas? What would it take to push 21st century coastal societies over the edge?

Our continuously increasing knowledge of lifeways and environment thousands of years ago on the Georgia Coast sounds a warning across the millennia to us today. It is imperative that we do everything in our power to slow human-enhanced climate change, the major driver behind the increasingly frequency and amplitude of storms, the more rapidly rising and heating seas, and the destabilization of the global environment—those extreme events that could lead to the a necessary decision to change our own ways of life.

APPENDICES

APPENDIX A

Raw Tucson compact format ringwidth measurements of Altamaha floating chronology.

As chronology is unanchored, the inner “dates” are not dates, but ring

numbers.

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140=N      16=I ALT019A      -3(20F4.0)~
394 6181067 562 225 254 243 477 7301040 505 59010671291 646 758 983 9831126 815
815 7581073 788 9331236 814 3651123 197 618140433962499 9541488 814 505 815 815
702 1971291 89911061416 813100017082041 959 917 916 792 647 5621250 813 562 896
2187120914593104 9801062135510211229 729 709 709 687 604 97915841250 93817711541
10831375 625 937 70912501125 876 625 500 709 8121688 896158310211792187611251690
166720001333 647 667 854 772114613332042179417701000 751 750137311121345 752 879
131943481681 8851345196817332509 9312797279311642509 750300015271656116410872276

293=N      1=I ALT019B      -3(20F4.0)~
825 826 226 253 844 454 466 398 582 181 316 687 416 256 796 802 849 885 822 534
473 199 556 8051094 728 6481235 808 505 97013551293200016711570 747 731 5561218
1399 922 5171017 536 861128235222278 8041710 4631152 7121438 274 677127117121378
1448 757 8311602 8561021 8311070 818 653 76216861164 596 612 918 748 2921273 564
6471258 255 532 292 392 574 240 95815231517 739 958 667 990 905 627 605 605 715
672 779 495 539 280 645 645 648 843 8001053 968 800 349 5391079 984 445 421 920
1173 9841777 888 920139719671399 381 476 85911741205 888 603 9831174 920 6031174
934 935 768 85116191122 86818831287242922661997200013981007 756 615 137 628 481
217 659 461 722 369 812 90 121 628 867 545 635 9941081 3851235 153 946 960 226
665 920147413581872 487105112081210 8661043 137 548 459 864 158 8501299 8231233
831 89 789 621 621 442 757 9521016 132 01233 842 742 6851187 753 4811259 933
672 351 985 817 991003 673 252 764 740 632 754 102 94711871031 677 510 714 550
224 183 28611421020 1041086 614 616133312731001 535 827 947 654 636 530 591 763
778 653 636 429 495 224 820 435 538 388 207 719 553 148 455 523 702 124 455 212
355 2771006 786 3221045 702 762 399 115 730 7941133

314=N      132=I ALT020A      -3(20F4.0)~
208017901318 7031378 938 4401027210923741465 850 70415531699 9121143178826671289
2052102611721820 998170313781202 615 79119552168264228702432167010851348 9101670
8201115 792 61510551172216917291113 733 879 41026962667205120511261225624031875
019931086 8211143 93820811436108822572432243223151700 480 739 652108815221264
1522 8261088 784 871 9131264 871 348 522 609 567 871 480 610 609 827 913 3171958
16091158 550 840 579 767 579 246 189 205 145 507 347 536 0 781 275 710 709 478
1057 536 449 405 868 463 8841172 984 8541014 580 62211601317 130 434 695 897 781
898 723 01013 868 507 767 376 840 190 666 260 174 203 275 463 492 680 334 623
449 145 348 145 217 666 782 696 666 550 839 781 883 3911013 14514621838 709 927
1331 810 782 3331465 52612001010 106 263 253 600 8841084 947 506 463 842 568 831
138 842 695 716 926 810 474 253 737 149 537 705 495 800 453 547 747 463 558 674
106 516 474 706 368 642 389 232 137 1611042 684 5531348 653 454 170 199 99 199
85 355 369 143 355 411 356 284 100 66813761475 979 795 638 426 397 369 383 312
143 270 143 114 128 242 114 114 85 170 100 553 234 281 111 391 142 173 109 234
422 564 453 235 266 80 547 95 234 80 501 517 109 470 563 250 501 173 234 281
219 266 80 328 422 80 485 548 236 0 188 447 254 190

357=N      173=I ALT020B      -3(20F4.0)~
886 735 791 998 807 808 520 507 194 802124822711135105010421027233319271588 206
214812741096 835 927194016261226180417522055176710061317109314041322 875 7481380
14521357 5751187 97018481278 798 856 451 406 901 365 261 423 8021295 011341019
336 538 599 222 510 669 392 757 508 0 662 554 570 99 863 234 786 861 6421239
832 437 208 348 293 6401494 421 800 987 714 77013491074 465 9281267 0 686 927
752 1191143 742 711 798 124 309 0 391 185 515 433 474 150 989 721 310 702 907
1113 948 4741030 640 763 5991072 865 8461052107310111669 496 968 9071362 6181320
721 846 763 681 577 6391113 580 24715681465 3091220 37112361133 8091113 762 577
391 556 5561071 907 618 5361236 846 323 298 153 375 145 432 0 54 340 537 357
832 400 457 483 764 859 577 238 93 203 635 152 804 271 76 330 237 68 68 51
51 226 68 85 68 42 25 68 51 68 85 503 68 127 280 76 128 279 93
127 357 754 644 753 457 144 42 381 51 51 135 152 51 152 440 76 330 212 85
482 325 179 595 97 341 0 447 106 180 81 163 130 293 74 97 0 146 65 196
163 528 870 561 106 636 512 479 81 252 406 512 227 934 138 123 238 115 562 792
123 562 108 492 396 285 108 531 385 625 146 94 346 118 177 93 54 370 138 46
364 479 860 131 138 69 392 61 39 46 474 92 283 515 619 308 688 523 476 382
417 316 215 148 260 440 346 223 226 262 78 494 672 708 807 811 363 683 568 557
540 522 71 319 189 234 266 394 188 777 777 483 235 386 331 260 486

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APPENDIX B

Raw Tucson compact format ringwidth measurements of Altamaha main coastal chronology (3161 B.C.E. – 2016 C.E.); decadal format inserts a “Year 0.” Measurements are in microns. “N=#” indicates the number of rings in the measured radius. “#=I” indicates the year of the inner ring of the measured radius. “ALT#” indicates the field site (Altamaha) followed by the sample ID and, in cases where more than 1 radius was measured, the radius (A or B). On the top right of the each informational line, information is provided to the computer program for correct file reading.

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26=N      1991=I ALT112A      -3(20F4.0)~
22132667266740281120 50730272533100024531013 81544274707270830545240444025602667
474059417012817371003957
26=N      1991=I ALT112B      -3(16F5.0)~
2397 1843 1704 2640 4592 1680 3784 3609 1023 2011 1501 5123 4075 4705 4682 5138
3914 4052 2561 4232 4147 63171008812555 2110 6604
42=N      1965=I ALT119A      -3(20F4.0)~
7501587 309 8925182874 2611576 2863257 540 696 261 143 7461808 622393315894154
1139129810141018 555102924121247 798 304 226 131 267 174 195 102 264 487 276 802
146 246
42=N      1965=I ALT119B      -3(20F4.0)~
560 796 283 1102417340110892136 266 411 38810601875 8502623385211951740 7771688
231 360 571 522 3171165136421411431 286 325 77 251 115 144 201 391 420 248 621
143 135
49=N      1962=I ALT113A      -3(20F4.0)~
542 536 298 360 565 442 456 657 343 453 190 499 349 319 340 396 277 244 293 253
765 989 724 643 521 637 500 367 439 418 459 493 484 374 452 111 389 0 492 492
492 611 302 468 636 484 359 144 129
49=N      1962=I ALT113B      -3(20F4.0)~
676 526 360 430 579 465 474 588 445 433 300 611 367 389 591 530 521 294 314 259
880 968 626 481 493 514 560 500 363 381 443 470 484 304 416 292 370 79 425 468
536 573 287 489 542 277 274 84 86
51=N      1958=I ALT120A      -3(20F4.0)~
8723111 1302398211411861184 918 714 469 87830214102200012662124 225355130622612
2980 425 212 6731948 54943043739 63915821480 531127942982860 103518823931290 477
1234 5423065212118861839 138 126 780 959 603
56=N      1958=I ALT120B      -3(20F4.0)~
8101698 10626842116 3533822118811291970 53528324307238713971614 79270335642728
3179 505 37110111964 56350964412 932 674 922 595128040992718 15154403485 882 361
1403 180 640 636 689 737 9311513 472 52 78 660 1941816 649 590
62=N      1954=I ALT122A      -3(20F4.0)~
408 828 881 23310253605 753 726 299 198104913472183 37442202173 736 896 7271105
1446157012381402 84313011900 87466133041114 240 758 286 843 272 518175651001458
16761518 457 1981543 617 778 922 32320562733 702 60618671472 1921025127417271422
1566 820
62=N      1954=I ALT122B      -3(20F4.0)~
5121156 811 215 6563800 567 449 312 267114512572456 71144812768 579 745 6111180
18251322 6101172 74913472030 132485338081468 312 961 5381294 186 54817495143 907
19831383 8411284209818212744 981 308 642 7831279 5572616294012792037214120051652
17001109
47=N      1911=I ALT032A      -3(16F5.0)~
6600 97601176012520 55601796022000206802528026390 4762 0 9863 9292 3382 4906
4905 9764 5480 2810 9764 1233 4047 1023 4256 8628 6421 1188 1792 1279 1536 746
4029 6698 1071 1303 1954 0 1629 1033 1329 918 1229 295 180 659 672
47=N      1911=I ALT032B      -3(16F5.0)~
6646 939210489 7183 715911871 8422118531803917888 2240 747411534 6684 7641 2311
2998 5528 3798 2675 6913 961 4985 1561 2489 3524 3434 1007 9365 7300 4009 2417
3044 4279 837 352 1013 0 2887 2868 3157 4493 5150 4625 1005 696 532
87=N      1865=I ALT041A      -3(16F5.0)~

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561 666 323 499 405 588 394 244 277 400 366 472 728 1200 1180 1304
 1397 1345 997 407 318 405 334 318 281 281 1058 9925 1899127461204411849
 851312531 6192 7045 9889 6018 5024 630610314 7471 6884 6075 6246 1922 889 5708
 611 2862 2667 1306 1779 806 1500 4058 564 2000 743 161 435 677 871 839
 129 258 774 177 452 180 2241 584 564 581 436 518 693 564 838 679
 193 661 1019 376 194 308 146
 50=N 1902=I ALT041B -3(16F5.0)~
 5271 4090 5835 7473 5826 5903 4816 8227 5210 425111708 1261 3913 3195 1745 1299
 529 873 2014 635 1749 397 689 159 476 610 503 370 344 1007 476 1775
 0 2246 1086 1167 691 852 1141 1321 1325 1086 1111 582 1376 1545 951 883
 1507 767
 131=N 1878=I ALT121A -3(20F4.0)~
 3366113444844236162816962860174170372725250814452237258110431024 367105910621215
 1151142825161754 9291033 440 604 645 591 956 647 317 40614851195 7081665 240 812
 552416918931052402316121937 9292228 46610491550 957 882 328 8221012 811 310 444
 1151 834 513 59511732297 855 892121716691146 868 955 531 988 860 410 505 506 97
 4732137 5441541 914 488 226 248 492 87 75 38 121 124 54 27 108 75 75 307
 315 360 524 102 833 230 506 103 138 782 908 149 5401581 874 615 676 460 172 69
 448 149 116 564 332 690 696 195 0 934 656
 131=N 1878=I ALT121B -3(20F4.0)~
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 1439423220591426157410612680 8351931 44013662226 993 725 0 471 6381286 240 981
 1843 577 751 54111471800 337 547 872162714881198 688 70 500 500 407 6281000 302
 80321751396 860 824 429 396 330 88 165 99 77 99 55 33 297 143 363 583 184
 0 1541044 300 606 222 825 582 179275326711152 01414 741 405 467 577 248 270
 585 464 833 323 0 833 839 487 6710351228
 148=N 1862=I ALT123A -3(20F4.0)~
 18463127 163 779 218 5211658 481 344 161 650 97810311597 700 89 414 110 785 803
 122 542 75723673142 5473183 91028482751205149083202 9591850 227 116 3661714 615
 1356 968 379 2671404 337155310003921 6111195 296 20329851761 463 21197532230 357
 2675 241264915621236117716084795 3092231 18646222907 701 763123826411855 7221361
 4141218 495 330 207 790 767 312 4142087 5981546 866101122501779 8492357 5781303
 86819621603 680 547 415 321 3211528173512641170 793 585 39713772115 340 567 487
 415 5092211 774 5861074 586 858 7571958 958 614 571 471 429 829 414 40110151514
 1272 886 600 443 157 286 154 173
 148=N 1862=I ALT123B -3(20F4.0)~
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 97 447 58920053166 6773131107925542776183244802446 6531778 567 235 2891888 412
 9731032 367 415 429 336138812523345 5151296 333 30825571665 982 23280342113 296
 2671 447225512751398 97916324705 567 853 178503821181592 815110421641776 729 693
 749 999 413 326 0 689 642 658 2211878 5331275 747113924971422 8632551 6871438
 85318891423 754 549 412 402 3061465168910021176 908 890 842 7003017 325 655 332
 433 4002044 813 422 889 897 724 6022009 813 422 526 538 421 817 567 416 8331286
 1003 866 515 477 202 230 236 225
 123=N 1787=I ALT011 -3(20F4.0)~
 1462183426902215 30938562974 99824092474 332 306271431111495105814971640 5172213
 1071183223152606214818321266 83312291653223310291326 617 357 682 844 584 7141104
 956 780 618 390 390 617 422 910 974 9761201 812 722 747 78012341137 876 6821428
 1363 5851428 910 715 64913961137 68210391169 876 782 312 649 6181006 7151266 454
 1136 357 2201007 782 374 9421594 276 58514961688125316861241 4361251234113643341
 97827052115100149111364136463631206 2272004 24317733546 52310971778 3192159 846
 593 4785651
 77=N 1724=I ALT073 -3(16F5.0)~
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 2758 2124 2478 2002 3793 2039 488810250 5031 5589 1144 2865 3142 2287 3281 3535
 244 1636 1268 4321 2719 931 1343 559 2089 4191 3959 4697 4689 4838 3995 1574
 1382 1717 1857 1762 2670 2714 2148 2239 2908 1907 2762 1527 381 621 571 103
 1000 619 2097 1153 2667 2857 3096 2398 2552 1464 1404 6872 7192
 77=N 1696=I ALT071A -3(20F4.0)~
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 14171002118811891250 417 167 5421041 583 0 771 229 708129310211250 897 9621625
 11041708 646 6891062115020011004206320652918 949 3452584224815561211
 91=N 1688=I ALT071B -3(20F4.0)~
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 14531976152421461502 9791559127010371229135112661635 930 313 7161132 635 0 439
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 21612330274641661794247519382379203622321850
 265=N 1522=I ALT007 -3(16F5.0)~
 387 2097 1448 462 648 722 537 702 465 560 476 124 453 247 408 732
 257 408 184 970 470 412 803 608 525 150 125 535 1096 455 206 475
 2391 1629 341 1405 1307 773 1552 569 1173 1389 845 1794 259 319 475 89
 2216 1681 1199 527 1586 1672 1707 1500 2034 518 631 1879 2853 698 940 382
 2871 1001 517 1103 811 906 1681 1216 629 630 957 491 138 295 1227 495
 1334 851 484 1965 1052 603 310 312 294 225 397 724 225 345 1224 121
 864 1174 1086 1103 794 1484 1052 879 2096 133 1309 1709 1369 945 1612 2521
 1781 1151 1345 1684 242 1575 509 1067 388 570 121 122 558 195 1079 1297
 2824 740 1563 351 2545 1963 401 1066 933 376 764 449 194 558 109 630
 279 146 424 109 685 1182 527 1854 1655 3672 4199 2364 2000 1927 640 177

1401 709 1837 2286 4912 2014 2068 2326 2068 1388 354 231 1769 1126 938 99
 553 2509 3902 4109 1057 1886 227 128 1472 859 741 3388 2054 4474 780 1659
 4109 3388 7743 3654 1245 2845 5946 7613 1788 5383 1771 3723 1991 1844 166 221
 240 627 426 682 719 2581 184 848 1531 224 204 16611097 4590 1714 1365
 221 645 793 1327 350 2157 700 117 130 297 1198 332 442 295 1807 700
 1069 2820 1716 2360 2667 4309 641 4206 3257 4721 662 364 221 785 1622 1097
 1695 369 177 1424 1529 454 219 805 273
 134=N 1509=I ALT029A -3(20F4.0)~
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 929 527 778 929 3011155 8041356 954 653 854 803 8791055 580 779 352115513561531
 7531029 804 904 578 804 251 879 7782059175819331506150614811757 2261282 527
 1052 144 536 4351464 39213201052150711751712101114641052
 113=N 1509=I ALT029B -3(20F4.0)~
 1962 7771060 5851692 818 60015821981 38230044495344043633186551044862913 8204381
 235217851905159717761089 9481305191810761198162215772359185021142502223111001383
 165012111273145714441015214127912152229519142894241824142460 9331741 8871670 552
 7993511258931891397119210172001 9831367 7521336148715841338 4171102 1211242 762
 794 966 483 810 742 897 535 604 103 462 606 655 776 931 777 983 69105310701483
 7081276 8451345129314141310 2081174 709219212421761
 157=N 1386=I ALT072A -3(20F4.0)~
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 228317361460130215531462 842155318451875347520971105314412561861 873163410932577
 75018151645 6611191180428081991166418751144290425404000366536641355209914362077
 2629210731631694 8801017 71212232042167015771004 4171520 1251525 190
 822 3801012 485 470 925 219 791 925127412371299 5761501 22122971299163024772023
 404 718 462 521 577 154 9261062 866 866 848 65611351120 770 250 501 8521116 983
 102114281195 501 19 577 494 811 7021021 310 810 154 8601219 19216781334 8871311
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 2399278317261348206531852207131815521223241279335726372638 8402238 95917092007
 15522335270812851126 645 769 494 24030361015221215151485 4391318 911030 106 803
 288 658 773 30 684 380 576 348 929 9622151 161 742 668 549 789 664 666 333 152
 778 455 587 741 879 8391261 464 221 616 5911027 519 800 982 336 428 510 446 296
 417 725 721 500 760 645 806 5881376 512
 132=N 1282=I ALT060 -3(20F4.0)~
 1627156617941869166914461187 864222117971092 894 8231457 982118710461367 9091048
 751 6861227102424601592 776 802 705 568 136 798 706 010461008109510481092 728
 1393 914 826 325 844105414431645116011751159101910621408 72410631111130414981352
 10621369 966 9661356 3381693 87112551353 966111235741982164513072127136014041642
 12081111159413531404 8711356111110151596 53116961307101910191653 389 971 7441961
 1707 7312042221518692797362918661279117219331137 48312762244 931 866 102 758 655
 862 65512061208 655 527 7591103 897144812211898
 212=N 1249=I ALT095A -3(20F4.0)~
 40334281303221274250271921122752106323263727478938292246129926423470 314 400 613
 555 42 441 457 206 840 413 597 6611282 802 274 275 4961220 99856264569 482 344
 303 648 138 372 179 444 401 419 924 771 496 880 216 411 334 400 808 915 703 869
 110111281032 3721088 924104710971335 76019161515138317072159 67010911707 2481043
 650 821 287 937 596 760 257 213 339 691 164 650 128 557 681 606 399 447 375 469
 532 6701053 106 766 341 287 585 4821246 44710871038 632 128 362 513 663 203 699
 35 725 610 451 384 429 203 746 266 575 500 309 575 532 2051025 713 628 596 504
 139 787 384 128 841 447 543 522 160 15 425 106 319 277 160 479 553 791 351 181
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APPENDIX C

Spatial Distribution of Altamaha WMA Trees

We collected the vast majority of specimens from two locations: 1) trees lying near roadways, which could be accessed in vehicles and 2) trees buried in ponds where large numbers of samples could be obtained. The P.I. and collaborators hiked around the WMA impoundments on several occasions to locate ancient trees and mark them for subsequent sampling. Since our DNR collaborators had limited time to help with the sampling using a 4.5-foot long chainsaw, we selected the ponds in which the largest number of samples had been located for sampling. Northing and Westing for these trees and other (modern) trees recovered from Sapelo Island is provided in Table 3.1 in this dissertation.

Mapping of the trees sampled at the Altamaha WMA (Figure C.1) may indicate that older specimens are generally located in the eastern (downriver) portions of Butler Island. It is unsure at this time whether this is simply a result of sampling happenstance or whether environmental factors (e.g., sea level regression that allowed establishment farther downriver) may play a role in the spatial distribution of ancient trees at the Altamaha WMA. Future sampling of more specimens at the site when they are recovered and located, particularly in as-of-now unsampled ponds, may help to elucidate the spatial distribution of the Altamaha WMA trees.



Figure C.1. Locations of buried recovered trees sampled on Butler and Champney Islands in the Altamaha Wildlife Management Area. There is no clear relationship of spatial distribution with ages of trees, but, in general, the older trees sampled were located farther east (downriver) than the younger trees. The reason for this distribution is at this time unknown.